

阿拉善地块东部早前寒武纪变质基底性质及归属

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摘要: 阿拉善地块是一个被造山带和断裂带围限的前寒武纪微陆块, 其早前寒武纪变质基底的性质和归属一直存在争议。本文根据笔者最近在阿拉善地块东部地区获取的研究资料, 结合阿拉善地块近年来的研究进展, 对阿拉善地块东部早前寒武纪变质基底的性质和归属进行了讨论, 认为阿拉善地块东部地区出露的 2.64 Ga 的斜长角闪岩和 ~2.5 Ga 的 TTG 片麻岩, 明确指示阿拉善地块东部也存在新太古代变质基底。这些新太古代变质岩石的锆石 Hf 同位素模式年龄主要集中在 3.38~2.88 Ga 之间, 与阿拉善地块东部的叠布斯格岩群和巴彦乌拉山岩群中古元古代锆石的 Hf 模式年龄 (3.4~2.6 Ga) 一致, 表明阿拉善地块东部经历了古太古代晚期—新太古代早期的陆壳生长事件和 ~2.5 Ga 的岩浆-变质事件。此外, 阿拉善地块东部的早前寒武纪变质基底还记录了 2.3~2.0 Ga 的岩浆事件和 ~1.95 Ga、~1.85 Ga 的岩浆-变质事件, 与阿拉善地块西部和华北克拉通变质基底具有相似的地壳演化历史。对比分析锆石年代学数据和 Hf 同位素组成的结果表明, 阿拉善地块东部早前寒武纪变质基底所记录的古元古代中晚期岩浆-变质事件与孔兹岩带极为相似, 指示阿拉善地块很有可能在古元古代时期作为孔兹岩带的西延部分卷入到了这一古元古代造山带中。

关键词: 阿拉善地块东部; 新太古代岩石; 古元古代; 构造归属

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The property and tectonic affinity of the Early Precambrian metamorphic basements in the eastern Alxa block

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Abstract: The Alxa block is a Precambrian microcontinent bounded by orogenic belts and fault zones, and its property and tectonic affinity have always been controversial. Based on the recent research data obtained by the author in the eastern Alxa block and combined with the research progress of Alxa block in recent years, it is considered that the 2.64 Ga amphibolite and ~2.5 Ga TTG gneiss exposed in the eastern Alxa block, and clearly indicates that Neoarchean metamorphic basement also exists in the eastern Alxa block. The zircon Hf isotopic data of these Neoarchean metamorphic rocks show that the model ages mainly range from 3.38 Ga to 2.88 Ga, which are consistent with the Hf model ages (3.4~2.6 Ga) of the Paleoproterozoic zircons from the Diebusige Complex and Bayanwulashan Complex in the eastern Alxa block. The above data indicate that the eastern Alxa block has experienced crustal growth from the Late Paleoarchean to Early Neoarchean and magmatic-metamorphic event of the ~2.5 Ga. In addition, the magmatic event of 2.3~2.0 Ga and magmatic-metamorphic events of ~1.95 Ga and ~1.85 Ga were

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also recorded in the Early Precambrian basements of the eastern Alxa block, similar to that of the western Alxa block and the North China Craton. Based on the similarities in the zircon chronology data and Hf isotopic compositions of the Early Precambrian basements from the Alxa block and the Khondalite belt, we proposed that the Alxa block has a similar crustal evolutionary history with the Khondalite belt, and was probably involved in the Khondalite belt as its western extension in the Paleoproterozoic.

