

吉林南部新太古代末地壳深熔作用

——来自变质石英闪长岩及淡色花岗岩的证据

刘建辉

(中国地质科学院 地质研究所, 北京 100037; 自然资源部深地动力学重点实验室, 北京 100037)

摘要: 深熔作用是大陆地壳分异、元素迁移富集和混合岩化作用的主要机制和关键地质过程。吉南地区出露的太古宙基底普遍经历了角闪岩相-麻粒岩相变质及深熔作用, 长英质淡色体及淡色花岗岩广泛分布。吉南和龙花岗-绿岩地体出露的太古宙变质石英闪长岩及相关的长英质浅色体和含斜方辉石(角闪石)淡色伟晶花岗岩的野外地质特征、相互关系及岩相学特征指示变质石英闪长岩经历了部分熔融, 熔体发生了原地分凝聚集。锆石 CL 图像、U-Pb 年代学、微量元素及 Hf 同位素分析显示, 变质石英闪长岩与淡色伟晶花岗岩具有相同的岩浆锆石和变质重结晶锆石为主的锆石类型组成, 其中后者以变质重结晶锆石为主, 变质石英闪长岩的岩浆结晶年龄为 2535 ± 14 Ma, 二者的岩浆锆石具有统一的 Hf 同位素组成, 变质重结晶锆石具有统一的 U-Pb 年龄, 分别为 2495 ± 22 Ma 和 2491 ± 10 Ma, 统一的锆石微区 Th、U 含量及 Th/U 值, 并且具有与岩浆锆石统一的 Hf 同位素成分, 指示两者变质重结晶锆石与岩浆锆石具有形成于同一岩浆源区的初始结晶锆石, 进一步指示变质石英闪长岩在约 2.5 Ga 经历了变质深熔作用, 含斜方辉石(角闪石)淡色伟晶花岗岩源自变质石英闪长岩的原地部分熔融。结合华北克拉通及吉南地区新太古代末准同时的强烈岩浆作用和具有逆时针 $p-T$ 轨迹的角闪岩相-麻粒岩相变质作用, 其可能反映了华北克拉通东部陆块新太古代末在形成稳定克拉通过程中地幔物质上涌的热效应。

关键词: 深熔作用; 长英质浅色体; 吉南地区; 混合岩化; 华北克拉通

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Late Neoarchean crustal anatexis in the southern Jilin Province: Evidence from metamorphic quartz diorite and associated leucogranite

LIU Jian-hui

(Institute of Geology, Chinese Academy of Geological Sciences, Beijing 100037, China; Key Laboratory of Deep-Earth Dynamics of Ministry of Natural Resources, Beijing 100037, China)

Abstract: Anatexis is the main mechanism and key geological process of continental crustal differentiation, element migration and enrichment and migmatization. The Archean basement exposed in the southern Jilin Province, in which felsic leucosome and leucogranites are extensively distributed, widely experienced amphibolite-granulite facies metamorphism and anatexis. The field geological characteristics, correlation and petrographic characteristics of the Archean metamorphic quartz diorite and associated felsic leucosome and orthopyroxene (Opx)- and amphibole (Amp)-bearing pegmatite granite occurred in Helong granite-greenstone terrane in the southern Jilin Province indicate that the metamorphic quartz diorite underwent partial melting and in-situ segregation and accumulation of melt. The integrated analyses of zircon cathodoluminescence (CL) images, in-situ U-Pb geochronology, trace elements

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作者简介: 刘建辉(1978-), 男, 研究员, 主要从事构造地质学和前寒武纪地质研究, E-mail: liujianhui1999@163.com。

and Hf isotope on the metamorphic quartz diorite and pegmatite granite show that they have the same types of magmatic and metamorphic recrystallized zircons, the latter is dominated by metamorphic recrystallized zircons, that the metamorphic quartz diorite has magmatic crystallization age of $2\ 535 \pm 14$ Ma, and that their magmatic zircons have consistent Hf isotope composition. Their metamorphic recrystallized zircons yield consistent metamorphic ages of $2\ 495 \pm 22$ Ma and $2\ 491 \pm 10$ Ma, respectively, and have identical Hf isotope compositions, which are consistent with that of magmatic zircons, Th and U contents and Th/U ratios, indicating that both metamorphic recrystallized and magmatic zircons have initial crystalline zircons formed in the same magmatic source, and further demonstrating that the metamorphic quartz diorite experienced metamorphic anatexis at ca. 2.5 Ga and that the Opx- and Amp-bearing pegmatite granite derived from in-situ partial melting of the metamorphic quartz diorite, in combination of the coeval strong and extensive magmatism and amphibolite-granulite facies metamorphism with counterclockwise *p-T* path at the end of Neoproterozoic in the North China Craton as well as the southern Jilin Province, which may reflect the thermal effect of mantle material upwelling on the continental crust during formation and stabilization of the craton lithosphere.

Key words: anatexis; felsic leucosome; the southern Jilin Province; migmatization; North China Craton

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深熔作用(anatexis)是早前寒武纪高级变质地体和造山带高级变质岩中非常普遍的地质作用(贺同兴, 1987; Sawyer, 1987, 2008; Ashworth and Brown, 1990; Patino Douce *et al.*, 1990; Brown, 1994a, 1994b; 程裕淇等, 2000, 2001, 2004; Liu *et al.*, 2012, 2019; 曾令森等, 2017; Zhou *et al.*, 2019),包括含水矿物的脱水部分熔融和水致部分熔融(Beard and Lofgren, 1991; Patino Douce and Beard, 1995; Clemens *et al.*, 1996, 2016; Clemens, 2006; Aranovich *et al.*, 2014; Weinberg and Hasalová, 2015; Schwindinger *et al.*, 2019),是形成深熔混合岩和深熔片麻岩的主要机制,也是地壳中花岗质岩浆形成、大陆地壳分异演化的主要机制和关键过程,具有深远的构造物理和地球化学效应(贺同兴, 1987; Sawyer, 1987, 2008; Ashworth and Brown, 1990; Brown, 2007, 2013; Zheng *et al.*, 2011; Zheng and Gao, 2021)。因此,过去几十年,针对不同岩石类型(泥质或基性岩)深熔作用的熔融反应、温压条件及热源(Bohlen, 1987; Harley, 1989; Bea, 2012; 魏春景, 2016; 魏春景等, 2016, 2017, 2021; Huang *et al.*, 2021)、流体(水)或挥发分(CO_2 、 Cl)在地壳熔融过程的潜在作用(Holtz *et al.*, 2001; Bachmann and Bergantz, 2006; Newton and Manning, 2008, 2010; Harlov, 2012; Aranovich *et al.*, 2014; Newton *et al.*, 2014; Collins *et al.*, 2021),熔体的

形成、分凝聚集、分离及运移(Brown, 2007, 2013; Clemens and Stevens, 2016),深熔条件下的转熔反应及转熔矿物的形成(夏琼霞, 2019 及其参考文献),深熔作用过程中副矿物(如锆石)的行为(Harley and Nandakumar, 2014; 王伟等, 2014; Zeng *et al.*, 2021 及其参考文献),深熔作用过程中元素的地球化学行为(Zheng *et al.*, 2011)及地壳的构造物理效应(如含熔体地壳的流变)(Brown, 2007, 2013)等方面,从对天然深熔混合岩和深熔片麻岩样品(露头)的研究到深熔实验及相平衡模拟开展了广泛研究。这些研究促进并丰富了对地壳深熔作用的理解,并为理解和研究大陆地壳的分异机制、元素的富集作用、深部地壳的流变学行为及物理性质等提供了重要依据。

华北克拉通早前寒武纪基底除西部孔兹岩带、中部造山带及胶-辽-吉带3条构造/活动带在古元古代经历了与造山作用有关的变质深熔作用外(Yin *et al.*, 2014; Liu *et al.*, 2019),其太古宙基底普遍经历了与新太古代末花岗质(TTG)岩浆事件准同期的深熔作用(程裕淇等, 2000, 2001, 2004; 任留东等, 2011; Dong *et al.*, 2017; Wang *et al.*, 2017; Jiang *et al.*, 2021; 图1a)。位于华北克拉通东部的吉林南部地区出露的太古宙基底深熔作用普遍而强烈(孙德有等, 1998; Guo *et al.*, 2018; Yu *et al.*, 2021),出露大量与深熔作用有关的淡色花岗质岩

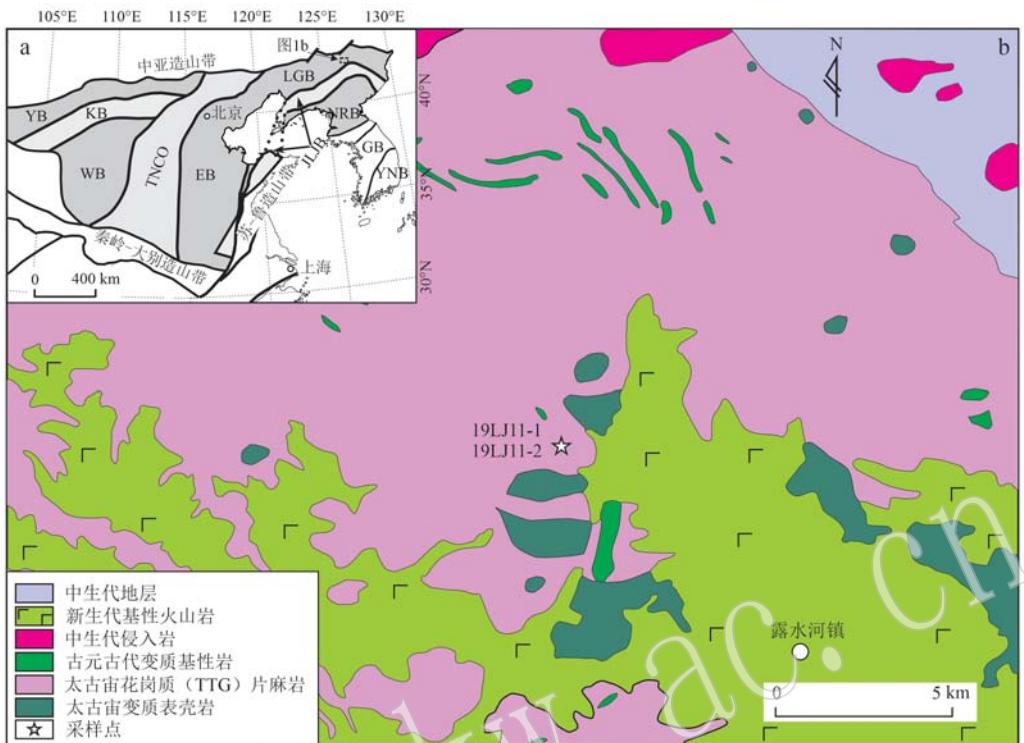


图 1 华北克拉通构造单元划分(a, 据 Zhao et al., 2005 修改)和研究区地质简图及采样位置(b)

