

· 专题研究 ·

北祁连造山带窑洞沟地区中奥陶世花岗岩成因

——锆石 U-Pb 年代学和岩石地球化学制约

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摘 要: 对北祁连造山带窑洞沟地区花岗岩进行了 LA-ICP-MS 锆石 U-Pb 测年以及主量和微量元素地球化学分析, 结果表明, 窑洞沟花岗岩的结晶年龄为 466 ± 2 Ma。窑洞沟花岗岩具有较高的 SiO_2 (72.45% ~ 80.18%) 和全碱含量 (Alk) (6.04% ~ 8.12%), 且 K_2O 含量低, $\text{K}_2\text{O}/\text{Na}_2\text{O}$ 值为 0.07 ~ 0.71, 里特曼指数在 0.96 ~ 2.16 之间, 小于 3.3, 铝饱和指数 (A/CNK) 为 0.94 ~ 1.01, 为低钾到中钾钙碱性系列的准铝质岩石; 在球粒陨石标准化的稀土元素配分模式图上表现出轻稀土元素富集的右倾型, 具有中等的负 Eu 异常, δEu 介于 0.67 ~ 0.81 之间, 在原始地幔标准化的微量元素蛛网图上明显富集 Rb、K 等大离子亲石元素和 Th、U 等放射性元素, 亏损 Nb、Ta、Ti、P 等高场强元素, 在 P_2O_5 - SiO_2 协变关系图上与 I 型花岗岩趋势一致, 在 Y - Rb 图上显示出分异 I 型花岗岩的特征。结合区域地质背景, 认为窑洞沟花岗岩的形成可能与北祁连洋盆持续向北俯冲导致弧后拉张伸展, 诱发幔源岩浆岩底侵上升, 致使古老地壳物质重熔形成母岩浆, 再经历中等程度分离结晶作用有关, 其岩浆具有岛弧岩浆属性, 可能是由于弧后盆地在岛弧的基础上裂开或者其岩浆源区受到了俯冲板片流体/熔体的改造作用。

关键词: 锆石 U-Pb 定年; 地球化学; I 型花岗岩; 弧后盆地; 北祁连; 甘肃

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Petrogenesis of the middle Ordovician granite in the Yaodonggou area of the North Qilian orogenic belt: Constraints from zircon U-Pb geochronology and geochemistry

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Abstract: This paper reports LA-ICP-MS zircon U-Pb dating data as well as whole rock major and trace element data from granite pluton in the Yaodonggou area of the North Qilian orogenic belt. The result shows that the pluton was emplaced at 466 ± 2 Ma. The Yaodonggou granite contains high SiO_2 (72.45% ~ 80.18%) and Alk (6.04% ~ 8.12%). Its K_2O content is less than that of Na_2O with $\text{K}_2\text{O}/\text{Na}_2\text{O}$ ratio of 0.07 ~ 0.71, Ritterman index of 0.96 ~ 2.16 (less than 3.3), and A/CNK of 0.94 ~ 1.01. The granite belongs to low-medial potassium calcium alkaline series quasi-aluminum rocks. The chondrite-normalized REE patterns show the right-inclined type of LREE enrichment, with a medium negative Eu anomaly and δEu of 0.67 ~ 0.81. The primitive mantle-normalized spidergram shows depletion of Nb, Ta, Ti, P and enrichment of Rb, K, U. The P_2O_5 - SiO_2 diagram shows the trend of

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I-type granite. The Y-Rb diagram shows the characteristics of differentiated I-type granite. Combined with the regional geological background, the authors hold that the formation of the Yaodonggou granite was probably related to the northward subduction of the North Qilian Ocean Basin, which led to the back-arc extension and mantle-derived magmatic intrusion and uplift. As a result, the ancient crustal material was remelted to form parent magma, then it was subjected to moderate crystallization differentiation. The magma had the property of island arc magma, which might have been attributed to the fact that the back arc basin was split on the basis of island arc or its magma source area was reformed by subduction plate fluid or melt.

Key words: U-Pb zircon dating; geochemistry; I-type granite; back-arc basin; North Qilian; Gansu

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北祁连造山带是一个具有完整沟-弧-盆体系的加里东期造山带(夏林圻等, 1991, 1992, 1995, 1996, 1998, 2001; 冯益民等, 1992; 许志琴等, 1994; 冯益民, 1997; 张建新等, 1997, 1998; 左国朝等, 1997; 张旗等, 2003; 杜远生等, 2004; 徐学义等, 2008), 自20世纪70年代以来, 随着板块构造理论的发展, 围绕火山岩、蛇绿岩、基底构造属性、变质变形作用、洋盆俯冲极性、大地构造演化等多个方面的关键科学问题, 开展了广泛而深入的研究(李春昱等, 1980; Wu *et al.*, 1993; 冯益民等, 1996; 左国朝等, 1997; 葛肖虹等, 1999; 夏林圻等, 1998, 2001, 2016; 杜远生等, 2002, 2004; Xia *et al.*, 2003; Zhang *et al.*, 2007; Song *et al.*, 2007a, 2009a; Xiao *et al.*, 2009), 近年来, 随着高压-超高压变质带的发现, 极大地激发了地质学家对祁连造山带早古生代板块俯冲-碰撞-拼贴过程研究的兴趣, 也取得了许多重要进展(Wu *et al.*, 1993; 许志琴等, 1994; 杨经绥等, 1998; Song *et al.*, 2004a, 2006, 2007a, 2009a, 2013; Zhang *et al.*, 2007, 2012), 但是, 关于北祁连洋盆的俯冲极性存在多种认识: ① 向北俯冲, 主要依据是北祁连造山带自南向北发育洋壳残片、俯冲杂岩带、弧火山岩带和花岗岩带、弧后盆地火山岩和花岗岩带、志留纪到泥盆纪的磨拉石建造以及向南-南西方向逆冲的韧性剪切带(夏林圻等, 1991, 1995, 1996, 1998, 2001; 许志琴等, 1994; 杜远生等, 2002, 2004; Yang *et al.*, 2002; Xu *et al.*, 2006; Song *et al.*, 2006, 2013); ② 向南俯冲, 主要依据南、北深、浅俯冲杂岩可能代表了同一俯冲带不同层次或向北的后退式俯冲, 南侧与中祁连地块结合带发育岛弧火山岩和陆缘弧花岗岩(周德进等, 1997; Gerhrels and Yin, 2003; Gerhrels *et al.*, 2003; Cowgling *et al.*, 2003; Yin *et al.*, 2007; Xiao *et al.*, 2009); ③ 双向俯冲, 主要依据是两边都存在古生代

