

甘肃北山西涧泉子富碱高钾花岗岩体的锆石 LA-ICP-MS 定年及其构造意义

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摘要:位于甘肃北山南带的西涧泉子花岗岩体锆石的 LA-ICP-MS 年龄为 266.1 ± 2.2 Ma, 其 Hf 同位素的模式年龄为 $746 \sim 871$ Ma, $\epsilon_{\text{Hf}}(t) = +1.3 \sim +4.7$, 平均 $+2.7$ 。西涧泉子花岗岩体以二长花岗岩为主, 具有较高的 SiO_2 含量 ($74.33\% \sim 76.67\%$) 和 $\text{Na}_2\text{O} + \text{K}_2\text{O}$ 含量 ($9.26\% \sim 9.57\%$), $\text{NK/A} = 0.9 \sim 1.08$, 富钾 ($\text{Na}_2\text{O}/\text{K}_2\text{O} < 1$), 属碱性系列。球粒陨石标准化分配模式图整体呈 V 字型, 具相对平缓的富集轻稀土元素稀土配分模式, 重稀土元素分馏不明显且相对亏损, Eu 具有明显的负异常, $\delta\text{Eu} = 0.10 \sim 0.16$ 。在微量元素原始地幔标准化蛛网图上, 亏损 $\text{Nb}, \text{Ba}, \text{Sr}, \text{Ti}$, 富集 $\text{Rb}, \text{Th}, \text{K}$ 。结合区域地质背景, 认为西涧泉子富碱高钾花岗岩体受音凹峡二叠纪裂谷带的影响强烈, 为壳幔混合成因, 是后碰撞环境裂谷作用阶段的产物, 进而确定该地区在中二叠世进入裂谷作用阶段。

关键词:北山, 富碱高钾花岗岩, 裂谷作用, 锆石 U-Pb 年龄, Hf 同位素

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LA-ICP-MS zircon U-Pb ages of Xijianquanzi alkali-rich potassium-high granites in Beishan, Gansu Province, and their tectonic significance

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Abstract: Beishan area is located in the west of Inner Mongolia and Gansu Province, which is the conjunction zone of Tarim plate, Sino-Korea plate and Kazakhstan plate and plays an important role in the study of the Central Asia Orogenic Belt (CAOB). The relationship of the three plates is so complicated that there exists much controversy concerning Paleozoic tectonic evolution in this area. Granitoids, especially Late Paleozoic granitoids, are widely distributed in Beishan area, which are of great importance in the study of Paleozoic magmatic evolution. Through analyzing the spatial and temporal distribution of the Late Paleozoic granitoids, we can find the relationships between these granitoids and explore the crust-mantle interaction in Beishan area during Late Paleozoic, which is important for determining the geotectonic evolution of this area. In the southern belt of Beishan, the most important tectonic event in Permian is the formation of Yin'aoxia rift zone, which makes the study of Late Paleozoic evolution more complex. The results of studying Xijianquanzi granite lying in the southern belt of Beishan in such aspects as major elements, trace elements, REE and LA-ICP-MS zircon U-Pb geochronology have provided new evidence for regional tectonics. The LA-ICP-MS zircon U-Pb age of Xijianquanzi granite is 266.1 ± 2.2 Ma, and the Hf modal ages are $746 \sim 871$ Ma. $\epsilon_{\text{Hf}}(t)$ values are $+1.3 \sim +4.7$ with an average

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of +2.7, which implies the mixture of crustal and mantle-derived magmas. This feature coincides with that of other areas in the Central Asia Orogenic Belt (CAOB). Therefore, the granites in the belt may have similar magma sources. The granitic body mainly consists of monzonitic granites which belong to alkaline series with high content of SiO_2 (74.33% ~ 76.67%), $\text{Na}_2\text{O} + \text{K}_2\text{O}$ (9.26% ~ 9.57%, $\text{NK/A} = 0.9 \sim 1.08$) and K_2O ($\text{Na}_2\text{O}/\text{K}_2\text{O} < 1$). In addition, chondrite-normalized REE patterns and primitive mantle normalized geochemical patterns of all samples are approximately parallel to each other, which shows that they evolved from the same magma source. Chondrite-normalized REE patterns of the intrusive body are of "V" types and show enrichment of light rare earth elements (LREE) with low ratio of $(\text{La/Yb})_{\text{N}}$ (3.18~5.29) and negative Eu anomalies ($\delta\text{Eu} = 0.10 \sim 0.16$). Except for Eu, the abundance of rare earth elements are demonstrably over 10 times that of chondrite. Depletion of Nb, Ba, Sr, Ti, P and enrichment of Rb, Th, K can be seen in the spidergram. Based on regional geological and geochemical characteristics, the authors have reached the conclusion that the Xijianquanzi intrusive body is an alkali-rich potassium-high granite which is a mixed product of crustal and mantle-derived magmas and was formed at the stage of rifting in the post-collisional extensional period, affected greatly by Yin'aoxia rift zone. This implies that the area entered into the stage of rifting in Middle Permian.

Key words: Beishan; alkali-rich potassium-high granites; rifting; zircon U-Pb age; Hf isotope