Fig. 1 Tectonic subdivision of the North China Craton (a, modified after Zhao et al., 2005) and simplified

geological map of the research region (b)

YB—阴山陆块; KB—孔兹岩带; WB—西部陆块; TNCO—华北中部造山带; EB—东部陆块; LGB—龙岗地块; NRB—狼林地块; JLJB—胶-辽-吉构造带; GB—京畿地块; YNB—永南地块
YB—Yinshan Block; KB—Khondalite Belt; WB—Western Block; TNCO—Trans-North China Orogen; EB—Eastern Block; LGB—Longgang Block; NRB—Nangrim Block; JLJB—Jiao-Liao-Ji Belt; GB—Gyeonggi Block; YNB—Yeongnam Block

脉或岩体，对其形成时代、源区及成因机制的深入研究能为理解研究区太古宙末地壳演化及构造体制提供重要的证据和约束。因此，本文选取吉林南部和龙太古宙地体内出露的经历了变质深熔作用的变质石英闪长岩及与之相伴的淡色花岗岩开展了详细的野外地质特征及相互关系的观察，并结合室内岩相学、锆石 U-Pb 年代学、微量元素特征及 Lu-Hf 同位素特征的综合研究，确定了深熔作用的时间及成因机制。

1 地质背景

华北克拉通早前寒武纪变质基底由西部陆块、东部陆块及联接它们的胶-辽-吉带、中部造山带及西部孔兹岩带(Zhao et al., 2005)或称胶辽造山/活动带、晋豫造山/活动带和丰镇造山/活动带等(翟明国等, 2007; Zhai and Santosh, 2011)3条古元古代造山/活动带所组成(图 1a)。东部陆块被胶-辽-吉古元古构造带进一步划分成北部的龙岗地块及南部

的狼林地块(Zhao et al., 2005; 图 1a)，北部的龙岗地块主要由太古宙花岗质片麻岩、变质基性-超基性岩及变质表壳岩组成，其中英云闪长质-奥长花岗质-花岗闪长质(TTG)片麻岩是最主要的岩石类型，尤其以新太古代晚期 TTG 片麻岩分布最广(Geng et al., 2012; Guo et al., 2015, 2016, 2017, 2018; Wang et al., 2016a, 2016b; Wan et al., 2018; Yu et al., 2021)。近年来，随着研究的深入，中太古代晚期-新太古代早期(2.9~2.7 Ga)的 TTG 片麻岩也在越来越多的地区被发现(Liu et al., 2013; Yang et al., 2013; Wan et al., 2014, 2018; Liou and Guo, 2019)，而中太古代早期-始太古代 TTG 片麻岩的报道相对较少，仅在鞍本及冀东地区小范围内被发现(Song et al., 1996; Wan et al., 2012)；已发现的最古老岩石为鞍本地区约 3.80 Ga 的奥长花岗质岩石(Liu et al., 1992, 2008; Song et al., 1996; Wu et al., 2008)，它们普遍遭受强烈的变质变形及深熔作用。南部的狼林地块主要由新太古代 TTG 片麻岩及古元古代花岗质岩石及变质表壳岩组成；传统

上, 龙岗地块与狼林地块被认为具有统一的太古宙变质结晶基底, 称为中朝克拉通(张秋生等, 1988; 伍家善等, 1998), 最新的研究进一步证实龙岗地块与狼林地块具有一致的早前寒武纪基底和演化历史(Zhai et al., 2019; Zhao et al., 2020)。胶-辽-吉古元古代造山/活动带主要由古元古代花岗质岩石、变质基性-超基性岩及变质火山-沉积岩系组成(张秋生等, 1988; 卢良兆等, 1996; Lu et al., 2006; Li and Zhao, 2007; Liu et al., 2014, 2021; 刘福来等, 2015), 前人围绕其构造属性及构造演化开展了广泛的研究, 目前, 对于其经历了起始的大陆地壳的伸展及其后的地壳增厚的造山作用已达成共识, 但对驱动大陆地壳的起始伸展及构造演化的动力学机制仍有不同的理解(Xu and Liu, 2019; Liu et al., 2020, 2021)。

位于龙岗地块东部的吉林南部地区是华北克拉通太古宙结晶基底出露的主要地区之一, 包括夹皮沟、白山、板石沟及和龙等太古宙花岗绿岩地体。和龙太古宙花岗绿岩地体是吉林南部太古宙花岗绿岩带重要的组成部分之一, 太古宙变质基底主要由大面积出露的钠质TTG、富钾质花岗质片麻岩及少量变质闪长岩、变质基性岩及变质表壳岩等组成, 变质表壳岩包括黑云(角闪)变粒岩、斜长角闪岩及浅粒岩夹条带状磁铁石英岩(葛文春等, 1993, 1996a, 1996b; 孙德有等, 1995, 1998; Ge et al., 2003; Guo et al., 2015, 2016, 2017, 2018; Yu et al., 2021; 图1b); 太古宙变质闪长岩及变质基性岩呈透镜体或脉体分布在太古宙花岗质片麻岩内(图1b); 出露的太古宙变质基底普遍遭受了新太古代末具有逆时针 $p-T$ 轨迹的角闪岩相-麻粒岩相变质作用(Ge et al., 2003)、韧性剪切变形及深熔作用(孙德有等, 1998; Yu et al., 2021)。此外, 大量的古元古代变质基性岩呈岩墙或岩脉产出与太古宙花岗质片麻岩中(图1b), 其成因仍有待深入研究。

2 地质特征及岩相学特征

本文研究的露头位于和龙太古宙花岗绿岩地体内露水河镇西北(图1b), 出露有暗色的变质石英闪长岩及含斜方辉石(角闪石)巨晶的淡色伟晶花岗岩, 变质石英闪长岩与淡色伟晶花岗岩的接触带呈渐变关系(图2a), 变质石英闪长岩显示了明显的剪切变形及变质深熔作用, 深熔作用形成的长英质浅

色体要么呈薄的浅色团块或迷雾状分布在暗色的变质石英闪长岩内, 要么呈浅色脉体沿剪切面理分布(图2b), 或沿张性裂隙、破裂面呈浅色块体或脉体分布, 内部含有少量暗色矿物, 矿物粒度明显变粗(图2c)。变质石英闪长岩的矿物组合包括角闪石、斜长石、黑云母、钾长石(主要为微斜长石)及石英, 呈它形不等粒结构, 含水矿物角闪石及黑云母呈港湾状及熔蚀残余结构, 并且斜长石具有明显的“糟化”和“暗化”现象(图2e, 2f), 指示它们经历了深熔作用的改造; 在这些早期残余矿物之间分布有以钾长石(微斜长石)、斜长石及石英为主的浅色矿物, 它们自形程度相对较好, 矿物表面干净, 呈花岗岩浆结构(图2g), 代表变质石英闪长岩深熔作用形成的矿物。含斜方辉石(角闪石)淡色伟晶花岗岩粒径达2 cm(图2d), 主要矿物组合包括钾长石、石英、紫苏辉石(斜方辉石)及角闪石(图2h, 2i), 其中紫苏辉石呈它形不等粒结构, 以包体形式分布于角闪石内(图2h), 巨晶钾长石及角闪石粒间分布有细粒的钾长石及石英(图2h, 2i), 含斜方辉石(角闪石)淡色伟晶花岗岩可能代表变质石英闪长岩变质深熔熔体分凝聚集结晶的产物。

3 测试方法

对采自吉林南部和龙太古宙花岗绿岩地体内露水河镇西北的1个变质石英闪长岩样品(19LJ11-1)和1个含斜方辉石(角闪石)淡色伟晶花岗岩样品(19LJ11-2)开展了锆石原位U-Pb年代学、微量、稀土元素及Lu-Hf同位素测试分析。

LA-ICP-MS锆石原位U-Pb测年及锆石微区微量元素、稀土元素分析在北京快科赛默科技测年实验室进行。分析仪器采用ESI公司生产的准分子激光剥蚀系统(New Wave NWR 193^{UC})和美国AGILENT科技有限公司生产的Agilent ICP-MS/MS 8900型串联四级杆电感耦合等离子体质谱仪联合构成的激光等离子质谱仪。分析时采用激光单点剥蚀采样, 激光束斑和频率分别为20 μm和5 Hz, 能量密度为3.5 J/cm², 载气为高纯氦气(He)。测试样品时使用标准锆石91500为外标进行同位素比值校正, 以Plešovice锆石为监控标样。测试所获数据采用Iolite(Paton et al., 2010)进行处理。锆石U-Pb年龄谐和图及加权平均年龄计算采用IsoPlot 4.15完成, 测试结果见表1。