花岗岩类(左国朝等, 1987; 吴才来等, 2006, 2010; 秦海鹏等, 2014)。其次, 关于北祁连洋盆开始俯冲的时限, 大多数学者认为最早俯冲时间可能为495 Ma或者更早(冯益民, 1997; 张招崇等, 2001; 史仁灯等, 2004; 吴才来等, 2006, 2010; 相振群等, 2007; Tseng *et al.*, 2009; Yan *et al.*, 2010; 夏小洪等, 2010; Song *et al.*, 2013; 夏林圻等, 2016); 也有学者认为俯冲开始时间稍晚于早中奥陶世(林宣慧等, 2010); 再者, 研究区岩浆岩发育, 但目前尚未有相关年代学和岩石地球化学方面的报道, 花岗岩作为了解地球内部的“探针”和“窗口”, 记录着大量构造演化方面的信息(Ma *et al.*, 2004; 莫宣学, 2011)。因此, 本文拟以北祁连窑洞沟地区早古生代花岗岩为研究对象, 在详细的野外调研的基础上, 以系统的锆石 U-Pb 年代学、岩石地球化学为研究手段, 精确厘定窑洞沟花岗岩的形成时代、探讨窑洞沟花岗岩的岩石成因、构造环境及地球动力学意义, 为北祁连早古生代洋盆俯冲机制及构造演化方面的研究提供新的证据。

1 区域地质背景

北祁连造山带处于秦-祁-昆巨型造山系的中段, 挟持于阿拉善地块、塔里木板块与中祁连地块之间, 呈 NW-SE 向展布, 长约 1 200 km, 宽约 100 ~ 300 km, 该造山带北边以龙首山断裂为界, 南界为中祁连北缘断裂, 西端被阿尔金左行走滑断裂截切, 东端为同心-固原右行走滑断裂。自北向南可以划分为弧后盆地、北祁连岛弧、俯冲杂岩和消减洋壳残片等不同的构造单元(冯益民等, 1996)。造山带内花岗质侵入体发育, 主要沿岛链延伸方向出露, 岩石类型主要有花岗闪长岩、花岗岩、二长花岗岩及英云闪长岩等。锆石 U-Pb 年代学表明它们的形成时代多集中

在 512 ~ 402 Ma, 最晚形成于 383 Ma (吴才来等, 2004, 2006, 2010; 秦海鹏, 2012; 熊子良等, 2012), 这些花岗岩体主要以 I 型花岗岩为主, 少量表现出类似于 S 型花岗岩特点, 如柴达诺花岗岩体和民乐窑沟花岗闪长岩 (吴才来等, 2010), 还有一些表现出 A 型花岗岩的特征, 如武威一带 422 ~ 418 Ma 二长花岗岩 (秦海鹏, 2012) 和黄羊河 404 Ma 的钾长花岗岩 (熊子良等, 2012), 这些花岗岩体被认

为是北祁连洋壳俯冲、碰撞及造山后垮塌作用的产物。本文研究的窑洞沟侵入体位于甘肃省肃南县西水乡窑洞沟到黄草沟一带, 主要由 4 个小岩体组成, 呈不规则状、条带状展布, 与奥陶系阴沟群火山岩呈侵入接触关系, 出露面积约 9 km²。岩性上以中粗粒的花岗岩为主, 局部地段可见二长花岗岩和花岗闪长岩出露。大地构造位置上位于前人划分的走廊弧后盆地上 (图 1)。

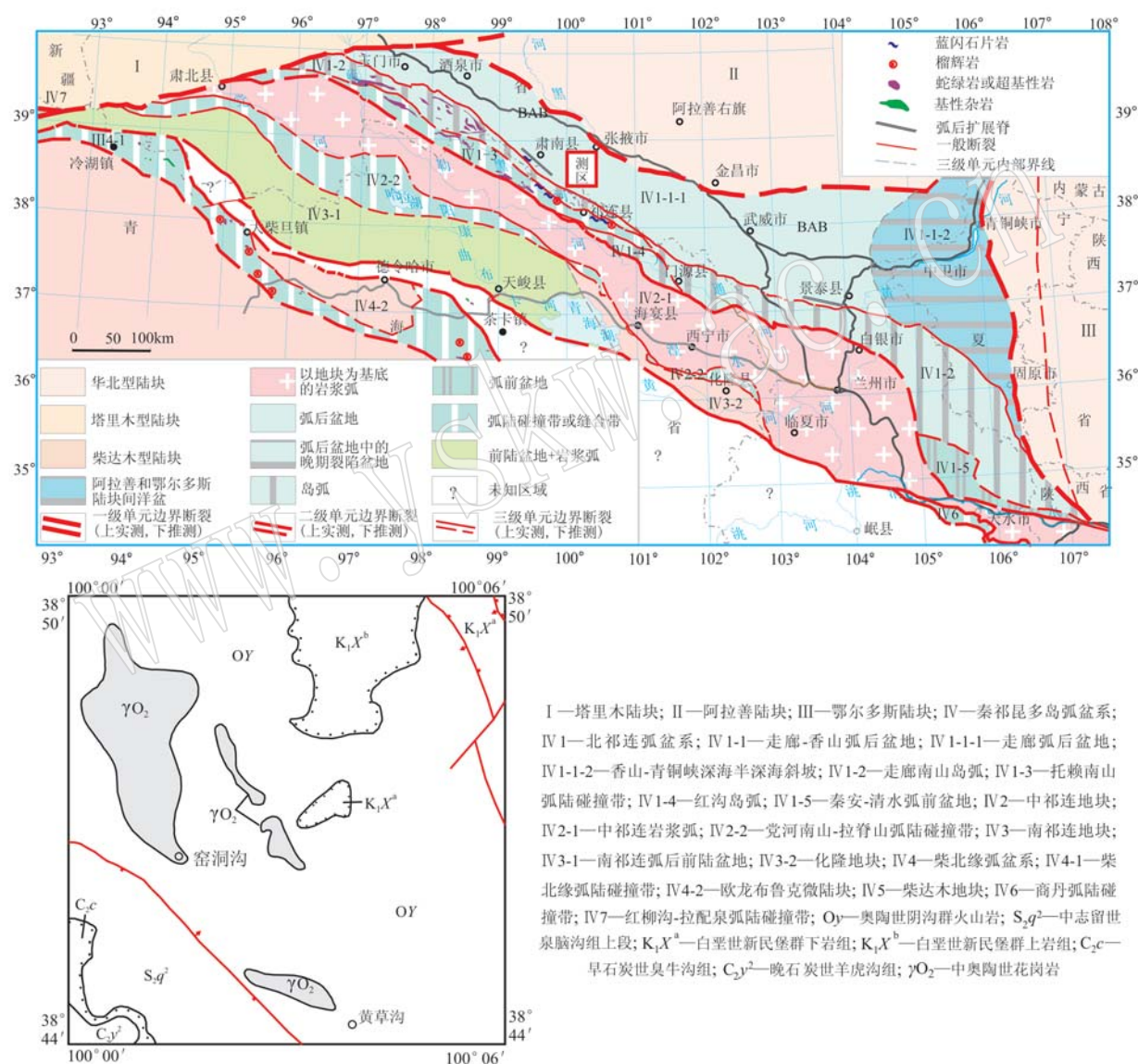


图 1 北祁连造山带构造单元划分图 (据冯益民等, 1996 修改) 及窑洞沟岩体分布图

Fig. 1 Tectonic subdivisions (after Feng Yimin *et al.*, 1996) and distribution of granite in the Yaodonggou area, Gansu