北山地区位于甘肃、新疆和内蒙古的交界地区，西邻天山和阿尔泰山，东接阿拉善，是华北板块、塔里木板块和哈萨克斯坦板块陆缘结合部，构造过程极其复杂，其大地构造分区一直存在争议（邓自华，1965；刘雪亚等，1995；左国朝等，1990a, 1990b；张新虎，1993；龚全胜等，2003；左国朝等，2003；江思宏，2004）。北山地区花岗岩类岩石分布广泛，约占该区全部侵入岩的95%（聂凤军等，2002a），尽管前人对本区的花岗岩进行了大量的研究，积累了较为丰富的资料（刘雪亚，1984；左国朝等，1990a；穆治国等，1992, 1994；孙桂英等，1995；于海峰等，1999；修群业，1999；梅华林等，1999；许保良等，2001；聂凤军等，2002a, 2002b；江思宏等，2003a, 2003b, 2006a, 2006b；戴霜等，2003；赵泽辉等，2007；王立社等，2009），但具有精确定年的花岗岩体多集中在北山北带，而北山南带精确定年的花岗岩体却很少，目前在北山南带尚未见有石炭—二叠纪花岗岩的精确定年报道。根据区域构造的对比，在北山南带应该存在一条晚古生代缝合带，但前人的研究表明，北山地区的3条主要蛇绿岩带——洗肠井-牛圈子-红柳河蛇绿岩带、明水-石板井-小黄山蛇绿岩带以及红石山蛇绿岩带都不属于北山南带晚古生代蛇绿岩带（左国朝等，1990a；于福生等，2006；任秉琛等，2001；郭召杰等，2006；张元元等，2008；代文军等，2008；王立社等，2007；宋泰忠等，2008；黄增保等，2006）。另外，北山南带二叠纪最重要的事件为音凹峡二叠纪裂谷带

的形成（左国朝等，1990a；刘明强等，1999；龚全胜等，2002；何世平等，2002）。该裂谷带的形成对整个北山南带的叠加影响使该区的古生代构造演化研究更为困难。因此，对北山南带花岗岩体的成因、构造环境背景及形成时代等问题的研究，不仅对阐明北山南带的大地构造归属有重要意义，而且对古生代的构造演化及东西构造带的对比显得尤为重要。西涧泉子花岗岩体位于北山南带音凹峡二叠纪裂谷带南侧，本文通过锆石LA-ICP-MS定年和地球化学分析研究，对该岩体的形成年代、成因机制及构造意义进行探讨，为该地区晚古生代的构造演化研究提供一些依据。

1 区域地质背景和岩体地质特征

西涧泉子花岗岩体位于甘肃省肃北县西涧泉西北9 km处（图1）处在北山南带音凹峡二叠纪裂谷带南侧10 km范围内，属于塔里木板块的北缘（左国朝等，1990a；张新虎，1993；刘雪亚等，1995；刘明强等，1999；龚全胜等，2002；何世平等，2002；聂凤军等，2002a）。该岩体侵入于前长城系敦煌岩群中，呈岩株状产出，出露面积约为2 km²，为肉红色中细粒二长花岗岩，无岩相分带。1:50 000区域地质填图中获得红柳沟北单元雷洞子岩体的U-Pb同位素年龄值为260.1±0.3 Ma。敦煌岩群Ⅲ岩组为中基性火山岩建造，含碎屑岩夹层；敦煌岩群Ⅳ岩组为碎屑岩建造。

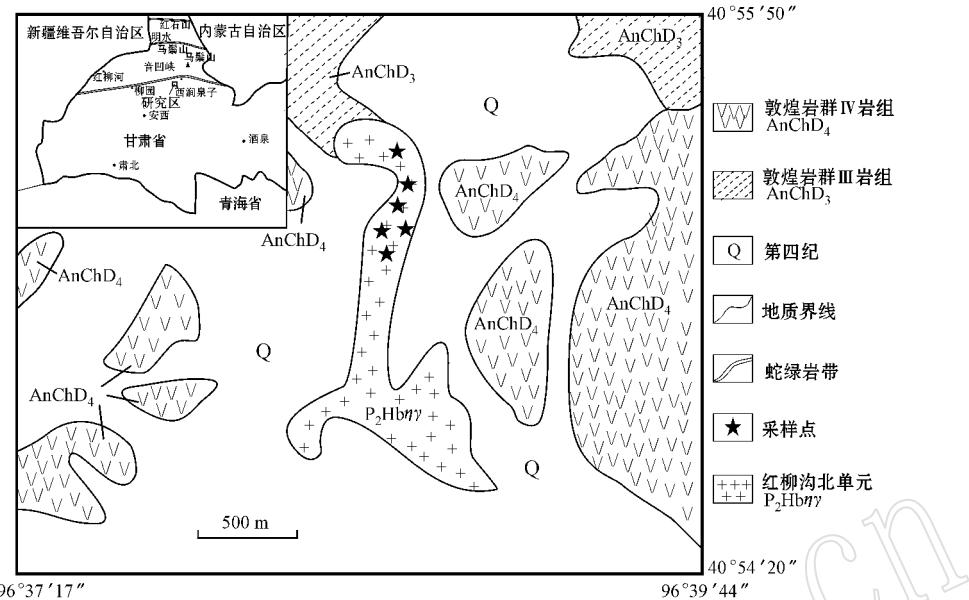


图 1 北山西涧泉子岩体地质简图(据 1:5 万西涧泉幅地质图^①修改)

Fig. 1 Geological sketch map of Xijianquanzi granite in Beishan (modified from 1:50 000 Geological Map of Xijianquan Sheet^①)

北山南带二叠纪地层主要出露有下二叠统双堡塘组、中二叠统金塔组和上二叠统方山口组。双堡塘组(P_{1sh})分布范围很广,多呈条带状分布于双堡塘、煤窑西山、后红泉南山、红柳园西、双井子、牛圈子和马鬃山等地,为一套以海相碎屑岩为主、夹灰岩和砂质灰岩扁豆体的岩石组合,与下伏干泉组、石板山组及海西中期花岗岩等为不整合接触,与上覆金塔组(P_{2j})分布于北山南带双堡塘、俞井子、煤窑西山、后红泉南山、红柳园和大沙沟等地,为一套巨厚的海相基性火山岩类夹正常碎屑岩组合,底界与双堡塘组整合接触,顶界被侏罗系不整合覆盖。方山口组(P_{3f})主要分布于红柳园地层小区白尖山南坡、安西县大奇山、野马井南和民勤黑山等地,由一套中酸性火山熔岩及火山碎屑沉积岩组成,局部可见少量的基性火山岩,底部有一层成分复杂的砾岩,与下伏石板山组不整合接触。二叠系海相碎屑岩-碳酸盐岩-火山岩组合构成一个完整的伸展型裂陷盆地的沉积序列(王洪亮等,2005)。