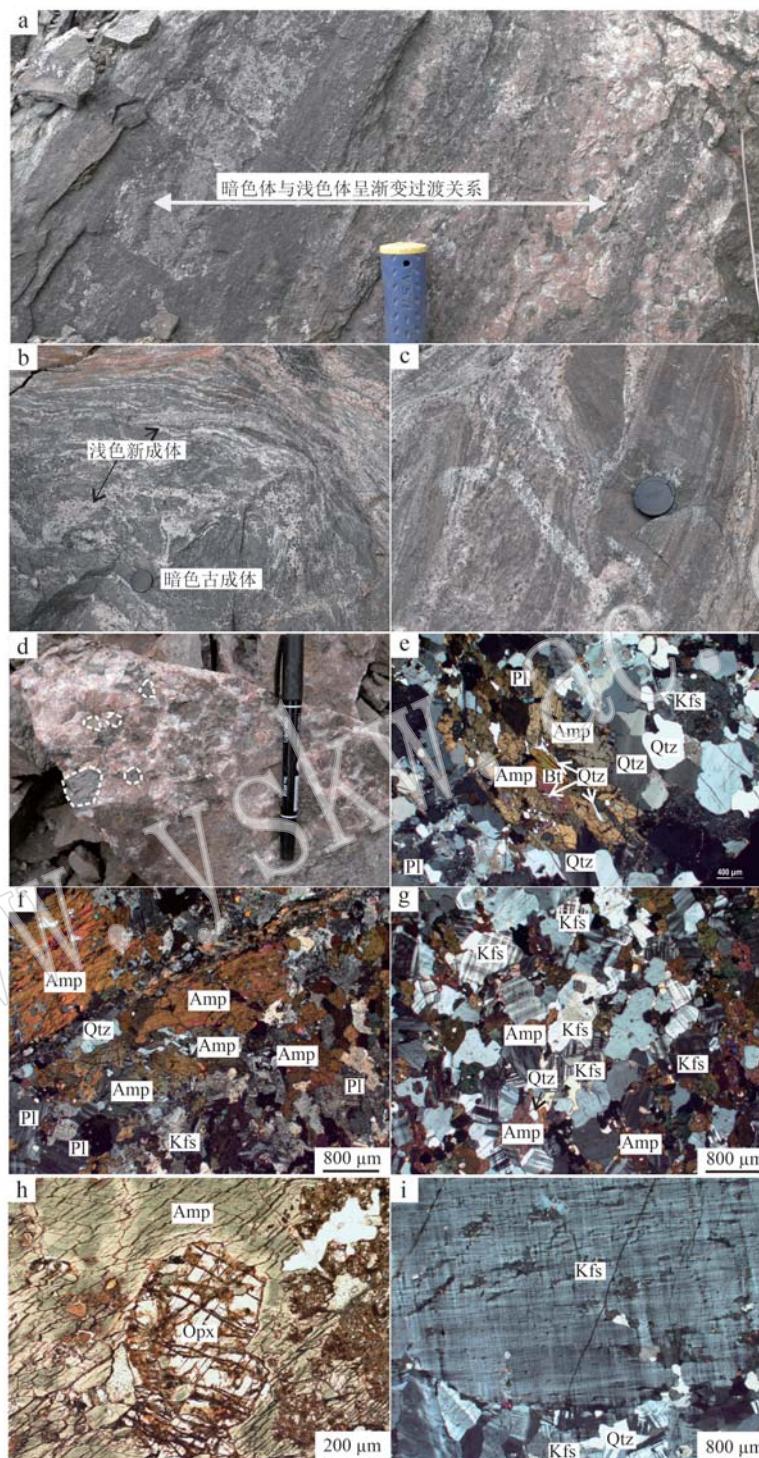


图2 吉南和龙太古宙基底变质石英闪长岩、长英质淡色体及淡色花岗岩野外照片及显微照片
Fig. 2 Photographs of outcrops, micrographs and their correlation of the metamorphic quartz diorite, leucosome and leucogranite in the Archean basement of Helong region, southern Jilin Province

a—变质石英闪长岩及含斜方辉石(角闪石)淡色伟晶花岗岩地质特征及相互关系; b, c—变质石英闪长岩及长英质浅色体地质特征及相互关系; d—含斜方辉石(角闪石)淡色伟晶花岗岩; e, f, g—变质石英闪长岩正交偏光照片; h, i—含斜方辉石(角闪石)淡色伟晶花岗岩单偏光及正交偏光照片; Qtz—石英; Pl—斜长石; Kfs—钾长石; Amp—角闪石; Opx—斜方辉石; Bt—黑云母
a—the geological features and relationship of the metamorphic quartz diorite and Opx- and Amp-bearing leucocite pegmatite granite; b, c—the geological features and relationship of the metamorphic quartz diorite and felsic leucosome; d—Opx- and Amp-bearing leucocite pegmatite granite; e, f, g—micrographs of the metamorphic quartz diorite; h, i—micrographs of the Opx- and Amp-bearing leucocite pegmatite granite; Qtz—quartz; Pl—plagioclase; Kfs—K-feldspar; Amp—amphibole; Opx—orthopyroxene; Bt—biotite

锆石 Hf 同位素测试在中国地质科学院矿产资源研究所自然资源部同位素地质重点实验室进行。分析仪器采用 Finnigan Neptune 多接收电感耦合等离子体质谱加载 New Wave UP213 紫外激光剥蚀系统(LA-MC-ICP-MS)，锆石 Hf 同位素分析在 U-Pb 测年点原位进行，实验过程中采用氦气作为剥蚀物质载气，激光剥蚀束斑直径采用 $55 \mu\text{m}$ ，剥蚀时间为 27 s。测试时选用锆石国际标样 GJ-1 作为参考物质。相关仪器运行条件及详细分析流程与侯可军等(2007)相同。分析过程中锆石标样 GJ-1 的 $^{176}\text{Hf}/^{177}\text{Hf}$ 测试加权平均值为 $0.282\ 024 \pm 0.000\ 004$ ($2\text{SD}, n=38$)，与文献报道值(Elhlou *et al.*, 2006; 侯可军等, 2007)在误差范围内完全一致。在 $\varepsilon\text{Hf}(t)$ 计算时，球粒陨石的 $^{176}\text{Hf}/^{177}\text{Hf}$ 值为 0.282 772, $^{176}\text{Lu}/^{177}\text{Hf}$ 值为 0.033 2(Blichert-Toft and Albarede, 1997)。在单阶段 Hf 模式年龄 t_{DM} 计算时，亏损地幔的 $^{176}\text{Hf}/^{177}\text{Hf}$ 值和 $^{176}\text{Lu}/^{177}\text{Hf}$ 值分别为 0.283 25 和 0.038 4(Griffin *et al.*, 2000)；在两阶段 Hf 模式年龄 $t_{\text{DM}}(\text{Hf}_2)$ 计算时，下地壳、平均地壳与亏损地幔的 $f_{\text{Lu/Hf}}$ 值分别为 -0.32、-0.548 2 及 0.156 6(Amelin *et al.*, 1999; Griffin *et al.*, 2000, 2002)。 ^{176}Lu 的衰变常量选用 $1.867 \times 10^{-11} \text{ a}^{-1}$ (Soderlunde *et al.*, 2004; Amelin, 2005)；相关计算中锆石的 U-Pb 年龄选择初始岩浆锆石结晶 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄，相关计算公式参考吴福元等(2007)，测试结果见表 2。

4 测试结果

4.1 变质石英闪长岩(19LJ11-1)

变质石英闪长岩样品(19LJ11-1)中的锆石主要

以自形柱状为主，少量为椭球状或不规则粒状，锆石长度主要分布在 $50\sim150 \mu\text{m}$ 之间，长短轴之比在 3~1 之间。CL 图像显示大部分锆石具有典型岩浆结晶环带的核、暗色的变质重结晶幔及极薄的亮色变质增生边组成的核幔边结构，部分锆石无锆石核，内部结构及色调均匀，与具有核幔边结构锆石幔部相同的内部结构及色调显示了变质重结晶锆石的特征(图 3a)。24 个具有典型振荡环带岩浆锆石域分析给出的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄范围在 $2\ 541\sim2\ 489$ Ma 之间(表 1)，大部分分析点显示了轻微的铅丢失，谐和性较好(图 4a；表 1)，其加权平均年龄为 $2\ 519 \pm 6$ Ma($\text{MSWD}=3.4, n=24$)，上交点年龄为 $2\ 535 \pm 14$ Ma($\text{MSWD}=1.6$ ，图 4a)。20 个具有均匀内部结构的暗色变质重结晶锆石域分析给出的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄分布在 $2\ 506\sim2\ 436$ Ma 之间，大部分分析点显示了不同程度的铅丢失(图 4a；表 1)，其加权平均年龄为 $2\ 476 \pm 7$ Ma($\text{MSWD}=5.3, n=20$)，上交点年龄为 $2\ 495 \pm 22$ Ma($\text{MSWD}=2.4$ ，图 4a)。具有典型岩浆锆石振荡环带的锆石域和无振荡环带、内部结构均匀的暗色变质重结晶锆石域具有类似的球粒陨石标准化稀土元素配分曲线模式，均显示轻重稀土元素分馏明显的左倾型模式，富集重稀土元素而亏损轻稀土元素(图 4b, 4c)；岩浆锆石域具有相对较为统一的 U 及 Th 含量，而变质重结晶锆石域相对岩浆锆石域具有较高的、变化范围较大的 U 及 Th 含量(图 4d, 4e)；它们的 Th/U 值均大于 0.1，并且岩浆锆石域相对变质锆石域显示了更高的 Th/U 值(表 1；图 4f)。笔者将岩浆锆石域获得的上交年龄 $2\ 535 \pm 14$ Ma($\text{MSWD}=1.6$)和变质锆石域获得的上交点年龄 $2\ 495 \pm 22$ Ma($\text{MSWD}=2.4$ ，图 4a)分别解释为岩浆侵位(结晶)年龄和变质年龄。

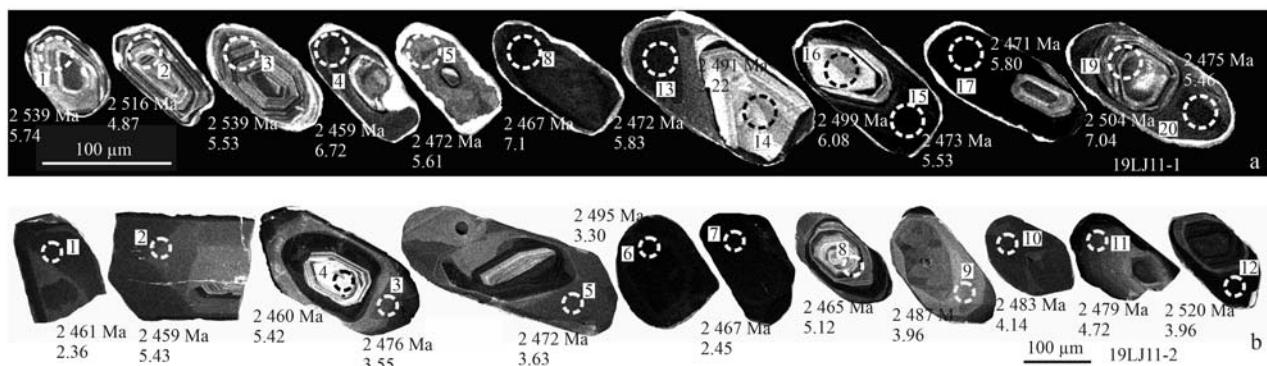


图 3 变质石英闪长岩(a)及淡色伟晶花岗岩(b)锆石 CL 图像、U-Pb 年龄及 $\varepsilon\text{Hf}(t)$ 值

Fig. 3 The cathodoluminescence (CL) images with U-Pb ages and $\varepsilon\text{Hf}(t)$ values for different zircon domains from the metamorphic quartz diorite sample (a) and leucogranite sample (b)

表1 吉南和龙地区变质石英闪长岩及淡色伟晶花岗岩 LA-ICP-MS 锆石 U-Pb 年龄数据表
Table 1 LA-ICP-MS U-Pb dating data for zircons from the metamorphic quartz diorite and leucogranite samples collected from the Helong region in the southern Jilin Province