2 岩相学特征

窑洞沟岩体主体上为灰白色中粗粒花岗岩, 岩石新鲜面呈灰色-灰白色, 中-粗粒花岗结构, 块状构

造, 主要组成物质有斜长石、碱性长石、石英及极少量的暗色矿物, 其中, 斜长石白色到乳白色, 呈半自形板柱状结构, 含量 10% ~ 12%, 板径 2 ~ 6 mm, 双晶发育, 边缘熔蚀明显, 蚀变较弱, 为微弱的钠黝帘石化, 测得 An = 18 ~ 20, 为酸性的更长石, 碱性长石

为半自形-它形板状结构,含量 50% ~ 52%,内部由于应力作用多数呈现多晶化或角砾化,条纹发育,无双晶,蚀变弱,负低突起, I 级灰白干涉色;石英为它形晶,无蚀变,无解理,含量 28% ~ 30%,波状消光和

变形纹比较发育,粒径稍微小于长石,正低突起, I 级灰白干涉色,岩石中还可可见少量黑云母及后期蚀变矿物,含量 6% 左右(图 2)。岩体整体蚀变较强,多见硅化,绿泥石化、褐铁矿化蚀变。

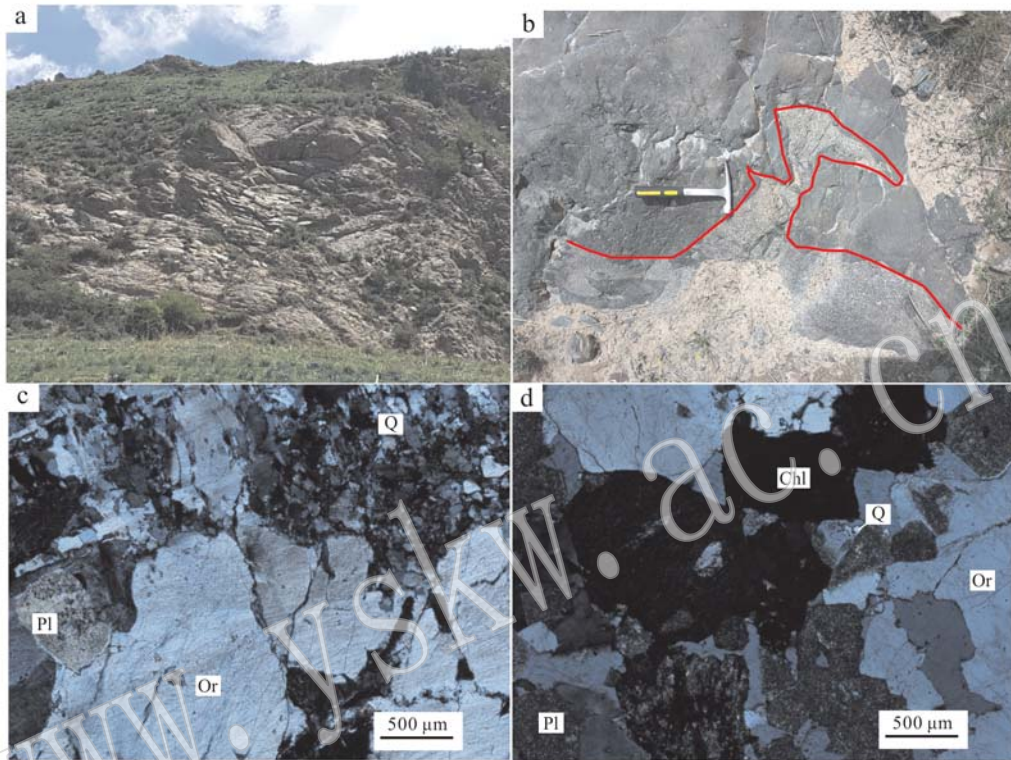


图 2 窑洞沟岩体野外及镜下照片(+)

Fig. 2 Field photographs and microphotographs of granite in the Yaodonggou area, Gansu(+)

Q—石英; Pl—斜长石; Or—碱性长石; Chl—绿泥石

Q—quartz; Pl—plagioclase; Or—alkaline feldspar; Chl—chlorite

3 LA-ICP-MS 锆石 U-Pb 测年

3.1 分析方法

对样品(2017YT3)中挑选出的锆石进行了 LA-ICP-MS 锆石 U-Pb 同位素检测工作。锆石样品靶的制作、阴极发光(CL)检测照片的拍摄及 U-Pb 微区同位素测定均在西北大学大陆动力学国家重点实验室完成。首先,在显微镜下挑选出透光性好、晶型完整无裂纹、粒径较大的锆石放置在双面胶纸上,灌注环氧树脂,待其固化后打磨并抛光,使锆石中心暴露,然后拍摄阴极发光图像,用于测定时选取锆石颗粒和测试部位。锆石阴极发光(CL)检测在电子探针实验室 MonoCL3 系统上完成,检测时其电子束加速电压为 10 kV。锆石 U-Pb 微区同位素测定在电感

耦合等离子体质谱仪(Agilent 7500a)与准分子激光剥蚀系统(GeoLas 2005)联机完成,测试时激光束斑直径为 32 μm ,剥蚀深度为 20 ~ 40 μm ,实验中采用 He 作为剥蚀物质的载气,用 NISTSRM 610 对仪器进行优化,锆石年龄采用标准锆石 91500 作为外标,用 ^{29}Si 作内标,NISTSRM610 作外标来校正微量元素含量,数据处理采用 Glitter 4.0 软件完成。按照 Common Pb Correction(ver3.15)方法对普通 Pb 进行校正(Andersen, 2002),应用 ISOPLOT 3.0 程序(Ludwig, 2003)进行锆石年龄加权平均值计算及 U-Pb 谐和图的绘制。

3.2 分析结果

显微镜下观测结果显示,样品中的锆石自形程度较高,多呈长柱状-柱状,无色透明,长度大多在 100 ~ 200 μm 之间,长宽比为 2:1,阴极发光(CL)图

像(图3)显示清晰的岩浆型锆石的振荡环带。样品(2017YT3)共进行了30颗锆石颗粒的测定,分析结果见表1。测点中Th的含量变化范围为 $116.55 \times 10^{-6} \sim 500.68 \times 10^{-6}$,U的含量变化范围为 $294.88 \times 10^{-6} \sim 657.04 \times 10^{-6}$,Th/U值较大,变化于0.39~0.88,平均值0.60,且Th、U之间正相关性较好,说明锆石为岩浆成因(Claesson *et al.*, 2000; Belousova

et al., 2002; 吴元保等, 2004)。剔除1个不谐和点(2017YT3-23),有29个测点在谐和线或其附近, $^{206}\text{Pb}/^{238}\text{U}$ 的表面年龄范围为 $470 \pm 5 \text{ Ma} \sim 462 \pm 5 \text{ Ma}$, $^{206}\text{Pb}/^{238}\text{U}$ 年龄加权平均值为 $466 \pm 2 \text{ Ma}$,置信度95%,MSWD=0.13(图4),这一年龄解释为花岗岩的结晶年龄,对应地质历史时期的中奥陶世。