西涧泉子岩体的镜下特征:石英(30%~45%),颗粒相对较大,半自形到他形,可见石英粒化现象、波状消光;长石颗粒相对石英来说较小,斜长石与碱性长石含量相近;斜长石(20%~30%),

多见环带结构和聚片双晶,部分发生蚀变,多为绢云母化;碱性长石(20%~30%),以条纹长石和微斜长石为主,两种长石含量相近,也有少量钾长石,基本未蚀变或蚀变较轻;云母(4%~6%),以黑云母为主,含少量的白云母,多充填在矿物颗粒之间,少部分发生蚀变形成绿泥石;另外,含少量的帘石,以绿帘石为主,也有黝帘石,多见在黑云母中;副矿物有榍石、锆石、磷灰石及磁铁矿。在 QAP 图解(图略)中,落入二长花岗岩区域。

2 LA-ICP-MS 年代学

选取样品 BS07-35 作为锆石激光探针年龄分析对象,通过 CL 和 BSE 图像分析,选择测点位置,然后采用激光剥蚀等离子体分析技术(LA-ICP-MS)对锆石进行微区原位单点 U-Pb 同位素定年。锆石 U-Pb 年龄测定和微量元素分析在中国科学院地质与地球物理研究所多接收等离子体质谱实验室进行。实验中采用 He 作为剥蚀物质的载气,用美国国家标准技术研究院研制的人工合成硅酸盐标准参考物质 NIST SRM 610 进行仪器最佳化,使仪器达到最高的灵敏度、最小的氧化物产率、最低的背景值和稳定的信号。锆石年龄测定采用外标校正方法,每隔 5 个

^① 甘肃地矿局酒泉地调队. 1994. 西涧泉幅地质图及报告(1:5 万)

样品分析点测一次标准,以保证标准和样品的仪器条件完全一致。在分析样品的前后各测 2 次 NIST SRM 610,以 Si 作内标来测定锆石中 U、Th、Pb 的含量。采用国际标准锆石 91500 作为外标标准物质。详细的实验原理及流程见参考文献(Yuan *et al.*, 2004),样品的同位素比值及元素含量计算采用 GLITTER(ver 4.0; Macquarie University)程序,并应用

Andersen(2002)的 $^{207}\text{Pb}/^{206}\text{Pb}$ 、 $^{207}\text{Pb}/^{235}\text{U}$ 、 $^{206}\text{Pb}/^{238}\text{U}$ 和 $^{208}\text{Pb}/^{232}\text{Th}$ 综合方法进行同位素比值的校正,以达到扣除普通 Pb 的影响。放射成因 Pb 的含量根据 Stacey 和 Kramer(1975)所给出的地壳平均 Pb 同位素组成进行了普通 Pb 校正。年龄计算及谐和图的绘制采用 Isoplo(ver 3.23)程序完成。同位素比值和年龄误差为 1(表 1)。

表 1 西涧泉子花岗岩 BS07-35 中的锆石 LA-ICP-MS U-Pb 分析数据

Table 1 LA-ICP-MS U-Pb dating data of granite BS07-35 in northern Xijianquanzi

分析点	$w_{\text{B}}/10^{-6}$		$^{232}\text{Th}/$ ^{238}U	$^{207}\text{Pb}/^{235}\text{U}$		$^{206}\text{Pb}/^{238}\text{U}$		$^{207}\text{Pb}/^{235}\text{U}$		$^{206}\text{Pb}/^{238}\text{U}$	
	^{232}Th	^{238}U		比值	1σ	比值	1σ	t/Ma	1σ	t/Ma	1σ
121-01	481.98	345.61	1.395	0.327 9	0.027 0	0.042 3	0.001 3	288	21	267	8
121-02	251.98	688.61	0.366	0.324 5	0.005 7	0.042 2	0.000 4	285	4	266	3
121-03	602.31	1 477.34	0.408	0.300 4	0.006 9	0.041 7	0.000 4	267	5	263	2
121-04	100.07	178.38	0.561	0.338 7	0.010 5	0.043 6	0.000 6	296	8	275	4
121-05	12 212.35	4 956.28	2.464	1.086 3	0.062 4	0.027 4	0.000 6	747	30	174	4
121-06	114.72	291.58	0.393	0.300 0	0.010 9	0.042 1	0.000 5	266	9	266	3
121-07	162	708.45	0.229	0.310 7	0.013 1	0.042 3	0.000 7	275	10	267	4
121-08	222	645.38	0.344	0.341 0	0.005 4	0.044 9	0.000 4	298	4	283	3
121-09	12 967.79	13 218.17	0.981	0.761 0	0.035 9	0.032 4	0.0005	575	21	206	3
121-10	7 562.37	4 281.57	1.766	1.003 3	0.009 4	0.041 6	0.000 4	705	5	263	3
121-11	298.71	305.9	0.976	0.325 0	0.030 7	0.042 1	0.001 4	286	24	266	9
121-12	985.58	1 920	0.513	0.421 9	0.010 6	0.040 0	0.000 4	357	8	253	3
121-13	840.02	1 476.44	0.569	0.323 0	0.016 2	0.042 6	0.000 8	284	12	269	5
121-14	136.87	276.01	0.496	0.346 2	0.024 5	0.045 8	0.001 2	302	18	289	7
121-15	1 639.29	1 304.84	1.256	0.590 7	0.020 1	0.040 6	0.000 5	471	13	256	3
121-16	353.51	326.57	1.082	0.299 0	0.031 6	0.041 5	0.001 5	266	25	262	9
121-17	155.49	443	0.351	0.396 1	0.058 4	0.040 3	0.001 7	339	42	254	11
121-18	374.35	464.05	0.807	0.303 2	0.015 2	0.038 8	0.000 8	269	12	245	5
121-19	100.57	384.93	0.261	0.325 5	0.050 1	0.043 0	0.002 3	286	38	271	14
121-20	661.7	1 757.87	0.376	0.328 9	0.004 7	0.042 3	0.000 4	289	4	267	3