Key words: eastern Alxa block; Archean rocks; Paleoproterozoic; tectonic affinity

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阿拉善地块在大地构造位置上位于华北克拉通、塔里木克拉通和古亚洲洋构造域的接壤位置,传统观点认为阿拉善地块是华北克拉通的组成部分(Huang, 1945; 任纪舜等, 1980),是研究地壳早期构造演化的重要地区。近年来,不同学者对阿拉善地块开展了大量调查研究,并在基底物质组成、造山带演化和构造变形、沉积学、古生物地层及成矿作用等方面取得了一些重要进展(耿元生等, 2006, 2007; Feng et al., 2013; Zhang J et al., 2013; 张进等, 2013; Dan et al., 2014; 谢力等, 2014; 张建新等, 2018; 刘平华等, 2019; 牛鹏飞等, 2019)。但目前有关阿拉善地块东部太古宙变质基底的研究资料还相对比较匮乏,仍然存在一些关键性问题尚未解决,比如阿拉善地块内太古宙岩石的出露情况和阿拉善地块的归属问题。

宫江华等(2012)和Zhang J X 等(2013)在阿拉善地块西部的北大山地区首次报道了2.5 Ga TTG片麻岩的存在,解决了长期以来阿拉善地块是否存在太古宙岩石的争议问题。但阿拉善地块东部是否存在太古宙岩石仍存在疑问。根据耿元生等(2006, 2007)提出的划分方案,叠布斯格岩群、巴彦乌拉山岩群和波罗斯坦庙岩群共同组成了阿拉善地块东部的一套新太古代-古元古代角闪岩相-麻粒岩相变质岩系,陈志勇等(2004)在对阿拉善地块东部的前寒武纪变质岩系进行重新划分时也将叠布斯格岩群的时代划归中太古代,认为阿拉善地块东部存在太古宙变质基底,但没有提出明确的年代学证据。李俊健(2006)根据在叠布斯格岩群中获得的 $3\ 018\pm49$ Ma的斜长角闪岩 Sm-Nd 模式年龄,也认为阿拉善地块存在太古宙古老地壳。而另外一些学者所获得的年代学数据表明,虽然锆石 Hf 同位素数据结果不排除深部存在太古宙基底的可能性,但不支持阿拉善地块内有太古宙岩石的出露(周红英等, 2007; 董

春艳等, 2007; Dan et al., 2012a; Wu et al., 2014)。

有关阿拉善地块的构造归属问题,不同学者根据基底组成及其记录的构造热事件提出了不同的观点,主要包括:①阿拉善地块是华北克拉通的组成部分:认为阿拉善地块自新太古代(~2.5 Ga)以来就是华北克拉通的一部分(任纪舜等, 1980; 张振法等, 1997; 伍家善等, 1998; 万渝生等, 2003; Zhai et al., 2005),并作为新太古代花岗绿岩带成为了华北克拉通的太古宙微陆块之一(杨振德等, 1988; 翟明国等, 2000; Zhai and Santosh, 2011);或是阴山地块的西延部分,同阴山地块一起在~1.95 Ga与鄂尔多斯地块碰撞形成华北克拉通西部陆块(Zhao et al., 2005 及相关参考文献);或是华北克拉通西部陆块古元古代造山带的一部分,即孔兹岩带的一部分(董春艳等, 2007; 耿元生等, 2010a; 宫江华等, 2012; Zhang J X et al., 2013; Dong et al., 2017);②阿拉善地块是独立发展的微陆块:认为阿拉善地块是独立的古元古代地体,既不属于阴山地块的西延,也不属于孔兹岩带的一部分(李俊健, 2006; Dan et al., 2012a);③阿拉善地块亲扬子或塔里木克拉通:一些学者通过古地磁学、年代学、古生物学和岩浆事件研究认为阿拉善地块与扬子克拉通或塔里木克拉通具有亲缘性(葛肖虹等, 2000; 李献华等, 2004; 段吉业等, 2005; 李锦轶等, 2009; Dan et al., 2014; Yuan and Yang, 2015a; Zhang B H et al., 2016; Zhang S H et al., 2016; Zhang J et al., 2016; Song et al., 2017),并且在古生代才与华北克拉通拼合(李锦轶等, 2012; Zhang J et al., 2013; Yuan and Yang, 2015b; Dan et al., 2016)。由此可见,前人对于阿拉善地块归属的认识存在较大的分歧,并且阿拉善地块归属的争议归根结底最大的矛盾点还是阿拉善地块与华北克拉通之间的关系。