分析点	锆石类型	$w_B/10^{-6}$	同位素比值						年龄/Ma	误差/%								
			Pb	Th	U	Th/U	$^{207}\text{Pb}/^{235}\text{U}$	$\pm 1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$\pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$\pm 1\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$\pm 1\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$\pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb}$	$\pm 1\sigma$
变质石英闪长岩(19LJ11-1)																		
19LJ11-1#1	岩浆锆石	161.00	138.50	148.10	0.94	10.780 0	0.150 0	0.462 6	0.006 8	0.168 4	0.001 7	2.502	13	2.449	30	2.539	17	103.67
19LJ11-1#2	岩浆锆石	174.00	140.00	174.40	0.80	10.710 0	0.140 0	0.463 9	0.006 5	0.166 2	0.001 6	2.496	12	2.455	29	2.516	16	102.48
19LJ11-1#3	岩浆锆石	188.10	162.20	223.20	0.73	10.590 0	0.130 0	0.454 4	0.005 8	0.168 3	0.001 4	2.486	11	2.413	26	2.539	14	105.22
19LJ11-1#6	岩浆锆石	83.70	67.00	114.00	0.59	10.960 0	0.180 0	0.476 6	0.007 5	0.166 2	0.001 9	2.517	15	2.511	33	2.516	19	100.20
19LJ11-1#7	岩浆锆石	117.60	99.10	160.00	0.62	11.130 0	0.190 0	0.477 2	0.008 6	0.168 7	0.001 6	2.531	16	2.512	38	2.541	16	101.15
19LJ11-1#11	岩浆锆石	118.00	102.90	125.70	0.82	10.870 0	0.150 0	0.469 9	0.006 4	0.167 1	0.001 6	2.511	13	2.481	28	2.526	16	101.81
19LJ11-1#12	岩浆锆石	172.00	150.00	162.00	0.93	10.660 0	0.160 0	0.461 7	0.006 7	0.166 9	0.001 7	2.492	14	2.445	29	2.526	16	103.31
19LJ11-1#14	岩浆锆石	112.10	106.50	111.10	0.96	10.230 0	0.150 0	0.450 3	0.006 8	0.163 8	0.001 8	2.453	13	2.395	30	2.491	18	104.01
19LJ11-1#16	岩浆锆石	75.60	77.00	89.60	0.86	10.510 0	0.160 0	0.461 6	0.006 8	0.164 6	0.001 9	2.478	14	2.445	30	2.499	20	102.21
19LJ11-1#18	岩浆锆石	99.00	89.20	154.60	0.58	10.820 0	0.120 0	0.469 4	0.006 3	0.166 2	0.001 5	2.506	11	2.479	27	2.520	14	101.65
19LJ11-1#19	岩浆锆石	313.00	265.00	272.30	0.97	10.750 0	0.150 0	0.469 8	0.006 4	0.164 7	0.001 2	2.500	13	2.481	28	2.504	12	100.93
19LJ11-1#21	岩浆锆石	146.70	122.10	130.10	0.94	10.690 0	0.160 0	0.465 1	0.007 7	0.166 6	0.001 8	2.496	14	2.460	34	2.520	18	102.44
19LJ11-1#23	岩浆锆石	170.00	137.50	139.70	0.98	10.760 0	0.180 0	0.465 1	0.007 9	0.167 1	0.001 5	2.502	16	2.460	35	2.526	15	102.68
19LJ11-1#25	岩浆锆石	118.30	101.80	127.90	0.80	10.530 0	0.150 0	0.457 7	0.006 7	0.166 3	0.001 8	2.480	13	2.427	30	2.517	18	103.71
19LJ11-1#29	岩浆锆石	158.90	137.70	219.50	0.63	10.750 0	0.140 0	0.463 4	0.005 7	0.167 4	0.001 4	2.500	12	2.453	25	2.530	14	103.14
19LJ11-1#30	岩浆锆石	175.00	146.50	200.60	0.73	10.620 0	0.160 0	0.467 2	0.007 1	0.164 5	0.001 8	2.489	14	2.470	31	2.499	18	101.17
19LJ11-1#32	岩浆锆石	333.00	305.00	266.00	1.15	10.140 0	0.140 0	0.441 4	0.006 9	0.166 0	0.001 7	2.448	13	2.355	31	2.515	17	106.79
19LJ11-1#33	岩浆锆石	353.00	327.00	283.00	1.16	9.970 0	0.200 0	0.440 6	0.008 5	0.163 5	0.001 8	2.428	19	2.351	38	2.489	18	105.87
19LJ11-1#35	岩浆锆石	164.00	129.50	207.30	0.62	10.590 0	0.160 0	0.460 8	0.006 4	0.167 3	0.001 5	2.494	14	2.442	28	2.530	15	103.60
19LJ11-1#37	岩浆锆石	174.00	137.70	216.30	0.64	10.260 0	0.120 0	0.450 1	0.006 2	0.164 4	0.001 6	2.457	11	2.398	29	2.498	17	104.17
19LJ11-1#38	岩浆锆石	242.00	196.20	235.60	0.83	10.680 0	0.150 0	0.460 7	0.006 6	0.167 5	0.001 4	2.494	13	2.441	29	2.532	14	103.73
19LJ11-1#40	岩浆锆石	90.10	70.00	108.00	0.65	11.320 0	0.220 0	0.497 0	0.009 8	0.164 4	0.002 1	2.547	18	2.598	42	2.502	23	96.30
19LJ11-1#43	岩浆锆石	148.80	126.40	183.70	0.69	10.640 0	0.150 0	0.460 5	0.007 3	0.167 0	0.001 7	2.492	13	2.440	32	2.525	17	103.48
19LJ11-1#44	岩浆锆石	62.20	46.60	96.20	0.48	11.150 0	0.230 0	0.487 7	0.008 2	0.166 0	0.002 6	2.533	19	2.559	36	2.512	27	98.16
19LJ11-1#44	变质锆石	153.00	131.50	290.00	0.45	10.180 0	0.130 0	0.459 2	0.006 6	0.160 5	0.001 3	2.449	12	2.434	29	2.459	14	101.03
19LJ11-1#5	变质锆石	127.00	105.20	204.50	0.51	10.460 0	0.180 0	0.466 8	0.008 3	0.161 8	0.001 4	2.473	16	2.467	37	2.472	15	100.20
19LJ11-1#8	变质锆石	162.00	138.00	884.00	0.16	10.180 0	0.140 0	0.457 0	0.006 6	0.161 2	0.001 2	2.451	13	2.425	29	2.467	12	101.73
19LJ11-1#9	变质锆石	284.00	246.70	972.00	0.25	10.300 0	0.150 0	0.466 4	0.006 5	0.159 4	0.001 3	2.460	14	2.466	29	2.447	14	99.23
19LJ11-1#10	变质锆石	229.00	196.20	346.00	0.57	10.840 0	0.170 0	0.477 3	0.007 8	0.164 0	0.001 1	2.506	15	2.513	34	2.495	12	99.28
19LJ11-1#13	变质锆石	102.30	88.00	272.40	0.32	10.060 0	0.150 0	0.449 0	0.006 3	0.161 8	0.001 4	2.438	14	2.389	28	2.472	14	103.47
19LJ11-1#15	变质锆石	336.00	304.00	736.00	0.41	10.020 0	0.150 0	0.447 0	0.006 8	0.161 9	0.001 3	2.436	13	2.380	30	2.473	13	103.91
19LJ11-1#17	变质锆石	408.00	370.00	1430.00	0.26	10.040 0	0.140 0	0.448 9	0.006 7	0.161 5	0.001 0	2.437	12	2.388	30	2.471	11	103.48
19LJ11-1#20	变质锆石	203.00	171.70	302.40	0.57	10.070 0	0.140 0	0.449 7	0.006 2	0.162 1	0.001 4	2.441	13	2.393	28	2.475	15	103.43
19LJ11-1#22	变质锆石	140.00	132.70	292.50	0.45	10.300 0	0.150 0	0.456 5	0.006 6	0.162 9	0.001 4	2.459	13	2.422	29	2.483	14	102.52
19LJ11-1#24	变质锆石	504.00	434.00	931.00	0.47	10.300 0	0.130 0	0.457 0	0.006 1	0.162 7	0.001 1	2.460	12	2.429	26	2.482	11	102.18