本次实验共选取 20 个锆石,共有 20 个测点。所测的锆石都是透明柱状的,晶形较好,发育良好的岩浆型韵律环带结构,为典型的岩浆成因锆石(图 2)。打点位置边部和核部均有分布,测试过程中未发现较老的锆石颗粒或核部的残留锆石。分析结果见表 1。其中 l21-05、l21-09、l21-10、l21-12、l21-15、l21-17 的谐和度均大于 110,不谐和,将这些年龄删去不用。其余 14 颗锆石的 U、Th 含量分别为 $178.38 \times 10^{-6} \sim 1 757.87 \times 10^{-6}$ 和 $100.07 \times 10^{-6} \sim 840.02 \times 10^{-6}$,Th/U 比值均大于 0.1(0.229~1.395),Th、U 之间正相关性较好,其特征应属典型的岩浆型锆石(Claesson *et al.*, 2000; Belousova *et al.*, 2002)。14 颗锆石的测年结果中,年龄集中在 $245 \pm 5 \sim 289 \pm 7 \text{ Ma}$ 之间,处理数据时,剔除 3 个离群值(l21-08、l21-14、l21-18),对谐和线上最集中的

11 个数据点进行 $^{206}\text{Pb}/^{238}\text{U}$ 加权平均计算,得到该花岗岩的年龄为 $266.1 \pm 2.2 \text{ Ma}$ (95% conf., MSWD=0.82, 1σ)(图 3),代表岩体的结晶年龄。

3 地球化学特征

3.1 主量元素特征

西涧泉子花岗岩体的主量元素分析结果列于表 2。SiO₂% 含量很高,在 74.33%~76.67%;Al₂O₃ 含量在 11.66%~14.02%,A/NCK=0.88~1.08,在铝饱和指数判别图解上落在分界线上(图 4);Na₂O+K₂O 含量高,平均 9.46%,富钾(Na₂O=3.69%~4.49%,K₂O=4.83%~5.57%,Na₂O/K₂O<1);NK/A=0.90~1.08 均大于 0.9,在 SiO₂-AR 与碱度关系图解(图 5)上,样品均落入碱性系列区域。另

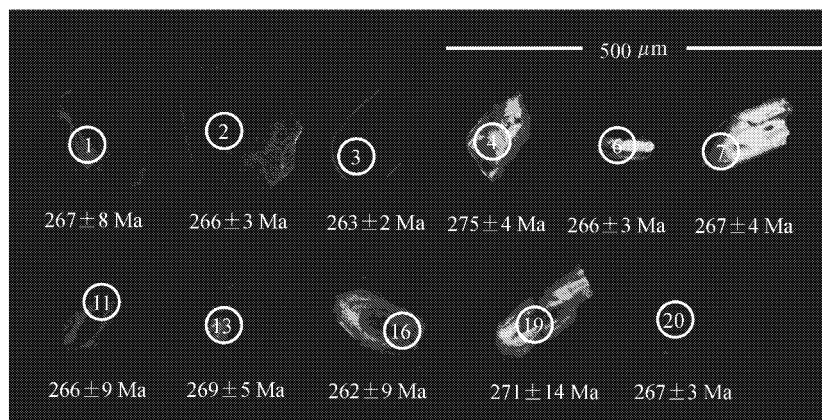


图 2 西涧泉子花岗岩体锆石特征和 LA-ICP-MS 测点位置

Fig. 2 Characteristics and dating spots of zircons from Xijianquanzi granite

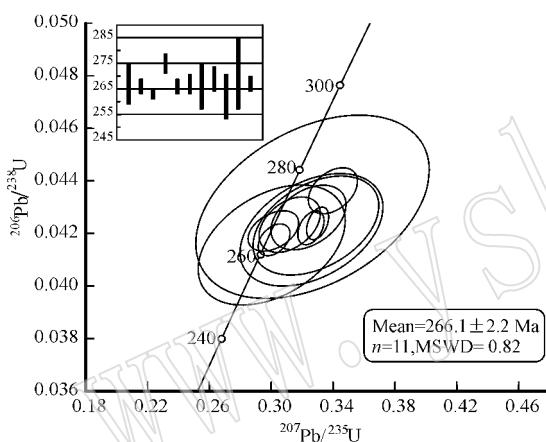


图 3 西涧泉子花岗岩体锆石 LA-ICP MS U-Pb 同位素谱和图

Fig. 3 Zircon LA-ICP MS U-Pb isotopic concordia plots for zircons from Xijianquanzi granite

外 P_2O_5 含量低且变化不大($0.001\% \sim 0.011\%$)；低 TiO_2 ($0.042 \sim 0.059$)；低铁(FeO 为 $0.30\% \sim 0.73\%$)，低 CaO ($0.25\% \sim 0.82\%$)，低 MgO ($0.03\% \sim 0.68\%$)。 $\text{Mg}^{\#}$ 值的变化较大， $\text{Mg}^{\#} = 7 \sim 77$ 。从主量元素的特征看，西涧泉子花岗岩具有富碱高钾岩石的特征。

3.2 稀土与微量元素特征

由表 2 可以看出西涧泉子花岗岩体的稀土元素总量相对较低，为 $175.34 \times 10^{-6} \sim 199.17 \times 10^{-6}$ ，在球粒陨石标准化分配模式图(图 6)上，富集轻稀土元素(LREE)，重稀土元素(HREE)分馏不明显且相对亏损，但都明显高于 10 倍球粒陨石丰度，因此具有相对较低的(La/Yb)_N 比值($3.18 \sim 5.29$)，Eu 具