随着研究的深入，在变质基底岩石中获得的大量锆石 U-Pb 年代学和相应的 Hf 同位素数据，为探讨阿拉善地块早前寒武纪的性质和归属提供了科学依据 (Zhao *et al.*, 2001; 耿元生等, 2006, 2007, 2010a, 2010b; 宫江华等, 2011; Dan *et al.*, 2012a, 2012b, 2016; Zhang J X *et al.*, 2013; Wu *et al.*, 2014)。本文将结合近年来该地区的研究成果和前人研究数据，综合探讨以下两个问题：① 阿拉善地块东部是否出露太古宙岩石；② 阿拉善地块早前寒武纪变质基底与华北克拉通同时代的变质基底具有怎样的关系。

1 地质背景

华北克拉通作为全球最古老的克拉通之一，具有 3.8 Ga 的地质演化历史 (Liu *et al.*, 1992; John *et al.*, 2008; Wan Y S *et al.*, 2015)，相比于其他克拉通，它具有更为复杂的构造演化历史，经历了 3.8 Ga 以来的多期重大构造事件 (翟明国, 2010)。除了少量早太古代岩石记录了早期陆壳的形成外，华北克拉通早前寒武纪变质基底还记录了晚太古代 2.9~2.7 Ga 的陆壳生长事件、~2.5 Ga 岩浆-变质事件和古元古代 ~1.95 Ga、~1.85 Ga 的岩浆-变质事件。Zhao 等 (2001, 2005)、Zhao 和 Zhai (2013) 综合华北克拉通基底构造、岩石组合、岩浆-变质作用和同位素年代学等研究成果，将华北克拉通划分出了东、西部两大陆块和胶-辽-吉带、孔兹岩带、中部造山带 3 条古元古代造山带。其中，西部陆块由北部的阴山地块和南部的鄂尔多斯地块在 ~1.95 Ga 沿东西向延伸的孔兹岩带碰撞拼合而成，近同时期东部陆块由龙岗地块和狼林地块碰撞拼合而成，在 ~1.85 Ga 西部陆块与东部陆块沿南北向延伸的中部造山带碰撞拼合，最终形成统一的克拉通 (Zhao *et al.*, 2001; Wilde *et al.*, 2002)。

阿拉善地块位于华北克拉通的最西端，北邻中亚造山带，西侧以阿尔金断裂为界与塔里木克拉通相隔，南接北祁连造山带，东侧以狼山-巴彦乌拉山断裂带为界与华北克拉通西部陆块的主体部分相连 (Zhang J *et al.*, 2013; Wang *et al.*, 2016)，是一个被造山带和断裂带围限的微陆块 (图 1a)。阿拉善地块构造位置复杂，陆表大面积被巴丹吉林沙漠和

腾格里沙漠覆盖，基岩出露较少，其早前寒武纪变质基底主要出露于西部的龙首山、北大山、雅布赖山等地区以及东部的巴彦乌拉山、波罗斯坦庙、狼山等地区，并被后期大规模的显生宙侵入岩脉解 (图 1a)。早期的研究将出露于阿拉善地块东部的早前寒武纪中高级变质岩统称为阿拉善群，时代归为太古宙 (宁夏计委地质局, 1976)^①。后人对这套中高级变质岩进行了重新厘定和划分：霍福臣等 (1987) 将原阿拉善群划分为下部的叠布斯格岩群、中部的阿拉善群和上部的阿拉坦敖包群；陈志勇等 (2004) 将其划分为乌拉山群、色尔腾山岩群的柳树沟岩组和点力素泰岩组；耿元生等 (2007) 将原阿拉善群划分为 3 套变质地层单元 (新太古代叠布斯格岩群、古元古代巴彦乌拉山岩群和古-中元古代阿拉善岩群) 和 2 个变形深成片麻杂岩 (主体为古元古代的波罗斯坦庙片麻岩杂岩和主体为新元古代的毕及格台片麻岩杂岩)。前人通过不同的实验方法获取了一系列年代学数据 (图 1、表 1)，显示阿拉善地块东部分布的巴彦乌拉山岩群、波罗斯坦庙杂岩和叠布斯格岩群的年龄主要以古元古代为主，并有少数尚未确认的太古宙信息。