续表 1
Continued Table 1

分析点	锆石类型	同位素比值						$w_{\text{B}}/10^{-6}$	Th/U	$^{207}\text{Pb}/^{235}\text{U} \pm 1\sigma$	$^{206}\text{Pb}/^{238}\text{U} \pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb} \pm 1\sigma$	$^{207}\text{Pb}/^{235}\text{U} \pm 1\sigma$	$^{206}\text{Pb}/^{238}\text{U} \pm 1\sigma$	$^{207}\text{Pb}/^{206}\text{Pb} \pm 1\sigma$	年龄/Ma	增和 度/%	
		Pb	Th	U	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$\pm 1\sigma$											
19LJ11-1#26	变质锆石	213.00	186.30	378.00	0.49	10.440 0	0.180 0	0.459 8	0.007 9	0.163 7	0.001 1	2.472	16	2.437	35	2.492	11	102.26
19LJ11-1#27	变质锆石	155.00	147.40	254.90	0.58	10.470 0	0.160 0	0.461 0	0.006 0	0.163 6	0.001 5	2.476	14	2.443	26	2.490	16	101.92
19LJ11-1#28	变质锆石	446.00	396.00	706.00	0.56	10.370 0	0.160 0	0.462 0	0.007 7	0.162 1	0.001 1	2.465	14	2.446	34	2.476	12	101.23
19LJ11-1#31	变质锆石	279.00	237.00	442.00	0.54	10.160 0	0.140 0	0.450 6	0.005 9	0.162 4	0.001 3	2.447	13	2.396	26	2.479	13	103.46
19LJ11-1#34	变质锆石	232.00	196.00	1201.00	0.16	9.800 0	0.120 0	0.446 7	0.007 1	0.158 1	0.001 3	2.415	12	2.379	32	2.436	15	102.40
19LJ11-1#36	变质锆石	94.50	76.30	332.00	0.23	10.370 0	0.150 0	0.462 2	0.007 2	0.161 6	0.001 3	2.466	13	2.453	33	2.471	14	100.73
19LJ11-1#39	变质锆石	23.80	23.30	204.40	0.11	10.200 0	0.120 0	0.446 2	0.005 8	0.164 8	0.001 4	2.453	11	2.377	26	2.506	14	105.43
19LJ11-1#41	变质锆石	310.00	271.00	425.00	0.64	9.830 0	0.140 0	0.436 1	0.006 1	0.162 8	0.001 3	2.417	13	2.332	27	2.485	13	106.56
19LJ11-1#42	变质锆石	46.80	38.10	167.70	0.23	10.320 0	0.180 0	0.462 1	0.007 6	0.161 0	0.001 7	2.464	16	2.447	34	2.466	18	100.78
浅色伟晶花岗岩 (19LJ11-2)																		
19LJ11-2#1	变质锆石	234.00	235.20	300.00	0.78	11.010 0	0.120 0	0.494 9	0.006 1	0.160 7	0.001 4	2.523	10	2.591	26	2.461	15	94.98
19LJ11-2#2	变质锆石	246.00	242.70	949.00	0.26	10.480 0	0.110 0	0.470 3	0.005 4	0.160 4	0.000 9	2.477	9.7	2.484	24	2.459	9.7	98.98
19LJ11-2#3	变质锆石	174.00	170.00	392.00	0.43	10.510 0	0.120 0	0.466 1	0.005 3	0.162 1	0.001 0	2.479	11	2.465	23	2.476	11	100.45
19LJ11-2#5	变质锆石	201.00	195.50	298.40	0.66	10.590 0	0.130 0	0.471 8	0.005 9	0.161 7	0.001 1	2.486	11	2.490	26	2.472	11	99.28
19LJ11-2#6	变质锆石	60.80	57.70	112.20	0.51	11.210 0	0.180 0	0.493 6	0.007 1	0.163 8	0.001 8	2.539	15	2.585	30	2.495	18	96.52
19LJ11-2#7	变质锆石	211.00	210.10	428.00	0.49	10.190 0	0.140 0	0.457 7	0.006 6	0.161 1	0.001 0	2.450	13	2.428	29	2.467	10	101.61
19LJ11-2#9	变质锆石	82.70	82.90	136.50	0.61	10.430 0	0.190 0	0.462 5	0.007 5	0.163 3	0.001 5	2.471	17	2.448	33	2.487	16	101.59
19LJ11-2#10	变质锆石	190.50	195.10	211.00	0.92	10.420 0	0.130 0	0.463 7	0.005 9	0.162 8	0.001 4	2.471	11	2.454	26	2.483	14	101.18
19LJ11-2#11	变质锆石	198.60	205.40	215.80	0.95	10.540 0	0.130 0	0.469 8	0.005 1	0.162 4	0.001 4	2.484	11	2.482	22	2.479	14	99.88
19LJ11-2#12	变质锆石	45.10	47.70	90.30	0.53	11.000 0	0.160 0	0.478 1	0.006 0	0.166 5	0.001 5	2.522	13	2.518	26	2.520	15	100.08
19LJ11-2#13	变质锆石	269.00	246.70	333.90	0.74	10.320 0	0.150 0	0.460 3	0.007 2	0.160 8	0.001 1	2.462	13	2.439	32	2.462	12	100.94
19LJ11-2#15	变质锆石	210.00	189.30	515.00	0.37	10.320 0	0.140 0	0.462 4	0.006 5	0.160 5	0.001 2	2.462	12	2.448	29	2.461	12	100.53
19LJ11-2#16	变质锆石	254.00	218.10	355.90	0.61	10.520 0	0.150 0	0.463 8	0.006 2	0.162 5	0.001 1	2.480	13	2.459	28	2.480	12	100.85
19LJ11-2#18	变质锆石	84.70	76.60	126.00	0.61	11.280 0	0.280 0	0.487 0	0.012 0	0.167 0	0.001 6	2.540	23	2.552	53	2.525	16	98.94
19LJ11-2#19	变质锆石	506.00	444.00	721.00	0.62	10.350 0	0.150 0	0.461 2	0.007 4	0.162 3	0.001 1	2.466	14	2.443	32	2.478	11	101.43
19LJ11-2#21	变质锆石	196.00	170.10	296.10	0.57	10.360 0	0.180 0	0.455 7	0.008 2	0.164 0	0.001 3	2.464	16	2.418	36	2.495	13	103.18
19LJ11-2#22	变质锆石	113.20	96.50	152.90	0.63	10.460 0	0.200 0	0.452 6	0.007 5	0.166 0	0.001 7	2.475	17	2.411	35	2.515	18	104.31
19LJ11-2#23	变质锆石	140.00	122.60	227.10	0.54	10.560 0	0.200 0	0.472 8	0.009 4	0.162 3	0.001 6	2.484	18	2.493	41	2.477	17	99.36
19LJ11-2#24	变质锆石	120.00	89.00	158.80	0.56	10.940 0	0.330 0	0.481 0	0.013 0	0.164 1	0.002 3	2.511	28	2.529	54	2.498	25	98.77
19LJ11-2#25	变质锆石	323.00	223.40	319.00	0.70	10.570 0	0.180 0	0.471 1	0.008 4	0.162 0	0.001 5	2.483	16	2.486	37	2.473	15	99.48
19LJ11-2#26	岩浆锆石	255.00	245.20	295.50	0.83	10.760 0	0.140 0	0.481 9	0.006 5	0.160 6	0.001 1	2.501	12	2.539	29	2.460	12	96.89
19LJ11-2#28	岩浆锆石	301.00	315.00	533.00	0.59	10.340 0	0.160 0	0.464 9	0.007 7	0.161 1	0.000 9	2.465	14	2.464	32	2.465	9.9	100.05
19LJ11-2#14	岩浆锆石	38.00	37.60	137.90	0.03	10.090 0	0.170 0	0.453 1	0.007 6	0.159 7	0.001 1	2.441	16	2.408	34	2.451	11	101.79
19LJ11-2#17	岩浆锆石	155.40	136.40	201.10	0.68	10.520 0	0.180 0	0.471 1	0.007 6	0.160 7	0.001 4	2.479	16	2.486	33	2.460	15	98.95
19LJ11-2#20	岩浆锆石	395.00	294.00	1385.00	0.21	11.040 0	0.220 0	0.491 0	0.011 0	0.162 6	0.001 9	2.524	18	2.574	48	2.480	20	96.35

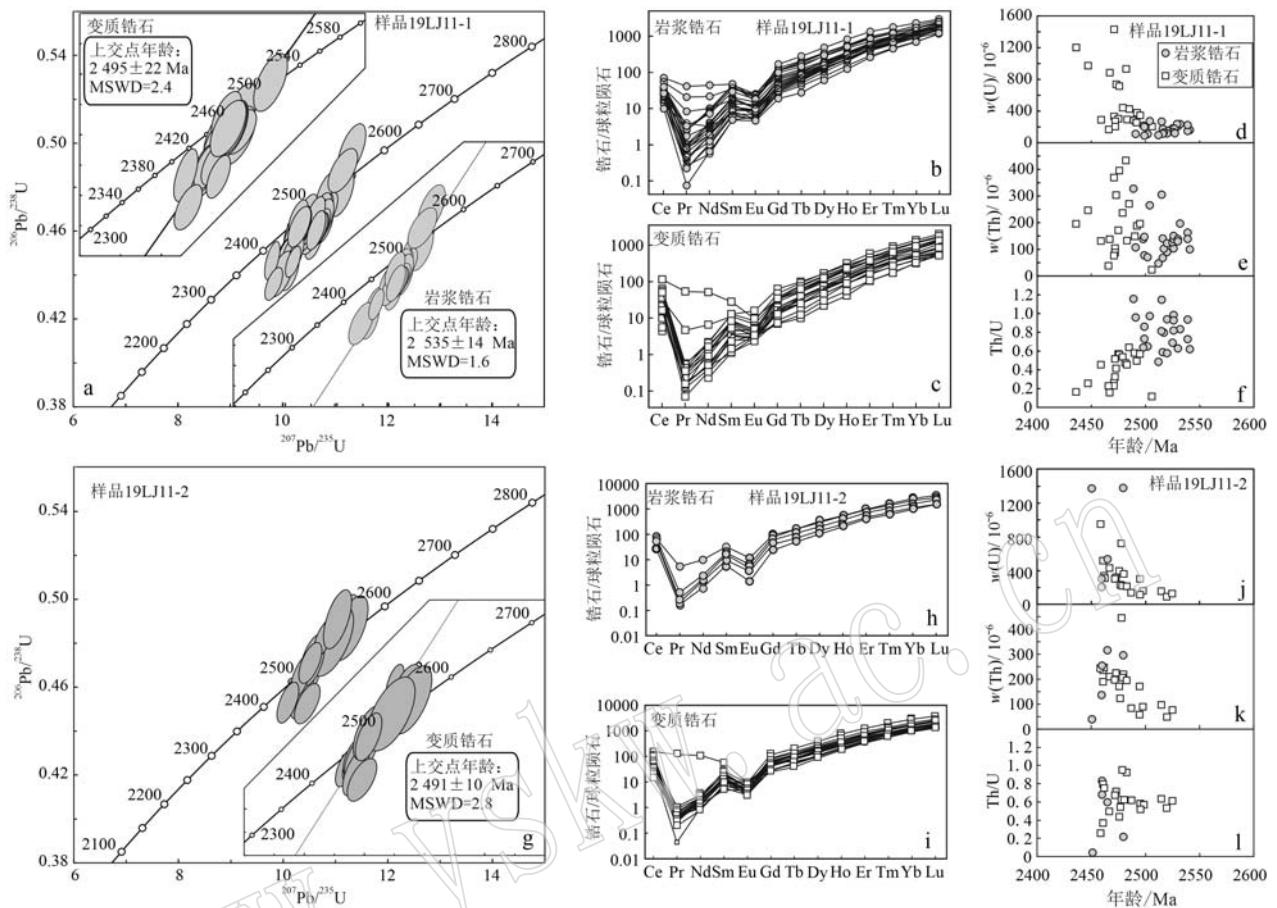


图 4 变质石英闪长岩(a~f)及淡色伟晶花岗岩(g~l)锆石 U-Pb 年龄谐和图、球粒陨石标准化稀土配分曲线(标准化值据 Sun and McDonough, 1989)、Th-U 含量及 Th/U 值图

Fig. 4 U-Pb Concordia diagrams, chondrite-normalized rare earth element (REE) patterns (normaoized data after Sun and McDonough, 1989), Th and U contents and Th/U ratios diagrams of zircons from the metamorphic quartz diorite sample (a~f) and leouc pegmatite granite sample (g~l)