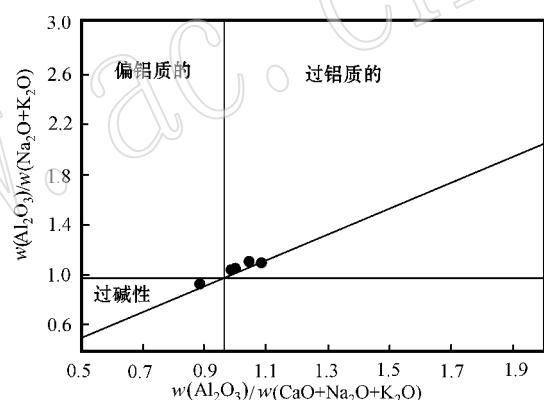
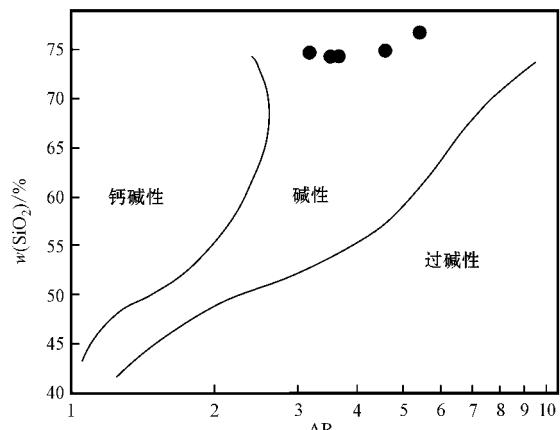
图 4 西涧泉子花岗岩体的铝饱和指数判别图解
(据 Maniar & Piccoli , 1989)Fig. 4 Shand's index diagram of Xijianquanzi granite
(after Maniar & Piccoli , 1989)图 5 AR - SiO_2 与碱度关系图(据 Wright , 1969)Fig. 5 Alkalinity ratio versus SiO_2 diagram
(after Wright , 1969)

表2 西涧泉子花岗岩体的主量元素($w_B/\%$)和微量元素($w_B/10^{-6}$)分析结果

Table 2 Analytical results of major elements ($w_B/\%$) and trace elements ($w_B/10^{-6}$) in Xijianquanzi granite

样品号	BS07-36	BS07-37	BS07-38	AB09-104	AB09-105
SiO ₂	74.35	74.91	74.33	74.71	76.67
Al ₂ O ₃	13.27	13.18	14.02	13.40	11.66
TFe ₂ O ₃	0.71	0.71	0.59	0.30	0.73
CaO	0.77	0.82	0.25	0.78	0.63
MgO	0.31	0.26	0.68	0.50	0.03
K ₂ O	5.67	4.83	5.55	5.57	5.27
Na ₂ O	3.92	4.49	4.07	3.69	4.22
MnO	0.041	0.039	0.021	0.040	0.031
TiO ₂	0.054	0.049	0.043	0.042	0.059
P ₂ O ₅	0.011	0.009	0.001	0.013	0.009
LOI	0.85	0.67	0.41	0.91	0.63
Total	99.96	99.97	99.96	99.96	99.94
AR	3.53	4.59	3.66	3.17	5.38
NK/A	0.95	0.96	0.91	0.90	1.08
A/NCK	1.00	0.99	1.08	1.05	0.88
Mg [#]	46	42	69	77	7
Rb	223.7	185.9	226.8	230.5	165.3
Ba	156.7	198.7	133.5	161.5	152.4
Th	28.93	23.04	26.01	23.55	21.59
U	1.87	1.63	1.72	2.10	1.61
Nb	10.36	13.90	10.72	10.50	7.49
Ta	1.088	0.974	0.455	0.987	0.431
Pb	22.77	26.63	3.58	18.72	17.44
Sr	30.8	37.7	56.7	28.4	37.5
Zr	106.8	102.2	106.0	101.4	94.1
Hf	5.234	5.283	5.153	5.242	4.805
Li	9.079	7.975	9.332	6.905	5.462
Be	2.357	2.581	1.912	2.806	2.059
Sc	4.51	5.07	5.05	5.35	4.67
V	4.73	40.55	4.40	5.29	7.50
Co	4.36	57.77	4.21	3.48	3.39
Ga	15.61	17.06	15.20	15.81	13.31
Cs	156.7	198.7	133.5	161.5	152.4
La	35.10	32.33	33.27	33.24	31.59
Ce	78.83	71.89	71.45	71.73	73.11
Pr	8.817	7.889	8.561	8.602	8.053
Nd	34.39	30.44	37.27	35.16	32.52
Sm	7.564	6.546	7.557	8.835	6.461
Eu	0.346	0.340	0.271	0.324	0.307
Gd	8.044	6.866	7.713	10.27	6.514
Tb	1.288	1.074	1.169	1.774	0.996
Dy	8.16	6.62	7.00	11.3	6.11
Ho	1.703	1.350	1.357	2.295	1.228
Er	5.259	4.186	4.067	6.856	3.725
Tm	0.797	0.642	0.601	1.012	0.562
Yb	5.427	4.490	4.067	.766	3.882
Lu	0.815	0.679	0.603	0.988	0.584
Y	48.78	38.90	38.29	65.46	37.56
REE	196.54	175.34	184.95	199.17	175.64
LREE	165.05	149.44	158.38	157.89	152.04
HREE	31.49	25.90	26.57	41.28	23.60
LREE/HREE	5.24	5.77	5.96	3.83	6.44
(La/Yb) _N	4.18	4.66	5.29	3.18	5.27
δ Eu	0.14	0.16	0.11	0.10	0.15
10 000 Ga/Al	2.22	2.45	2.05	2.23	2.16

测试由北京大学造山带与地壳演化教育部重点实验室完成 δ Eu =

$$\text{Eu}_{\text{N}} / (\text{Sm}_{\text{N}} \times \text{Gd}_{\text{N}})^{1/2}$$