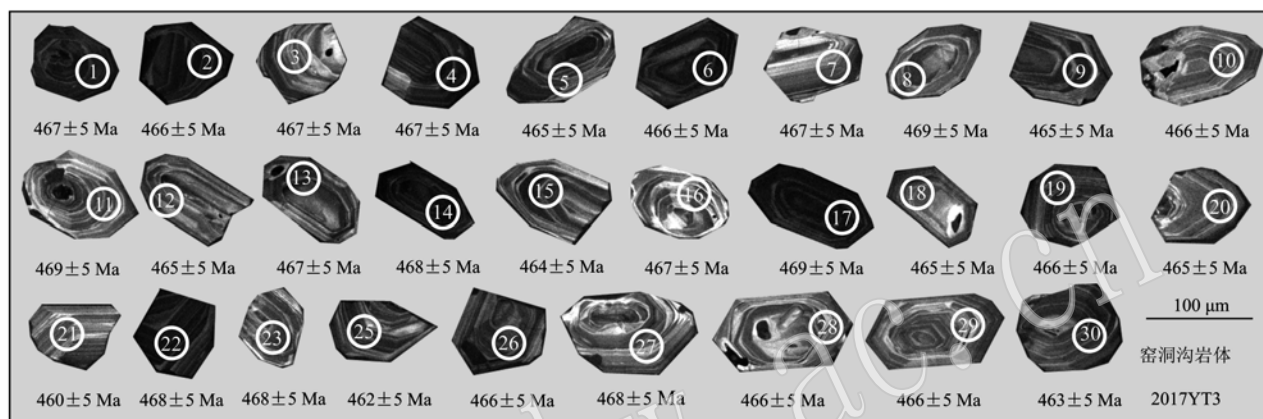


图3 窑洞沟花岗岩锆石阴极发光检测图像(CL)及测试位置

Fig. 3 Measuring position and CL images of zircon grains from granite in the Yaodonggou area, Gansu

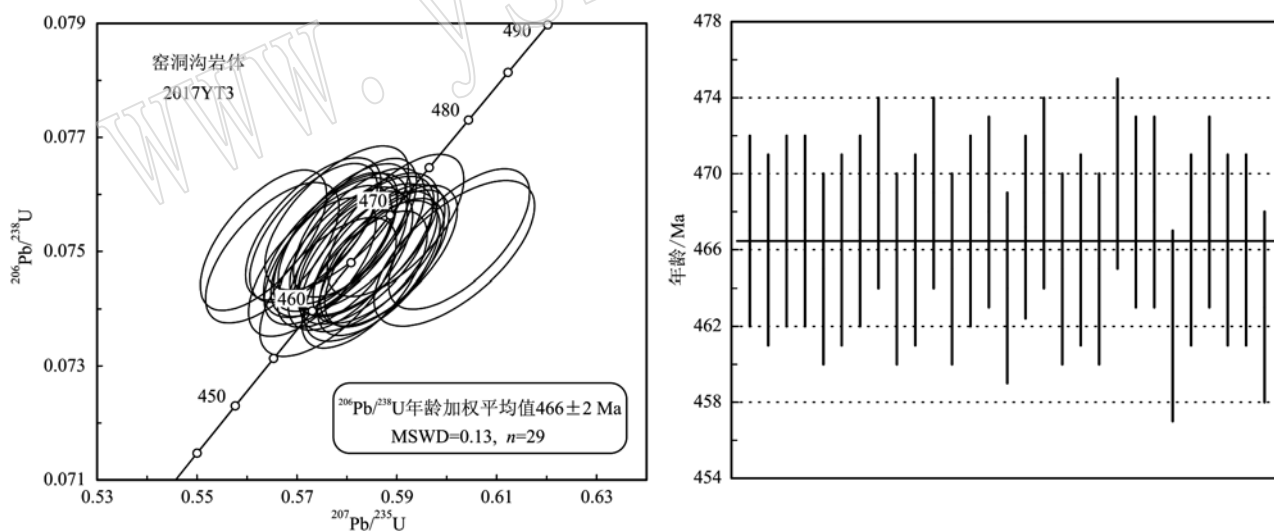


图4 窑洞沟花岗岩 LA-ICP-MS 锆石 U-Pb 年龄谐和图

Fig. 4 LA-ICP-MS U-Pb concordia diagram of zircon from granite in the Yaodonggou area, Gansu

4 地球化学特征

花岗岩的主量、微量元素分析测试均在中国地质调查局西安地质调查中心实验测试中心完成,其中,主量元素中的 Fe_2O_3 和FeO采用湿化学分析法

测试完成,其余主量元素采用荧光光谱法完成,分析仪器为Panalytical公司生产的PW440型X荧光光谱仪(XRF),分析误差低于5%,微量元素和稀土元素测试仪器为ThermoFisher公司生产的X-Series II型电感耦合等离子质谱仪(ICP-MS),检测限优于 5×10^{-9} ,相对标准偏差低于5%。

表 1 窑洞沟花岗岩 LA-ICP-MS U-Pb isotopic composition of the zircon of the granite in the Yaodonggou area, Gansu

分析点号	$w_B/10^{-6}$		Th/U	同位素比值				年龄/Ma			
	Th	U		$^{207}\text{Pb}/^{206}\text{Pb}$	$\pm 1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$\pm 1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$\pm 1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$\pm 1\sigma$
2017YT3-01	316.24	487.76	0.648 352	0.054 33	0.000 80	0.563 31	0.008 35	0.075 20	0.000 81	385	16
2017YT3-02	191.18	389.83	0.490 419	0.056 08	0.000 83	0.579 10	0.008 58	0.074 90	0.000 81	456	16
2017YT3-03	123.59	294.88	0.419 120	0.055 79	0.000 92	0.577 72	0.009 51	0.075 11	0.000 83	444	18
2017YT3-04	218.49	457.99	0.477 063	0.055 77	0.000 80	0.577 40	0.008 33	0.075 08	0.000 80	443	15
2017YT3-05	299.96	509.61	0.588 607	0.055 62	0.000 82	0.573 08	0.008 43	0.074 73	0.000 80	437	16
2017YT3-06	201.11	325.66	0.617 546	0.054 66	0.000 85	0.565 07	0.008 82	0.074 98	0.000 81	398	17
2017YT3-07	125.22	304.22	0.411 610	0.056 37	0.000 92	0.583 90	0.009 46	0.075 13	0.000 82	467	18
2017YT3-08	288.21	388.85	0.741 160	0.056 68	0.000 87	0.589 73	0.009 03	0.075 46	0.000 82	479	16
2017YT3-09	332.03	657.04	0.505 342	0.056 28	0.000 79	0.580 72	0.008 23	0.074 84	0.000 80	463	15
2017YT3-10	315.57	524.33	0.601 854	0.055 77	0.000 82	0.576 24	0.008 44	0.074 94	0.000 80	443	16
2017YT3-11	350.48	534.58	0.655 617	0.055 14	0.000 86	0.573 28	0.008 95	0.075 40	0.000 82	418	17
2017YT3-12	235.32	353.39	0.665 893	0.056 64	0.000 88	0.584 39	0.009 06	0.074 83	0.000 81	478	17
2017YT3-13	465.51	527.31	0.882 818	0.056 38	0.000 87	0.583 79	0.008 99	0.075 09	0.000 81	467	17
2017YT3-14	224.48	415.42	0.540 395	0.055 80	0.000 84	0.579 50	0.008 75	0.075 51	0.000 81	444	16
2017YT3-15	266.91	440.29	0.606 191	0.056 83	0.000 87	0.585 08	0.008 95	0.074 66	0.000 81	485	16
2017YT3-16	500.68	567.98	0.881 510	0.055 74	0.000 79	0.577 73	0.008 20	0.075 17	0.000 80	442	31
2017YT3-17	418.38	649.02	0.644 633	0.055 93	0.000 78	0.581 62	0.008 15	0.075 42	0.000 80	450	15
2017YT3-18	444.71	604.14	0.736 088	0.056 94	0.000 81	0.586 77	0.008 35	0.074 73	0.000 79	489	15
2017YT3-19	178.05	379.52	0.469 145	0.058 19	0.001 04	0.601 51	0.010 61	0.074 97	0.000 84	537	20
2017YT3-20	441.21	583.66	0.755 937	0.056 95	0.000 79	0.588 02	0.008 22	0.074 88	0.000 79	490	15
2017YT3-21	179.36	301.36	0.595 169	0.055 45	0.000 91	0.577 96	0.009 48	0.075 59	0.000 83	430	18
2017YT3-22	167.83	387.01	0.433 658	0.056 11	0.000 86	0.582 06	0.008 91	0.075 23	0.000 81	457	17
2017YT3-23	214.31	409.08	0.523 883	0.055 20	0.000 81	0.573 33	0.008 43	0.075 33	0.000 80	420	16
2017YT3-25	116.55	295.73	0.394 109	0.056 18	0.000 88	0.576 22	0.009 02	0.074 38	0.000 80	459	17
2017YT3-26	226.58	401.63	0.564 151	0.056 47	0.000 86	0.583 68	0.008 89	0.074 96	0.000 81	471	16
2017YT3-27	223.36	412.16	0.541 925	0.058 09	0.000 90	0.602 47	0.009 32	0.075 22	0.000 81	533	17
2017YT3-28	219.59	394.29	0.556 925	0.056 74	0.000 84	0.586 42	0.008 73	0.074 96	0.000 80	481	16
2017YT3-29	275.04	458.21	0.600 249	0.056 95	0.000 88	0.588 37	0.009 07	0.074 93	0.000 81	490	17
2017YT3-30	363.10	571.42	0.635 435	0.056 80	0.000 80	0.583 53	0.008 28	0.074 51	0.000 79	484	15