锆石 Lu-Hf 同位素分析结果显示, 岩浆锆石域和变质重结晶域均具有极低的 $^{176}\text{Lu}/^{177}\text{Hf}$ 值(绝大部分小于 0.001), 而且, 尽管它们具有变化的单点 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄, 但它们具有较为统一的 $^{176}\text{Hf}/^{177}\text{Hf}$ 值, 分别在 0.281 432~0.281 289 和 0.281 405~0.281 329 之间(表 2; 图 5a), 以及一致的岩浆结晶年龄计算的 $\varepsilon\text{Hf}(t)$ 值, 分别在 +8.35~+2.22 之间(平均值 5.82)和 +7.10~+5.11 之间(平均值 6.07)(表 2; 图 5b), 两阶段亏损地幔模式年龄 $t_{\text{DM}}(\text{Hf}_2)$ 平均值分别为 2 721 Ma 和 2 707 Ma(表 2; 图 5c)。该样品的锆石 Hf 同位素分析结果显示, 岩浆锆石域和变质重结晶锆石域具有统一的 Hf 同位素成分, 表明样品中的岩浆锆石和变质重结晶锆石具有形成于同一岩浆源区的初始结晶锆石, 并经历了后期变质热事件的改造, 使它们的 U-Pb 体系

遭受扰动或重置, 但它们的 Hf 同位素未受到扰动。

4.2 含斜方辉石(角闪石) 淡色伟晶花岗岩(19LJ11-2)

含斜方辉石(角闪石) 淡色伟晶花岗岩样品(19LJ11-2)与变质石英闪长岩样品(19LJ11-1)相比, 具有更为复杂的锆石形态, 以它形、半自形柱状为主, 部分为椭球状, CL 图像显示大部分呈暗色, 内部结构及色调均匀, 不含具有振荡环的岩浆锆石核, 整个锆石颗粒已完全变质重结晶, 少量锆石具有残留的振荡环带锆石核和宽的暗色变质重结晶边(图 3b)。5 个具有振荡环带的锆石核分析给出的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄分布在 2 480~2 451 Ma 之间, 大部分分析点显示了不同程度的铅丢失(图 4g; 表 1); 20 个变质重结晶锆石域分析给出的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄分布在 2 525~2 459 Ma 之间(表 1), 其加权平

表 2 吉南和龙地区变质石英闪长岩及淡色伟晶花岗岩锆石 Hf 同位素数据表
 Table 2 Hf isotopes analytical data for the zircons from the the metamorphic quartz diorite and leucoc pegmatite granite samples collected from the Helong region in the southern Jilin Province

分析点	锆石类型	单点年 龄/Ma	岩浆结晶 年龄/Ma	$^{176}\text{Yb}/^{177}\text{Hf}$	2σ	$^{176}\text{Lu}/^{177}\text{Hf}$	2σ	$^{176}\text{Hf}/^{177}\text{Hf}$	2σ	$\varepsilon\text{Hf}(0)$	2σ	$\varepsilon\text{Hf}(\tau)$	2σ	$t_{\text{DM}}^{\text{C}}(\text{Hf}_2)$ /Ma	$t_{\text{DM}}^{\text{C}}(\text{Hf})$ /Ma	$t_{\text{DM}}^{\text{C}}(\text{Hf}_2)$ 2σ	$t_{\text{DM}}^{\text{C}}(\text{Hf})$ 2σ	$f_{\text{Lu/Hf}}$	
变质石英闪长岩(19JJ11-1)																			
19JJ11-1-01	岩浆锆石	2.539	2.535	0.031 880	0.000 234	0.000 685	0.000 002	0.281 358	0.000 025	0.281 324	-50.02	0.88	5.74	0.89	2.622	2.735	34	2.676	-0.98
19JJ11-1-02	岩浆锆石	2.516	2.535	0.037 859	0.000 305	0.000 833	0.000 004	0.281 340	0.000 020	0.281 300	-50.63	0.71	4.87	0.71	2.656	2.811	27	2.729	-0.97
19JJ11-1-03	岩浆锆石	2.539	2.535	0.034 559	0.000 452	0.000 862	0.000 008	0.281 360	0.000 024	0.281 318	-49.93	0.85	5.53	0.85	2.631	2.753	32	2.689	-0.97
19JJ11-1-06	岩浆锆石	2.516	2.535	0.026 429	0.000 126	0.000 725	0.000 004	0.281 395	0.000 020	0.281 360	-48.69	0.69	7.01	0.70	2.574	2.625	27	2.598	-0.98
19JJ11-1-07	岩浆锆石	2.541	2.535	0.025 541	0.000 275	0.000 715	0.000 010	0.281 432	0.000 022	0.281 398	-47.38	0.78	8.35	0.79	2.523	2.508	30	2.516	-0.98
19JJ11-1-11	岩浆锆石	2.526	2.535	0.021 732	0.000 117	0.000 553	0.000 003	0.281 345	0.000 019	0.281 318	-50.46	0.68	5.52	0.68	2.630	2.754	26	2.689	-0.98
19JJ11-1-12	岩浆锆石	2.526	2.535	0.021 827	0.000 095	0.000 570	0.000 001	0.281 334	0.000 018	0.281 306	-50.87	0.65	5.09	0.66	2.647	2.792	25	2.716	-0.98
19JJ11-1-14	岩浆锆石	2.491	2.535	0.054 024	0.001 212	0.001 314	0.000 025	0.281 289	0.000 024	0.281 225	-52.45	0.85	2.22	0.85	2.760	3.041	33	2.891	-0.96
19JJ11-1-16	岩浆锆石	2.499	2.535	0.021 405	0.000 059	0.000 582	0.000 001	0.281 362	0.000 022	0.281 334	-49.86	0.77	6.08	0.77	2.609	2.705	29	2.655	-0.98
19JJ11-1-18	岩浆锆石	2.520	2.535	0.025 022	0.000 225	0.000 717	0.000 005	0.281 378	0.000 020	0.281 343	-49.29	0.71	6.42	0.71	2.597	2.676	27	2.634	-0.98
19JJ11-1-19	岩浆锆石	2.504	2.535	0.033 538	0.000 215	0.000 878	0.000 005	0.281 403	0.000 021	0.281 361	-48.40	0.74	7.04	0.75	2.573	2.622	29	2.596	-0.97
19JJ11-1-21	岩浆锆石	2.520	2.535	0.037 207	0.000 257	0.000 957	0.000 007	0.281 419	0.000 020	0.281 373	-47.85	0.72	7.46	0.72	2.557	2.585	28	2.571	-0.97
19JJ11-1-23	岩浆锆石	2.526	2.535	0.018 435	0.000 176	0.000 482	0.000 003	0.281 315	0.000 019	0.281 292	-51.51	0.69	4.59	0.69	2.665	2.835	26	2.746	-0.99
19JJ11-1-25	岩浆锆石	2.517	2.535	0.025 115	0.000 277	0.000 655	0.000 006	0.281 352	0.000 023	0.281 320	-50.23	0.80	5.58	0.80	2.628	2.749	30	2.686	-0.98
19JJ11-1-29	岩浆锆石	2.530	2.535	0.023 724	0.000 107	0.000 715	0.000 004	0.281 393	0.000 021	0.281 358	-48.76	0.74	6.95	0.74	2.576	2.629	28	2.602	-0.98
19JJ11-1-04	变质锆石	2.459	2.535	0.020 192	0.000 155	0.000 517	0.000 006	0.281 377	0.000 018	0.281 352	-49.33	0.64	6.72	0.64	2.585	2.650	24	2.616	-0.98
19JJ11-1-05	变质锆石	2.472	2.535	0.019 990	0.000 053	0.000 558	0.000 003	0.281 348	0.000 019	0.281 321	-50.37	0.67	5.61	0.67	2.627	2.747	25	2.684	-0.98
19JJ11-1-08	变质锆石	2.467	2.535	0.016 770	0.000 066	0.000 455	0.000 001	0.281 385	0.000 018	0.281 363	-49.06	0.64	7.10	0.64	2.570	2.617	24	2.593	-0.99
19JJ11-1-09	变质锆石	2.447	2.535	0.024 263	0.000 123	0.000 647	0.000 003	0.281 338	0.000 021	0.281 307	-50.72	0.74	5.11	0.74	2.646	2.790	28	2.715	-0.98
19JJ11-1-10	变质锆石	2.495	2.535	0.013 160	0.000 054	0.000 353	0.000 000	0.281 335	0.000 017	0.281 338	-50.11	0.60	6.23	0.60	2.603	2.693	23	2.646	-0.99
19JJ11-1-13	变质锆石	2.472	2.535	0.008 039	0.000 033	0.000 266	0.000 001	0.281 340	0.000 020	0.281 327	-50.66	0.72	5.83	0.72	2.618	2.727	27	2.671	-0.99
19JJ11-1-15	变质锆石	2.473	2.535	0.010 181	0.000 052	0.000 281	0.000 001	0.281 332	0.000 019	0.281 318	-50.93	0.68	5.53	0.68	2.630	2.754	26	2.689	-0.99
19JJ11-1-17	变质锆石	2.471	2.535	0.011 190	0.000 045	0.000 326	0.000 001	0.281 342	0.000 021	0.281 326	-50.58	0.74	5.80	0.75	2.620	2.730	28	2.673	-0.99
19JJ11-1-20	变质锆石	2.475	2.535	0.009 347	0.000 017	0.000 264	0.000 001	0.281 329	0.000 019	0.281 317	-51.02	0.66	5.46	0.67	2.632	2.759	25	2.693	-0.99
19JJ11-1-22	变质锆石	2.483	2.535	0.018 185	0.000 088	0.000 527	0.000 005	0.281 357	0.000 018	0.281 331	-50.05	0.65	5.98	0.65	2.613	2.714	25	2.661	-0.98
19JJ11-1-24	变质锆石	2.482	2.535	0.021 055	0.000 045	0.000 591	0.000 002	0.281 360	0.000 017	0.281 331	-49.94	0.61	5.98	0.61	2.613	2.714	23	2.661	-0.98
19JJ11-1-26	变质锆石	2.492	2.535	0.032 193	0.000 513	0.000 948	0.000 015	0.281 405	0.000 021	0.281 359	-48.35	0.75	6.96	0.75	2.576	2.628	29	2.601	-0.97
19JJ11-1-27	变质锆石	2.490	2.535	0.015 687	0.000 050	0.000 442	0.000 001	0.281 348	0.000 019	0.281 327	-50.36	0.68	5.82	0.68	2.619	2.728	26	2.671	-0.99
19JJ11-1-28	变质锆石	2.476	2.535	0.011 854	0.000 072	0.000 323	0.000 003	0.281 368	0.000 018	0.281 352	-49.65	0.63	6.74	0.63	2.584	2.648	24	2.615	-0.99