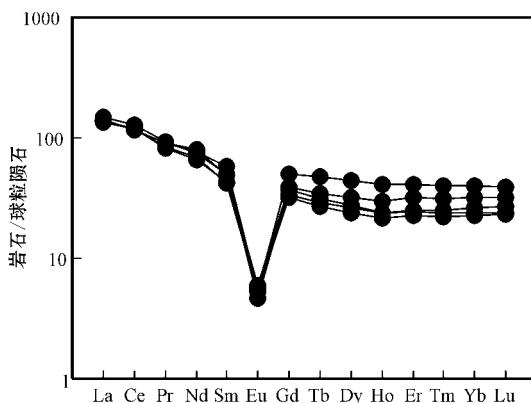


图6 稀土元素球粒陨石标准化图解(标准化值据 Sun & McDonough, 1989)

Fig. 6 Chondrite-normalized REE patterns
(normalized values after Sun & McDonough, 1989)

有很明显的负异常 δ Eu = 0.10~0.16, 可能与斜长石在源区的残余有关。稀土元素球粒陨石标准化分配模式整体呈V字型。

在微量元素原始地幔标准化蛛网图(图7)上,所有样品的稀土模式图和原始地幔标准化图基本平行,显示了同源演化的特征。Nb、Ba、Sr、Ti、P出现较大的亏损, Rb、Th、K具有明显的正异常,其中Nb元素明显亏损反映了伸展环境(Brown et al., 2003)。在花岗岩 Sr-Yb 类型判别图解(张旗等, 2006a, 2006b)上投入非常低 Sr 高 Y 区(图略)张旗等(2006a, 2006b, 2008)认为具有非常低 Sr 高 Yb 型花岗岩形成的压力极低, 在相图上大多位于高温

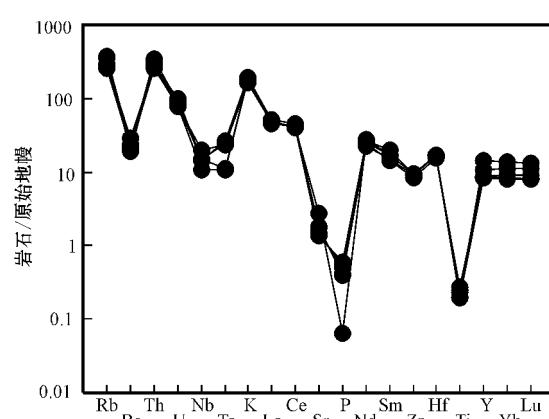


图7 微量元素原始地幔标准化图解(标准化值据 Sun & McDonough, 1989)

Fig. 7 Trace element spider diagram (normalized values after Sun & McDonough, 1989)

低压区,代表地壳减薄的产物。在 Rb-Y+Nb 图解(图 8)上,所有样品均落入了后碰撞范围,而且可以看出所有样品均具有过渡型花岗岩的特征(吴泰然,1995)。在样品与典型的后碰撞花岗岩大洋中脊标准化图解(图 9)上,样品的趋势与典型的后碰撞花岗岩并不一致,也说明西涧泉子花岗岩并非为典型的后碰撞花岗岩,可能为后碰撞晚期伸展阶段的产物。

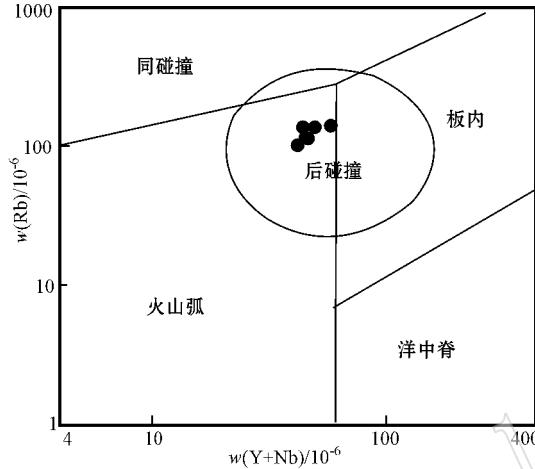


图 8 Rb-Y+Nb 图解(据 Pearce, 1996)

Fig. 8 Diagram of Rb-Y+Nb (after Pearce, 1996)

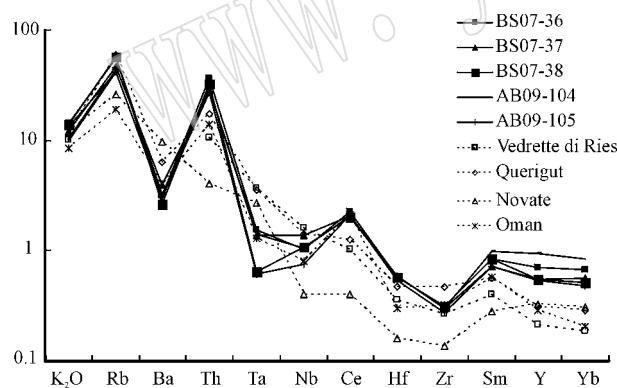


图 9 样品与典型的后碰撞花岗岩大洋中脊标准化图解
(典型花岗岩数据和标准化值据 Pearce, 1984)

Fig. 9 Ocean ridge granite (ORG) normalized geochemical patterns for samples and representative post-collision granites (representative post-collision granites; data and normalized values after Pearce, 1984)

4 锆石 Hf 同位素特征

参照锆石阴极发光(CL)图像进行了锆石 Hf 同位素原位测量实验。实验在中国科学院地质与地球

物理所配有 193 nm 激光取样系统的 Neptune 多接收电感耦合等离子体质谱(LA-MC-ICP MS)上进行,激光束直径为 63 μm,激光剥蚀时间为 26 s,激光脉冲频率 6 Hz,激光脉冲能量 100 mJ,测定时采用锆石国际标样 91500 做外标。详细的分析流程见 Wu 等(2006)。测定结果列于表 3,并表示于图 10。可以看出西涧泉子花岗岩的 $\text{Hf}^{176}/\text{Hf}^{177}$ = 0.282 643~0.282 731 相当于 $\epsilon\text{Hf}(t) = +1.3 \sim +4.7$,平均 +2.7,均为正值, Hf 同位素的模式年龄为 746~871 Ma。