4.1 主量元素

窑洞沟岩体的主量元素分析结果及特征参数见

表2。据表可知,灰白色中粗粒花岗岩具有极高的SiO₂含量,介于72.45%~80.18%之间,全碱含量为

表2 窑洞沟地区花岗岩主量元素($w_B/\%$)和微量元素($w_B/10^{-6}$)分析数据

Table 2 Chemical compositions of major elements ($w_B/\%$) and trace elements ($w_B/10^{-6}$) for granite in the Yaodonggou area, Gansu

样号	2017YT-3-1H	2017YT-3-2H	2017YT-3-3H	2017YT-3-4H	2017YT-3-5H	2017YT-3-6H
SiO ₂	74.70	73.70	79.36	77.88	80.18	72.45
Al ₂ O ₃	12.46	12.75	11.23	11.76	10.16	13.57
Fe ₂ O ₃	0.34	0.30	0.11	0.13	0.17	0.49
FeO	1.06	1.23	0.66	0.82	0.55	1.24
CaO	1.34	1.45	0.74	0.80	1.08	1.21
MgO	0.71	0.81	0.46	0.53	0.35	0.91
K ₂ O	2.70	2.89	0.43	0.96	1.65	3.30
Na ₂ O	4.64	4.65	5.75	5.78	4.31	4.67
TiO ₂	0.20	0.22	0.16	0.16	0.13	0.24
P ₂ O ₅	0.05	0.05	0.03	0.03	0.03	0.06
MnO	0.04	0.04	0.03	0.02	0.02	0.05
灼失	1.64	1.83	0.98	1.07	1.29	1.72
H ₂ O ⁺	0.74	0.84	0.48	0.50	0.56	0.98
Alk	7.47	7.69	6.25	6.82	6.04	8.12
K ₂ O/Na ₂ O	0.58	0.62	0.07	0.17	0.38	0.71
A/CNK	0.96	0.95	1.00	0.98	0.94	1.01
La	20.60	16.60	10.80	14.00	16.20	21.40
Ce	34.00	25.70	21.20	27.50	31.60	40.70
Pr	5.30	4.63	2.30	3.14	3.53	5.62
Nd	19.00	16.80	7.88	11.00	11.50	20.00
Sm	3.86	3.21	1.37	1.91	2.08	3.75
Eu	0.85	0.78	0.29	0.43	0.53	0.87
Gd	3.55	2.97	1.21	1.80	1.83	2.92
Tb	0.58	0.47	0.19	0.27	0.28	0.45
Dy	3.42	2.66	1.02	1.54	1.62	2.56
Ho	0.68	0.56	0.21	0.32	0.34	0.56
Er	1.96	1.59	0.59	0.96	0.95	1.65
Tm	0.30	0.26	0.096	0.15	0.16	0.26
Yb	1.98	1.72	0.66	1.03	1.02	1.79
Lu	0.31	0.28	0.10	0.17	0.17	0.28
Y	24.80	16.30	4.86	8.65	9.26	14.30
ΣREE	121.19	94.53	52.78	72.87	81.07	117.11
δEu	0.69	0.76	0.67	0.70	0.81	0.77
(La/Sm) _N	3.45	3.34	5.09	4.73	5.03	3.68
(La/Yb) _N	7.46	6.92	11.74	9.75	11.39	8.58
(Gd/Yb) _N	1.48	1.43	1.52	1.45	1.48	1.35
Cu	4.79	4.85	12.10	7.79	40.10	6.34
Pb	8.38	7.79	9.05	7.38	8.80	10.20
Zn	15.40	16.00	16.20	15.70	10.80	19.00
Cr	3.92	2.70	2.50	3.38	2.19	3.55
Ni	0.72	3.55	1.78	0.12	0.53	1.04
Co	1.98	2.19	1.36	1.75	1.96	2.51
Li	6.18	7.12	2.52	2.75	3.42	7.26
Rb	58.90	49.30	14.00	24.50	37.90	74.20
Cs	0.92	0.89	0.65	0.49	0.85	1.04
Sr	79.40	66.80	81.10	75.30	91.50	103.00
Ba	494	518	87.9	224	360	612
V	14.80	16.70	5.36	5.24	4.64	20.20
Sc	7.63	11.30	9.40	11.40	5.81	7.68
Nb	5.88	7.25	4.56	5.67	4.78	7.30
Ta	0.48	0.56	0.38	0.45	0.36	0.56
Zr	127	133	112	122	97	147
Hf	3.50	3.88	2.92	3.33	2.67	4.17
Be	1.10	1.08	0.86	0.93	1.04	1.28
Ga	10.8	11.00	8.38	9.86	8.55	12.40
U	2.16	2.46	1.72	1.99	1.98	3.04
Th	11.30	11.20	5.51	7.10	7.37	12.00