续表 2

Continued Table 2

分析点	锆石类型	单点年 龄 Ma	岩浆结晶 年龄 Ma	$^{176}\text{Yb}/^{177}\text{Hf}$	2σ	$^{176}\text{Lu}/^{177}\text{Hf}$	2σ	$^{176}\text{Hf}/^{177}\text{Hf}$	2σ	$\varepsilon\text{Hf}(0)$	2σ	$\varepsilon\text{Hf}(\iota)$	2σ	t_{DM}	$t_{\text{DM}}(\text{Hf}_2)$	2σ	t_{DM}^{C}	$t_{\text{DM}}^{\text{C}}(\text{Hf})$	2σ	$f_{\text{Lu/Hf}}$
														年龄 Ma	$/\text{Ma}$	$/\text{Ma}$	$/\text{Ma}$	$/\text{Ma}$	$/\text{Ma}$	
淡色伟晶花岗岩(19J11-2)																				
19J11-2-04	岩浆锆石	2.460	2.535	0.018 282	0.000 104	0.000 585	0.000 004	0.281 344	0.000 016	0.281 315	-50.51	0.55	5.42	0.56	2.634	2.763	21	2.695	-0.98	
19J11-2-08	岩浆锆石	2.465	2.535	0.014 450	0.000 072	0.000 456	0.000 002	0.281 329	0.000 023	0.281 307	-51.03	0.80	5.12	0.81	2.645	2.789	30	2.714	-0.99	
19J11-2-14	岩浆锆石	2.451	2.535	0.018 537	0.000 229	0.000 611	0.000 007	0.281 393	0.000 020	0.281 363	-48.77	0.72	7.13	0.72	2.569	2.614	27	2.591	-0.98	
19J11-2-17	岩浆锆石	2.460	2.535	0.013 867	0.000 121	0.000 470	0.000 004	0.281 350	0.000 017	0.281 327	-50.28	0.60	5.85	0.61	2.618	2.725	23	2.669	-0.99	
19J11-2-20	岩浆锆石	2.480	2.535	0.013 769	0.000 192	0.000 401	0.000 005	0.281 362	0.000 019	0.281 342	-49.87	0.68	6.38	0.69	2.598	2.679	26	2.637	-0.99	
19J11-2-01	变质锆石	2.461	2.535	0.021 887	0.000 106	0.000 683	0.000 003	0.281 262	0.000 015	0.281 229	-53.39	0.53	2.36	0.53	2.751	3.029	20	2.883	-0.98	
19J11-2-02	变质锆石	2.459	2.535	0.016 517	0.000 199	0.000 551	0.000 010	0.281 342	0.000 018	0.281 316	-50.56	0.65	5.43	0.66	2.634	2.762	25	2.695	-0.98	
19J11-2-03	变质锆石	2.476	2.535	0.013 608	0.000 056	0.000 446	0.000 001	0.281 284	0.000 015	0.281 263	-52.61	0.54	3.55	0.54	2.704	2.925	21	2.810	-0.99	
19J11-2-05	变质锆石	2.472	2.535	0.021 544	0.000 296	0.000 688	0.000 007	0.281 298	0.000 018	0.281 265	-52.11	0.64	3.63	0.64	2.703	2.918	24	2.805	-0.98	
19J11-2-06	变质锆石	2.495	2.535	0.035 686	0.000 358	0.001 089	0.000 009	0.281 308	0.000 015	0.281 256	-51.76	0.54	3.30	0.55	2.717	2.947	21	2.825	-0.97	
19J11-2-07	变质锆石	2.467	2.535	0.027 696	0.000 546	0.000 857	0.000 015	0.281 273	0.000 015	0.281 232	-53.00	0.53	2.45	0.53	2.748	3.021	20	2.877	-0.97	
19J11-2-09	变质锆石	2.487	2.535	0.013 995	0.000 067	0.000 445	0.000 003	0.281 296	0.000 018	0.281 274	-52.20	0.63	3.96	0.64	2.689	2.890	24	2.785	-0.99	
19J11-2-10	变质锆石	2.483	2.535	0.029 228	0.000 140	0.000 910	0.000 007	0.281 324	0.000 019	0.281 279	-51.22	0.69	4.14	0.69	2.684	2.874	26	2.774	-0.97	
19J11-2-11	变质锆石	2.479	2.535	0.019 168	0.000 197	0.000 585	0.000 005	0.281 324	0.000 017	0.281 296	-51.21	0.60	4.72	0.60	2.661	2.824	23	2.738	-0.98	
19J11-2-12	变质锆石	2.520	2.535	0.015 756	0.000 206	0.000 501	0.000 008	0.281 299	0.000 015	0.281 274	-52.10	0.53	3.96	0.54	2.689	2.889	20	2.785	-0.98	
19J11-2-13	变质锆石	2.462	2.535	0.015 711	0.000 164	0.000 469	0.000 003	0.281 274	0.000 016	0.281 251	-52.98	0.57	3.14	0.57	2.720	2.961	22	2.835	-0.99	
19J11-2-15	变质锆石	2.461	2.535	0.016 877	0.000 043	0.000 505	0.000 001	0.281 328	0.000 018	0.281 303	-51.07	0.62	5.00	0.63	2.650	2.800	24	2.722	-0.98	
19J11-2-16	变质锆石	2.480	2.535	0.019 183	0.000 078	0.000 589	0.000 002	0.281 301	0.000 017	0.281 272	-52.02	0.59	3.89	0.60	2.692	2.895	23	2.789	-0.98	
19J11-2-18	变质锆石	2.525	2.535	0.031 547	0.000 081	0.000 880	0.000 002	0.281 309	0.000 018	0.281 267	-51.73	0.62	3.69	0.63	2.701	2.913	24	2.801	-0.97	

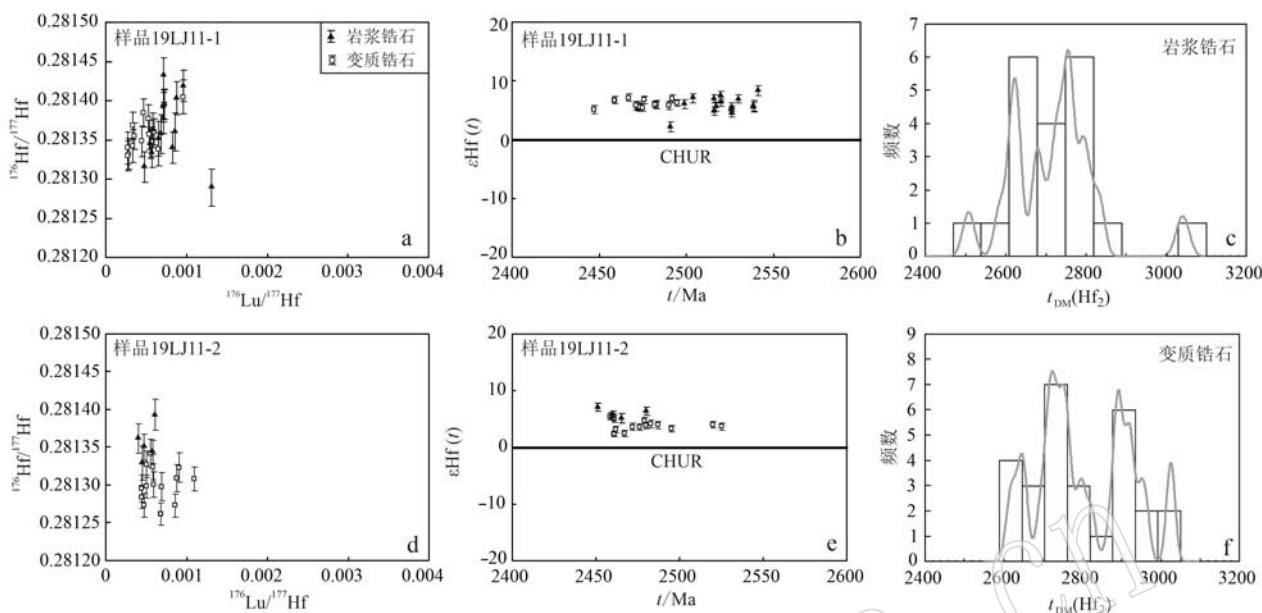


图 5 变质石英闪长岩(a~c)及淡色伟晶花岗岩(d~f)锆石 Hf 同位素成分

Fig. 5 Results of in situ Hf isotope analyses of zircons from the metamorphic quartz diorite sample (a~c) and leucogranite pegmatite sample (d~f)

a, d— $^{176}\text{Lu}/^{177}\text{Hf}$ vs. $^{176}\text{Hf}/^{177}\text{Hf}$ 图; b, e—单点 U-Pb 年龄 vs. $\varepsilon\text{Hf}(t)$ 值; c, f—两阶段 Hf 模式年龄直方图; CHUR—球粒陨石均一储库
 a, d— $^{176}\text{Lu}/^{177}\text{Hf}$ vs. $^{176}\text{Hf}/^{177}\text{Hf}$; b, e—U-Pb ages vs $\varepsilon\text{Hf}(t)$ values; c, f—two-stage Hf model ages histograms; CHUR—chondritic uniform reservoir

均年龄为 2478 ± 8 Ma (MSWD = 7.4, $n = 20$)，上交点年龄为 2491 ± 10 Ma (MSWD = 2.8, 图 4g)，与变质石英闪长岩样品 (19LJ11-1) 给出的变质年龄一致，并且其岩浆锆石与变质重结晶锆石同样显示了类似的球粒陨石标准化稀土元素配分曲线模式 (图 4h, 4i)，以及类似的锆石微区 U 和 Th 含量变化范围及 Th/U 值 (均大于 0.2) (图 4j~4l)。

锆石 Lu-Hf 同位素分析结果显示，它们具有极低的 $^{176}\text{Lu}/^{177}\text{Hf}$ 值 (绝大部分小于 0.001)，统一的 $^{176}\text{Hf}/^{177}\text{Hf}$ 值和 $\varepsilon\text{Hf}(t)$ 值，分别在 0.281 393 ~ 0.281 262 之间和 +7.13 ~ +2.36 之间 (平均值 +4.41) (表 2; 图 5d, 5e)，两阶段亏损地幔模式年龄 $t_{\text{DM}}(\text{Hf}_2)$ 在 3 029 ~ 2 614 Ma 之间变化，平均值为 2 821 Ma (表 2, 图 5f)。该样品显示了与变质石英闪长岩样品类似的锆石 Hf 同位素成分 (表 2; 图 5)。