狼山地区位于阿拉善地块的东北缘，是华北克拉通与阿拉善地块之间的过渡区 (图 1a、1b)。近年来的研究认为，阿拉善地块的东界位于桌子山-贺兰山地区，向北延伸至狼山东侧，狼山地区属于阿拉善地块 (Zhang *et al.*, 2011; 李锦铁等, 2012; 张进等, 2012, 2013; Feng *et al.*, 2013; 公王斌, 2014; Hu *et al.*, 2014; Yuan and Yang, 2015b)。最近 Wang 等 (2016) 根据详细的基底岩石年代学特征的差异，将狼山基底分为北部的主体部分和狼山南段地区两部分，并认为狼山主体部分属于华北克拉通，而狼山南段部分属于阿拉善地块。本文在狼山地区所采集样品均位于其南段的叠布斯格地区，该地区出露阿拉善地块内重要的早前寒武纪变质基底 (耿元生等, 2006, 2007; Dan *et al.*, 2012a)。

2 阿拉善地块东部早前寒武纪变质基底同位素年代学研究进展

2.1 狼山地区

狼山南段地区出露的叠布斯格岩群是阿拉善地

^① 宁夏计委地质局. 1976. 1:20 万区域地质调查报告吉兰泰幅.

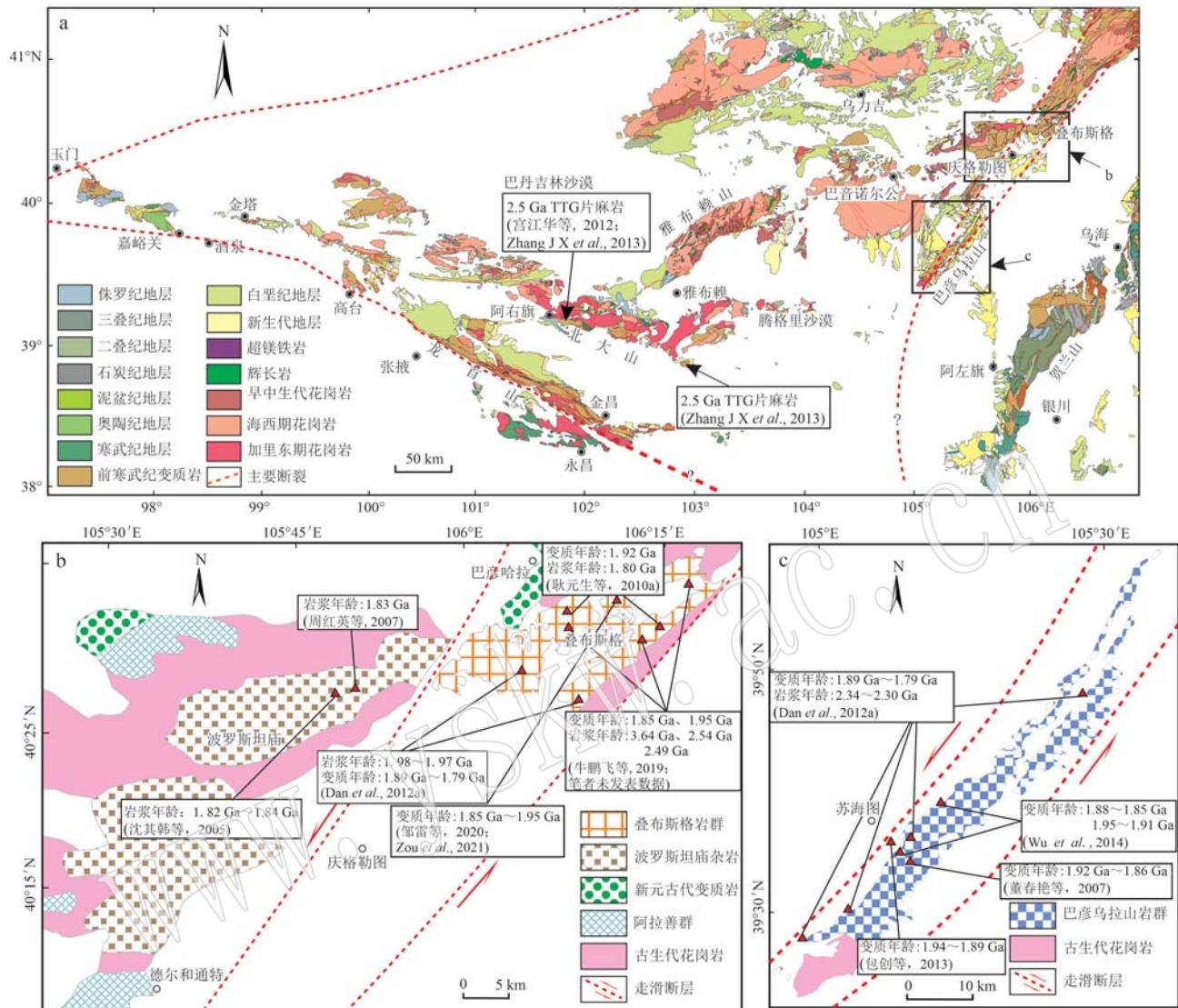


图1 阿拉善地块基底及其东缘早前寒武纪变质基底分布图

Fig. 1 The distribution of early Precambrian basements in the Alxa block and sketch maps of the basements in the eastern Alxa block

a—阿拉善地块地质简图; b—叠布斯格岩群和波罗斯坦庙杂岩(据耿元生等, 2010a); c—巴彦乌拉山岩群(据内蒙古地质矿产局, 1991)

a—simplified geological maps of the Alxa block; b—the Diebusige Complex and Boluositanmiao Complex(after Geng Yuansheng et al., 2010a);