5 岩体成因及其形成的构造背景

西涧泉子花岗岩高硅,富碱高钾,具很低的 CaO、MgO 稀土元素球粒陨石标准化分配模式整体呈 V 字型,这些特点类似于 A 型花岗岩,但该花岗岩体主要为二长花岗岩,且 $10000 \text{ Ga/Al} = 2.05 \sim 2.45$,在 Whalen 等(1987)的区分 A 型与 I、S 型花岗岩的判别图解(图略)中,并不能投入 A 型花岗岩区域,故推测为一种向 A 型花岗岩过渡的富碱高钾花岗岩。

西涧泉子富碱高钾花岗岩的年龄为 266.1 ± 2.2 Ma。据目前的区域地质研究,北山地区在早二叠世之前洋壳已经闭合,之后进入碰撞阶段(左国朝等,1990a;刘雪亚等,1995;龚全胜等,2003;何世平等,2005)。很多学者认为北山地区在二叠纪处于一种伸展拉伸的构造背景中(左国朝等,1990a;姜常义等,2006;赵泽辉等,2006)。北山南带音凹峡二叠纪裂谷带就是在该背景下形成的(左国朝等,1990a;刘明强等,1999;龚全胜等,2002;何世平等,2002),该带发现有同期形成的二叠纪花岗岩与海相玄武岩,同时发育大量的辉绿岩脉(左国朝等,1990a;赵泽辉等,2006;姜常义等,2006)。由于 Pearce 等的一系列花岗岩构造环境判别中都没有裂谷这一环境,因此这类花岗岩在判别图解中均只能判断为后碰撞环境形成的。而后碰撞环境是一个复杂的时期,该时期包括了诸如板块之间沿剪切带的大规模运动、合拢、岩石圈拆沉作用、小型海洋板块的俯冲以及裂谷的生成等,这些事件包括持续的或间隙的伸展体系(肖庆辉等,2002)。综合区域地质背景,可以判断 266.1 ± 2.2 Ma 的西涧泉子富碱高钾花岗岩是后碰撞裂谷作用阶段的产物。吴泰然(1995)将这类花岗岩归为拉张型过渡壳花岗岩(ECG),认为这种花岗岩的形成是由于地壳的拉伸

表3 甘肃北山地区西涧泉子花岗岩体LA-MC-ICPMS锆石Hf同位素分析结果

Table 3 Zircon Hf isotopic compositions of Xijianquanzi granite in Beishan area

点位	$^{176}\text{Yb}/^{177}\text{Hf}$	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}_{\text{m}}$	$^{176}\text{Hf}/^{177}\text{Hf}$	2σ	$\epsilon\text{Hf}(t)$	$\epsilon\text{Hf}(0)$	t_{DM}/Ma	$f_{\text{Lu/Hf}}$
121-01	0.041 551	0.001 625	0.282 676	0.282 669	0.000 020	-3.4	1.5	830	-0.95
121-02	0.032 579	0.001 300	0.282 689	0.282 684	0.000 014	-2.9	2.2	804	-0.96
121-03	0.047 961	0.001 880	0.282 652	0.282 643	0.000 015	-4.2	1.3	871	-0.94
121-04	0.031 016	0.001 235	0.282 664	0.282 658	0.000 018	-3.8	2.1	838	-0.96
121-05	0.121 574	0.004 165	0.282 711	0.282 698	0.000 015	-2.1	1.0	837	-0.87
121-06	0.051 875	0.002 045	0.282 741	0.282 731	0.000 019	-1.1	4.2	746	-0.94
121-07	0.035 121	0.001 381	0.282 690	0.282 682	0.000 020	-2.9	2.9	806	-0.96
121-08	0.023 572	0.000 946	0.282 675	0.282 670	0.000 014	-3.4	2.7	817	-0.97
121-09	0.208 912	0.007 764	0.282 745	0.282 720	0.000 012	-1.0	1.9	878	-0.77
121-10	0.084 654	0.003 032	0.282 701	0.282 688	0.000 015	-2.5	2.0	826	-0.91
121-11	0.036 926	0.001 399	0.282 672	0.282 665	0.000 015	-3.5	2.2	832	-0.96
121-12	0.056 114	0.002 094	0.282 704	0.282 695	0.000 012	-2.4	2.3	800	-0.94
121-13	0.047 677	0.001 803	0.282 701	0.282 692	0.000 018	-2.5	3.0	799	-0.95
121-14	0.035 607	0.001 380	0.282 710	0.282 703	0.000 020	-2.2	3.8	777	-0.96
121-15	0.112 262	0.003 461	0.282 743	0.282 724	0.000 020	-1.0	4.7	773	-0.90
121-16	0.032 357	0.001 249	0.282 685	0.282 679	0.000 017	-3.1	2.4	809	-0.96
121-17	0.106 319	0.003 268	0.282 722	0.282 706	0.000 020	-1.8	3.5	799	-0.90
121-18	0.068 804	0.002 441	0.282 726	0.282 715	0.000 017	+1.6	3.3	776	-0.93
121-19	0.032 454	0.001 281	0.282 710	0.282 704	0.000 021	-2.2	3.3	774	-0.96
121-20	0.068 628	0.002 496	0.282 718	0.282 705	0.000 017	-1.9	3.4	789	-0.92