5 讨论

5.1 变质深熔作用时间

吉南地区太古宙基底普遍经历了角闪岩相到麻粒岩相的高级变质作用，变质深熔作用普遍而强

烈，因而太古宙基底样品中的原始岩浆锆石往往遭受了变质深熔热事件的改造，具有复杂的形态和内部结构，一些锆石的 CL 图像显示它们原始的岩浆结晶环带部分或完全被修改 (图 3)，它们的 U-Pb 体系相应地受到扰动或完全重置，导致单个样品往往具有变化的、宽范围的锆石 U-Pb 年龄以及不同程度的 Pb 丢失。本文研究的变质石英闪长岩样品 (19LJ11-1) 中的锆石主要可以分为两类：一类是具有典型岩浆锆石结晶环带的锆石 (图 3a)，另一类是内部结构均匀具有典型变质重结晶锆石特征的锆石，它们通常形成岩浆锆石变质边或幔 (图 3a)，或整个颗粒完全变质重结晶，锆石 Hf 同位素分析结果显示变质重结晶锆石与岩浆锆石具有统一的 Hf 同位素组成 (图 5a, 5b; 表 2)，表明该样品中的岩浆锆石及变质重结晶锆石具有形成于同一岩浆源区的初始结晶锆石，并经历了后期变质深熔热事件的改造或变质重结晶，其岩浆锆石 U-Pb 体系遭受了不同程度扰动， $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄在 2 541 ~ 2 489 Ma 之间变化 (表 1)，其给出的加权平均年龄为 2519 ± 6 Ma (MSWD = 3.4, $n = 24$)，而上交点年龄为 2535 ± 14 Ma (MSWD = 1.6, 图 4a)，因此，将上交点年龄

解释为变质石英闪长岩原始岩浆结晶的年龄；而变质重结晶锆石的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄在2 506~2 436 Ma之间变化，大部分分析点显示了不同程度的铅丢失（图4a；表1），它们具有高的Th/U值（均>0.1），具有与岩浆锆石类似的锆石微区球粒陨石标准化稀土元素配分曲线及Hf同位素成分（图4b~4f, 图5），表明其变质重结晶可能发生在变质作用峰期较短时间内高温压条件下（Vavra *et al.*, 1999; Whitehouse and Kamber, 2002; Harley and Kelly, 2007; Harley *et al.*, 2007），其加权平均年龄为2 476±7 Ma (MSWD=5.3, n=20)，上交点年龄为2 495±22 Ma (MSWD=2.4, 图4a)，笔者将其上交点年龄2 495±22 Ma (MSWD=2.4, 图4a)解释为变质重结晶年龄。淡色伟晶花岗岩样品(19LJ11-2)大部分为变质重结晶锆石，具有振荡环带的岩浆锆石核通常被宽厚的变质重结晶锆石所包围（图3b），指示锆石遭受了更强烈的改造或重置，其5个岩浆锆石域的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄分布在2 480~2 451 Ma之间，显示了较差的谐和性，20个变质重结晶锆石的 $^{207}\text{Pb}/^{206}\text{Pb}$ 年龄在2 525~2 459 Ma之间变化（表1），它们与变质石英闪长岩样品的变质重结晶锆石一样，具有高的Th/U值（均>0.1），与岩浆锆石类似的锆石微区球粒陨石标准化稀土元素配分曲线及Hf同位素组成（图4h~4l, 图5），表明其经历了在较短时间内高温压条件下发生的变质重结晶，其加权平均年龄为2 478±8 Ma (MSWD=7.4, n=20)，上交点年龄为2 491±10 Ma (MSWD=2.8, 图4g)，后者代表其变质重结晶的年龄。变质石英闪长岩样品和淡色伟晶花岗岩样品中变质重结晶锆石给出了在误差范围内统一的上交点年龄2 495±22 Ma (MSWD=2.4)和2 491±10 Ma (MSWD=2.8)，表明研究区变质深熔作用发生的时间为约~2.5 Ga，与吉南地区以及华北克拉通其它地区新太古代末变质深熔事件年龄一致（Zhao *et al.*, 1999, 2005; Guo *et al.*, 2016, 2017, 2018; Yu *et al.*, 2021）。

5.2 石英闪长岩原地深熔作用及深熔机制

研究区太古宙基底广泛分布的长英质浅色体及淡色花岗岩通常被认为是地壳深熔作用的产物，要么是寄主（或围岩）岩石的原地熔融，要么是外来深熔熔体的注入。野外露头观察显示，深熔作用形成的长英质浅色体在变质石英闪长岩中呈不规则网脉、与剪切面理平行的细脉，并与暗色变质石英闪长岩构成浅色和暗色相间分布的条带状构造，花岗

质团块状或薄迷雾状分布在变质石英闪长岩中，具有混合岩化的特征（图2b）；此外，一些长英质浅色体在变质石英闪长岩内的破裂面分布，粒度明显变粗（图2c），随着长英质浅色体的增加，呈伟晶花岗岩脉出露，内部含有斜方辉石和角闪石巨晶集合体，并且它们与暗色变质石英闪长质围岩呈渐变关系（图2a, 2d）。岩相学分析显示，暗色变质石英闪长岩中的斜长石、角闪石及黑云母岩相学特征均指示它们经历了部分熔融作用（图2e~2g），这些早期残余矿物之间分布有以钾长（微斜长石）、斜长石及石英为主的浅色矿物（图2g），代表变质石英闪长岩部分熔融熔体的结晶矿物。此外，变质石英闪长岩（样品19LJ11-1）与淡色伟晶花岗岩（样品19LJ11-2）具有相同的岩浆锆石和变质重结晶锆石U-Pb年龄信息（图4a, 4g；表1）；并且，无论岩浆锆石还是变质重结晶锆石，它们的U、Th、微量元素含量及Th/U值变化范围均分别可以对比（图4），尤其它们具有统一的锆石Hf同位素组成（图5；表2），表明淡色伟晶花岗岩源自变质石英闪长岩的部分熔融。因此，根据变质石英闪长岩与长英质浅色体及淡色花岗岩的野外地质特征、相互关系、岩相学特征及锆石U-Pb年代学、微量元素及Hf同位素的精细对比分析，推断变质石英闪长岩中的长英质浅色体及含斜方辉石（角闪石）的淡色伟晶花岗岩均源自石英闪长岩的原地部分熔融。

变质石英闪长岩及含斜方辉石角闪石淡色伟晶花岗岩的岩相学分析显示，变质石英闪长岩中参与部分熔融的矿物为角闪石、黑云母、斜长石及石英（图2e~2g），代表熔融产物的淡色花岗岩的主要矿物组合为钾长石、斜方辉石（转熔矿物）、石英及少量斜长石（图2h, 2i），熔融作用形成的长英质浅色体及淡色花岗岩主要分布在张性构造中（如裂隙或破裂面），淡色伟晶花岗岩呈块状构造，矿物无优选方位，结合变质石英闪长岩及淡色伟晶花岗岩中的矿物组合，指示其部分熔融的压力较低；根据变质基性岩及安山质闪长岩的熔融实验及相关研究（Beard and Lofgren, 1991; Wolf and Wyllie, 1994; Aranovich *et al.*, 2014; Newton *et al.*, 2014），在石英闪长岩的部分熔融过程中黑云母及角闪石等含水矿物发生脱水形成花岗质熔体，其部分熔融反应可能为：黑云母+角闪石+斜长石+石英=斜方辉石+钾长石+熔体。野外地质特征表明，熔体沿裂隙、剪切面理，张性破裂带分凝聚集并结晶，形成长英质浅色体或淡

色花岗岩脉体或岩块(图2b, 2c), 随着熔体的不断增加, 彼此独立的熔体将会相互联通、聚集并增温, 形成淡色伟晶花岗岩, 并在结晶冷却过程中伴有转熔矿物斜方辉石与熔体反应形成角闪石(Brown, 2007, 2013; Clemens and Stevens, 2016; 图2a, 2d)。华北克拉通在新太古代末约2.55~2.50 Ga期间经历了准同期的巨量花岗质(TTG)岩石的形成和角闪岩相-麻粒岩相变质作用, 麻粒岩相变质作用具有逆时针p-T轨迹, 被认为可能与幔源岩浆的底侵有关(Zhao et al., 1999, 2005; 耿元生等, 2010)。本文研究的吉南地区, 太古宙基底岩石在新太古代末经历了角闪岩相-麻粒岩相变质作用, 麻粒岩相变质作用具有逆时针p-T轨迹(葛文春等, 1993; Ge et al., 2003), 紫苏花岗岩及深熔作用形成的淡色钾质花岗岩广泛分布(葛文春等, 1996a, 1996b; 孙德有等, 1998; Guo et al., 2016, 2017, 2018; Yu et al., 2021)。因此, 吉南地区新太古代末约2.5 Ga广泛发生的变质深熔作用同华北克拉通其它地区一样, 可能是地幔物质底侵作用的结果。新太古代末-古元古代早期被认为是华北克拉通构造体制转换的关键时期(Zhai and Santosh, 2011; Zhao and Zhai, 2013), 华北克拉通东部块体新太古代末普遍经历了具有逆时针p-T轨迹的变质作用和深熔作用(贺同兴, 1987; Zhao et al., 1999, 2005; Zhai et al., 2005; 程裕淇等, 2000, 2001, 2004; 任留东等, 2011; Dong et al., 2017; Wang et al., 2017; Jiang et al., 2021; Yu et al., 2021), 可能反映了华北克拉通东部陆块在形成稳定克拉通过程中地幔物质上涌的热效应。

6 结论

(1) 野外及岩相学特征表明, 石英闪长岩经历了变质深熔作用, 其部分熔融反应为: 黑云母+角闪石+斜长石+石英 = 斜方辉石+钾长石+熔体, 并伴有转熔矿物斜方辉石与熔体反应形成角闪石。

(2) 变质石英闪长岩与长英质浅色体及淡色伟晶花岗岩的野外地质特征、相互关系、岩相学特征及锆石U-Pb年代学、微量元素及Hf同位素的精细对比分析, 表明变质石英闪长岩中的长英质浅色体及淡色伟晶花岗岩均源自石英闪长岩的原地部分熔融。

(3) 变质深熔作用发生的时间为~2.5 Ga, 与

吉南地区以及华北克拉通其它地区新太古代末变质深熔事件年龄一致, 可能反映了华北克拉通东部陆块在形成稳定克拉通过程中地幔物质上涌的热效应。

致谢 在野外考察中, 与香港大学的赵国春教授、北京大学的魏春景教授及吉林大学的葛文春教授开展了交流讨论, 使作者受益匪浅, 在此表示衷心的感谢! 周喜文研究员在野外和成文过程中提供了帮助及交流探讨, 中国地质科学院矿产资源研究所郭春丽研究员在锆石Hf同位素测试中提供了帮助, 论文审稿人给出了良好的修改建议, 在此一并表示衷心的感谢!

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