6.04% ~ 8.12% 且 K_2O 小于 Na_2O , K_2O/Na_2O 值为 0.07 ~ 0.71, Al_2O_3 含量较低, Al_2O_3 含量介于 10.16% ~ 13.57%, A/CNK 介于 0.94 ~ 1.01, 平均为 0.97; 岩石里特曼指数在 0.96 ~ 2.16 之间, 均小于 3.3; 在 $(Na_2O + K_2O) - SiO_2$ 图上(图5), 均落入

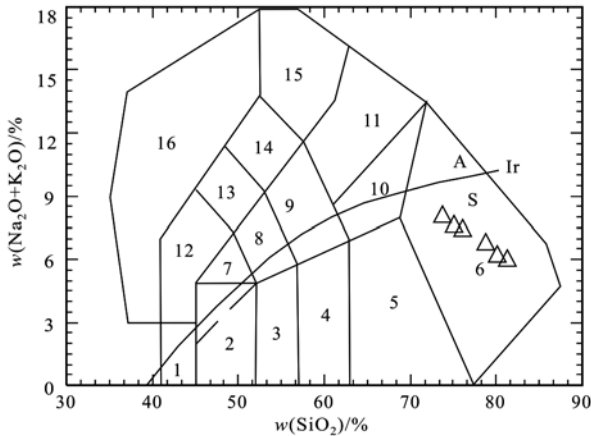


图5 窑洞沟花岗岩体 $(Na_2O + K_2O) - SiO_2$ 图

(底图据 Cox *et al.*, 1979)

Fig. 5 $(Na_2O + K_2O) - SiO_2$ diagram of granite in the

Yaodonggou area, Gansu (after Cox *et al.*, 1979)

1—橄榄辉长岩; 2—辉长岩; 3—辉长闪长岩; 4—闪长岩; 5—花岗闪长岩; 6—花岗岩; 7—二长辉长岩; 8—二长闪长岩; 9—二长岩; 10—石英二长岩; 11—正长岩; 12—似长辉长岩; 13—似长二长闪长岩; 14—似长正长闪长岩; 15—似长正长岩; 16—似长岩; A—碱性系列; S—亚碱性系列

1—olivine gabbro; 2—gabbro; 3—gabbro diorite; 4—diorite; 5—granodiorite; 6—granite; 7—monzogabbro; 8—monzodiorite; 9—monzonite; 10—quartz-monzonite; 11—syenite; 12—fooidites gabbro; 13—fooidites-monzodiorite; 14—fooidites-syenodiorite; 15—fooiditesy-enite; 16—fooidites; A—alkaline series; S—subalkalic series

亚碱性花岗岩区域内, 在 $K_2O - SiO_2$ 图解(图 6a)上, 窑洞沟岩体在低钾拉斑玄武系列到中钾钙碱性系列区域内均有出现, 在 $A/NK - A/CNK$ 图解上(图 6b), 均落入准铝质花岗岩区域内。

4.2 稀土元素

据表 2 可知, 窑洞沟花岗岩的稀土元素总量 (ΣREE) 介于 $52.78 \times 10^{-6} \sim 121.19 \times 10^{-6}$, 其中, 轻稀土元素 (LREE) 总量为 $43.84 \times 10^{-6} \sim 92.34 \times 10^{-6}$, 重稀土元素 (HREE) 总量为 $8.94 \times 10^{-6} \sim 37.58 \times 10^{-6}$, 轻重稀土元素间分馏明显, $LREE/HREE = 2.22 \sim 4.91$, $(La/Yb)_N$ 值为 6.92 ~ 11.74。轻稀土元素内部分馏较明显, 其 $(La/Sm)_N$ 值为 3.34 ~ 5.09, 重稀土元素内部分馏不明显, $(Gd/Yb)_N$ 值为 1.35 ~ 1.52; 在球粒陨石标准化的稀土元素配分图上表现出轻稀土元素富集的右倾型, 具有轻微-中等的负 Eu 异常, δEu 介于 0.67 ~ 0.81 之间(图 7a)。

4.3 微量元素

据表 2 可知, 窑洞沟岩体的大离子亲石元素 (LILE) Rb、Sr、Ba 含量分别为 $14.00 \times 10^{-6} \sim 74.20 \times 10^{-6}$, $66.80 \times 10^{-6} \sim 103 \times 10^{-6}$, $87.90 \times 10^{-6} \sim 612 \times 10^{-6}$, 放射性生热元素 (RPH) U、Th 含量分别为 $1.72 \times 10^{-6} \sim 3.04 \times 10^{-6}$, $5.51 \times 10^{-6} \sim 11.30 \times 10^{-6}$, 高场强元素 (HFSE) Nb、Ta、Zr、Hf 含量分别为 $4.56 \times 10^{-6} \sim 7.30 \times 10^{-6}$, $0.38 \times 10^{-6} \sim 0.56 \times 10^{-6}$, $97 \times 10^{-6} \sim 147 \times 10^{-6}$, $2.67 \times 10^{-6} \sim 4.17 \times 10^{-6}$ 。在原始地幔标准化微量元素蛛网图(图 7b)上, 普遍

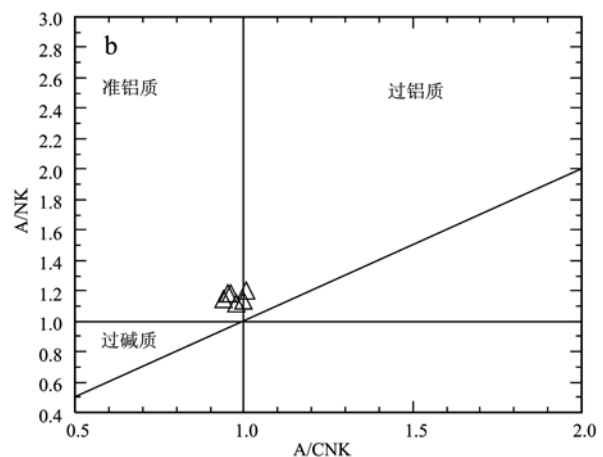
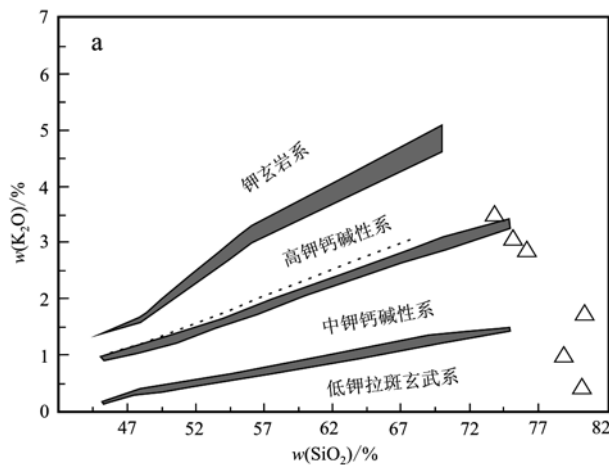


图6 窑洞沟岩体 $K_2O - SiO_2$ 图(a, 据 Rickwood, 1989)和 $A/NK - A/CNK$ 图(b, 据 Maniar and Piccolli, 1989)

Fig. 6 K_2O versus SiO_2 diagram (a, after Rickwood, 1989) and A/NK versus A/CNK diagram (b, after Maniar and