表中锆石Hf同位素成分计算所用的参数： ^{176}Lu 衰变常数 $\lambda = 1.865 \times 10^{-11}$ ，球粒陨石 $^{176}\text{Lu}/^{177}\text{Hf} = 0.0332$ ， $^{176}\text{Hf}/^{177}\text{Hf} = 0.282772$ (Blichert-Toft and Alberade, 1997)；亏损地幔 $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ ， $^{176}\text{Hf}/^{177}\text{Hf} = 0.28325$ (Griffin et al., 2000)；镁铁质和硅铝质地壳的 $f_{\text{Lu/Hf}}$ 分别是 -0.34 (Amelin et al., 2000) 和 -0.72 (Vervoort et al., 1996)。

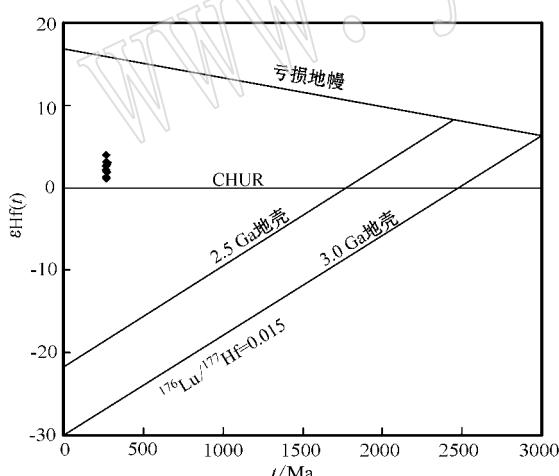


图10 西涧泉子花岗岩锆石的Hf同位素数据

Fig. 10 In-situ zircon Hf isotopic data of Xijianquanzi granite

减薄,上地幔热物质上涌,使地壳形成一种高温低压的环境,并使地壳发生部分重熔,同时上地幔上涌的热物质沿着拉张的裂隙与地壳的热物质发生混染作用,对陆壳的物质进行改造,使之向过渡类型转化,形成拉张型过渡壳的花岗岩。

西涧泉子富碱高钾花岗岩体锆石 $\epsilon\text{Hf}(t) =$

+1.3~+4.7,平均+2.7,均为正值。一般认为具有正 $\epsilon\text{Hf}(t)$ 值的花岗质岩石来自亏损地幔或从亏损地幔中新增生的年轻地壳物质的部分熔融。结合区域地质背景可以得出,音凹峡二叠纪裂谷带形成时有新的地幔岩浆加入,而西涧泉子地区明显受到音凹峡二叠纪裂谷带影响,形成了具有亏损地幔性质的 266.1 ± 2.2 Ma 的西涧泉子富碱高钾花岗岩体。同时,Nb、Ta、Ti 的负异常也代表源区有金红石残留,来源较深。另外,大部分样品的 Mg 值较大,反映幔源岩浆受到的混染程度降低,岩浆分异过程的时间相对较短(吴泰然等,2007),部分样品的 Mg 值较小,可能是由于壳幔混合不均一造成的。综上所述,西涧泉子富碱高钾花岗岩是在裂谷作用下壳幔混合成因的花岗岩,这也同时说明晚古生代甘肃北山地壳为垂向生长。据笔者未发表数据,处于音凹峡二叠纪裂谷带中的音凹峡南黑云母花岗岩所测年龄为 281.7 ± 2.9 Ma, $\epsilon\text{Hf}(t) = +4.4 \sim +7.8$,为一后碰撞环境下形成的壳幔混合成因花岗岩。由此可以看出,北山南带主碰撞以后形成的碰撞后伸展阶段最晚开始于 281.7 ± 2.9 Ma,之后即进入裂谷发育阶段,西涧泉子岩体正是这一过程开始的产物,也

就是说北山南带在 266.1 ± 2.2 Ma 开始了裂谷作用过程。在音凹峡二叠纪裂谷带柳园地区下二叠统中发育侵入年代在 $240 \sim 220$ Ma 的煌斑岩,标志着北山地区二叠纪开始的裂谷作用壳幔过程从初期岩石圈拆沉和亏损地幔强烈活动为主,发展到晚期岩石圈地幔交代富集物质的上涌(刘畅等,2006)。据顾连兴等(2006),北山西邻的东天山海西期碰撞后伸展阶段为 $285 \sim 250$ Ma,由 266.1 ± 2.2 Ma 的西涧泉子富碱高钾花岗岩可以推测北山南带与东天山海西期构造环境可能能够进行对比研究,但尚待进一步确定北山南带其他二叠纪花岗岩体的精确年龄及性质。

在二叠纪的中亚造山带(CAOB),大范围存在具有正 $\epsilon_{Nd}(t)$ 和低的($^{87}\text{Sr}/^{86}\text{Sr}$)、与幔源特征有关的花岗岩的垂向增生(Han et al.,1997;吴福元等,1997,1999;韩宝福等,1999;Wu et al.,2000,2002;Hong et al.,2004;Jahn et al.,2004;Chen and Jahn,2004),而且也存在与西涧泉子富碱高钾花岗岩体、音凹峡南花岗岩类似的例子——二叠纪幔源岩浆活动及同期的花岗质岩浆活动,多数学者认为其是在后碰撞伸展体制下形成的(韩宝福等,1999,2006;顾连兴等,2006),因此西涧泉子富碱高钾花岗岩与音凹峡南花岗岩所代表的北山南带为中亚造山带在二叠纪处于后碰撞伸展环境又提供了一个证据。

6 结论

(1) 西涧泉子花岗岩体主要为富碱高钾二长花岗岩,其岩石化学特征反映了岩体形成于后碰撞裂谷作用环境。

(2) 西涧泉子岩体的 $\epsilon_{Hf}(t)$ 在 $+1.3 \sim +4.7$,具有与中亚造山带其他地区相一致的壳幔混合成因特征,反映了这一花岗岩带具有相似的岩浆源。

(3) 岩体的锆石LA-ICP-MS年龄为 266.1 ± 2.2 Ma,可以确定该地区在早二叠世晚期已进入后碰撞的裂谷伸展发育阶段。

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