Piccolli, 1989) of granite in the Yaodonggou area, Gansu

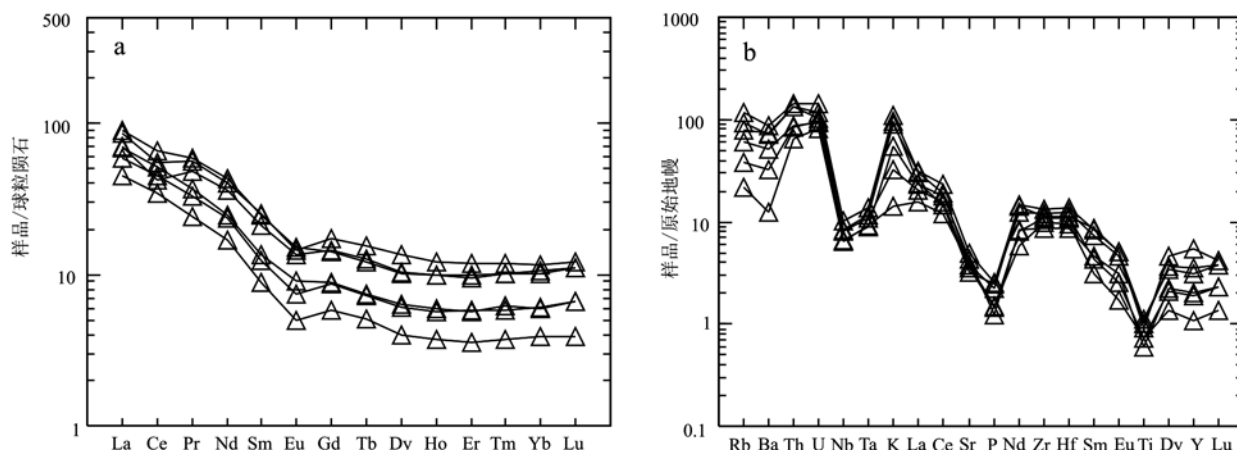


图7 窑洞沟岩体稀土元素配分图(a)和微量元素蛛网图(b)(标准化数据据 Sun and McDonough, 1989)

Fig. 7 REE patterns (a) and trace elements spidergrams (b) of granite in the Yaodonggou area, Gansu (normalized values after Sun and McDonough, 1989)

表现出 Nb、Ta、Ti、P 等高场强元素的亏损和 Rb、K、U 等元素的富集, 显示出弧岩浆岩的特点。

5 岩石成因

花岗岩源区性质的判定及成因类型的划分一直是花岗岩研究工作中最重要的问题, 目前被大家广泛接受的划分方案主要有 I、S、M、A 型 4 种 (Pitcher, 1993; 吴福元等, 2007), 其中, 由于 M 型花岗岩多与蛇绿岩套相关, 比较少见, 因此, 常见的花岗岩主要为 S 型、I 型和 A 型。窑洞沟花岗岩高硅、富碱、贫铝, SiO_2 含量介于 72.45% ~ 80.18%, ALK 介于 6.04 ~ 8.12, Al_2O_3 介于 10.3 ~ 13.82, 类似于 A 型花岗岩, 然而其 $10^4 \times \text{Ga}/\text{Al} = 1.41 \sim 1.73 < 2.6$, $\text{Zr} + \text{Nb} + \text{Ce} + \text{Y} = 142.64 \times 10^{-6} \sim 209.3 \times 10^{-6} < 350 \times 10^{-6}$, 均低于 A 型花岗岩边界值 (Whalen *et al.*, 1987; 吴福元等, 2007), 从而排除为 A 型花岗岩的可能性。窑洞沟花岗岩体富铝矿物少见, A/CNK 介于 0.94 ~ 1.01, 平均为 0.97, 小于 1.0, 为准铝质岩石, CIPW 标准矿物计算有 4 个样品不出现刚玉分子, 两个有刚玉分子的样品 (2017YT3-3H、2017YT3-6H) C 分子介于 0.031 ~ 0.264 之间, 均小于 1, 不具有 S 型花岗岩的特点 (Chappell *et al.*, 1974, 1999)。实验岩石学研究表明, 磷灰石在弱过铝质岩浆和准铝质岩浆中的溶解度很低, 且与 SiO_2 的含量呈负相关关系; 而在强过铝质岩浆中, 磷灰石溶解度明显增高且与 SiO_2 的含量呈正相关关系或基本保持不变 (Wolf and London, 1994), 磷灰石在不同铝质岩浆中

的这种独特的地球化学行为被成功地用于区分 S 型和 I 型花岗岩类 (Chappell, 1999; Wu *et al.*, 2003; Li *et al.*, 2007)。在 $\text{P}_2\text{O}_5 - \text{SiO}_2$ 图解上 (图 8a), 二者具有较好的负相关性, 显示出与 I 型花岗岩一致的变化趋势, 这一趋势还能得到 Y-Rb 图解的支持, 因为富 Y 矿物不会在准铝质 I 型岩浆演化的早期阶段结晶出来, 从而引起分异的 I 型花岗岩的 Y 含量升高且与 Rb 含量呈正相关关系 (Chappell, 1999; Li *et al.*, 2006; 李献华等, 2007), 在 Rb-Y 图解上 (图 8b), 也表现出与分异 I 型花岗岩一致的趋势。窑洞沟花岗岩矿物组成上主要为石英和长石, 暗色矿物极少, SiO_2 含量极高, P_2O_5 、 MgO 、 TiO_2 、 Fe_2O_3 、 FeO 含量极低, 也有类似于高分异花岗岩的特征, 然而, 其 Nb/Ta 值为 12.00 ~ 13.28, 平均为 12.68; Zr/Hf 比值为 34.28 ~ 38.36, 平均为 36.19, 又与高分异花岗岩具有较低的 Nb/Ta 及 Zr/Hf 比值不同 (Bau, 1996; Dostal and Chatterjee, 2000; Linnen and Kerppler, 2002; Claiborne *et al.*, 2006; Deering and Bachmann, 2010; Ballouard *et al.*, 2016; 吴福元等, 2017)。研究表明, 高分异花岗岩具有极低的锆石 Zr/Hf 值, 窑洞沟岩体锆石 Zr/Hf 值介于 29.79 ~ 37.01 之间, 平均 31.87, 根据 Breiter 等 (2014) 划分方案 (普通花岗岩的 Zr/Hf 值大于 55; 中等分异花岗岩 Zr/Hf 值大于 25, 小于 55; 高分异花岗岩 Zr/Hf 值小于 25), 窑洞沟岩体属于中等分异花岗岩。综上, 认为窑洞沟花岗岩体属于中等分异的 I 型花岗岩, 其初始岩浆在就位过程中可能经历了中等程度的分

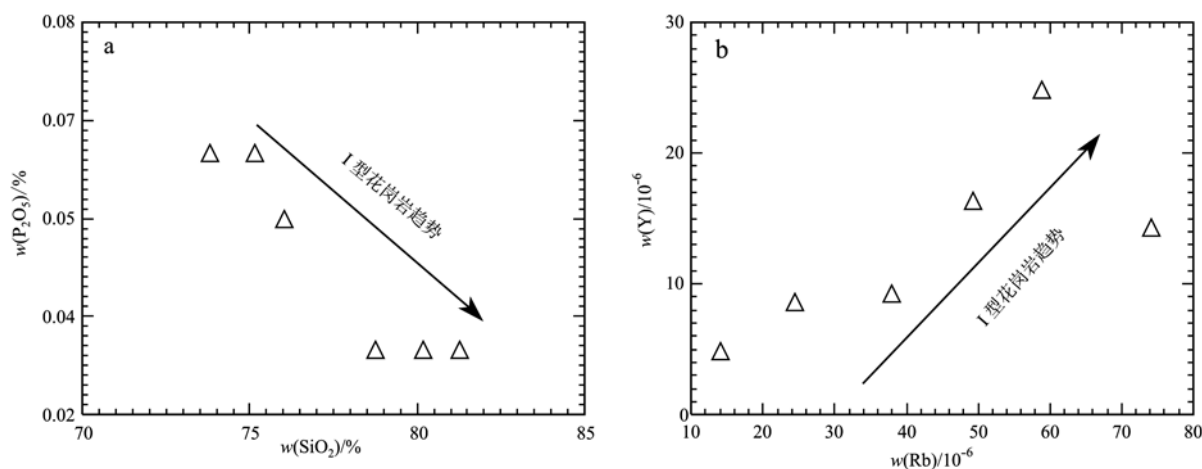


图8 窑洞沟岩体 P_2O_5 - SiO_2 图(a, 据 Wolf and London, 1994) 和 Y - Rb 图(b, 据 Li *et al.*, 2006)

Fig. 8 P_2O_5 versus SiO_2 diagram (a, after Wolf and London, 1994) and Y versus Rb diagram (b, after Li *et al.*, 2006) of granite in the Yaodonggou area, Gansu

离结晶作用。

6 动力学背景分析

区域构造研究表明,北祁连从早元古代中期开始,大陆岩石圈拉伸、减薄,并发生裂谷化(左国朝等, 1987; 夏林圻等, 1995, 2000; 葛肖虹等, 1999); 至新元古代,裂谷作用进一步加强,发育以双峰式火山岩为特征的大陆裂谷火山作用; 到晚寒武世,最终发生大陆裂解和分离,形成北祁连早古生代洋盆,于奥陶纪北祁连洋盆进入俯冲消减和弧后盆地协同演化阶段,发育大量岛弧和弧后盆地火山

岩(左国朝等, 1987; 夏林圻等, 1991, 1992, 1995; 葛肖虹等, 1999; 张旗等, 2000; Xia *et al.*, 2003; 曾建元等, 2007; 夏小洪等, 2010; Song *et al.*, 2013), 至 445 ~ 424 Ma 之间,洋盆闭合进入陆内碰撞和深俯冲过程(Xia *et al.*, 2003; 吴才来等, 2006; Zhang *et al.*, 2007; Song *et al.*, 2007a, 2007b, 2009a, 2009b)。窑洞沟岩体普遍表现出 Nb、Ta、Ti、P 等高场强元素的亏损和 Rb、K、U 等元素的富集,显示出弧岩浆岩的特点,利用 Pearce 等(1984)微量元素构造判别图解,在 Nb - Y 构造环境判别图上(图 9a),投影点主要位于火山弧到同碰撞花岗岩区域内,在 Rb - (Y + Nb) 构造环境判别图上(图 9b)

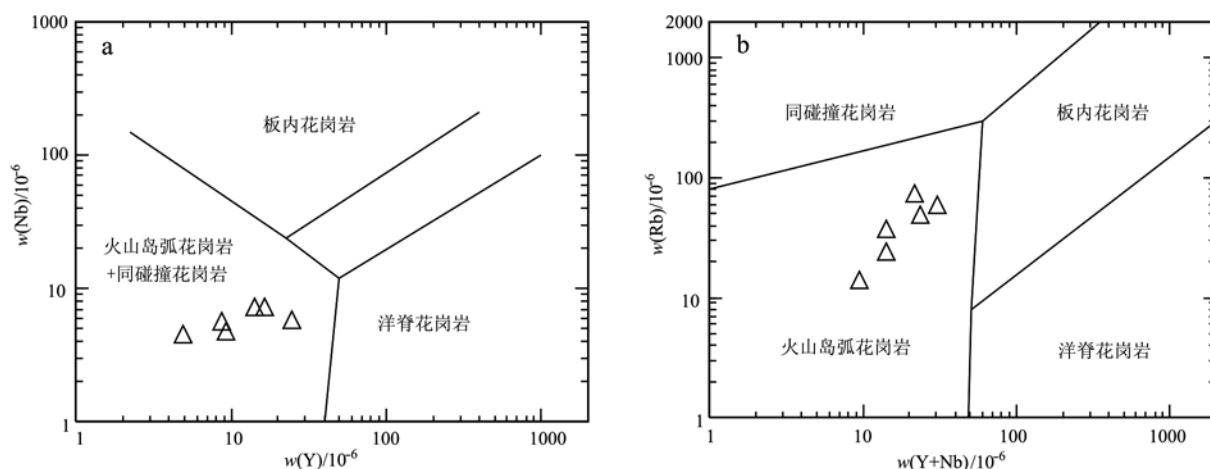


图9 窑洞沟岩体 Nb - Y 图(a)和 Rb - (Y + Nb) 构造环境判别图(b)(底图据 Pearce *et al.*, 1984)

Fig. 9 Nb versus Y diagram (a) and Rb versus Y + Nb diagram (b) of granite in the Yaodonggou area, Gansu (after Pearce *et al.*, 1984)

主要位于火山弧区域内。结合本区最新获得的窑洞沟岩体的高精度年代学数据(466 Ma)及其所处的大地构造位置, 本文认为窑洞沟岩体的形成可能与北祁连洋盆持续的向北俯冲, 导致弧后拉张伸展, 诱发幔源岩浆岩底侵上升, 底侵的幔源岩浆提供热量, 导致古老地壳物质熔融形成初始岩浆, 而后经历了中等程度分离结晶作用有关。其岩浆具有岛弧岩浆属性, 可能是由于弧后盆地是在岛弧的基础上裂开的或者其岩浆源区受到了俯冲板片流体/熔体的改造作用。

7 结论

(1) 花岗岩的 $^{206}\text{Pb}/^{238}\text{U}$ 年龄加权平均值为 466 ± 2 Ma, MSWD = 0.13, 代表其结晶年龄, 这为北祁连造山带早古生代岩浆活动事件提供了一个可靠年代学约束。

(2) 窑洞沟岩体具有高硅、低钾、低磷含量特征, 铝饱和指数(A/CNK)为 0.94 ~ 1.01, 富集 Rb、Th、U, 明显亏损 Nb、Ta、Sr、Ti、P 和 Eu 等元素, 属准铝质中等分异 I 型花岗岩。

(3) 窑洞沟岩体的形成可能与北祁连洋盆持续向北俯冲导致弧后拉张伸展, 诱发幔源岩浆岩底侵上升, 底侵的幔源岩浆提供热量, 导致古老地壳物质熔融形成初始岩浆, 而后经历中等程度分离结晶作用有关。其岩浆具有岛弧岩浆属性, 可能是由于弧后盆地在岛弧的基础上裂开或者其岩浆源区受到了俯冲板片流体/熔体的改造作用。

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