c—the Bayanwulashan Complex(after NMBGMR, 1991)

块内形成时代最老、变质程度最高的变质岩系, 可达到麻粒岩相(霍福臣等, 1987), 以含条带状铁建造和石墨矿为特征, 由黑云斜长片麻岩、斜长角闪岩、大理岩、磁铁石英岩等变质岩和大量深成侵入岩等组成, 变质程度达到麻粒岩相(Zou et al., 2021)。杨振德等(1988)和李俊健(2006)分别获得叠布斯格岩群中斜长角闪岩的Rb-Sr等时线年龄为3 219 Ma, Sm-Nd等时线年龄为 $3\ 018\pm49$ Ma, 也因此认为叠布斯格岩群形成于中太古代; 耿元生等(2006,

2007)通过锆石SHRIMP U-Pb和锆石蒸发法测年获得叠布斯格岩群的变质表壳岩形成于2.70 Ga左右, 并经历了2.69~2.50 Ga和2.0~1.9 Ga两期变质改造, 在该区还发现了最古老的3.5 Ga的残留碎屑锆石, 但由于数据较少无法确定其可靠性。耿元生等(2010a)在叠布斯格岩群片麻岩中识别出2.0~1.9 Ga和1.85~1.80 Ga两期构造热事件, 通过角闪石 ^{39}Ar - ^{40}Ar 测年获得了含辉石斜长角闪岩1 918 Ma的变质年龄; Dan等(2012a)通过锆石SIMS U-Pb

表1 阿拉善地块东部地区早前寒武纪变质基底岩石年龄数据表

Table 1 Dating data for the early Precambrian basements in the eastern Alxa Block

位置	岩性	测试方法	成岩年龄	变质年龄	文献
叠布斯格岩群					
	斜长角闪岩	全岩 Rb-Sr 等时线	3 219 Ma		杨振德等, 1988
	斜长角闪岩	全岩 Sm-Nd 等时线	3 018±49 Ma		李俊健, 2006
N40° 31. 596' E106° 6. 327'	透辉角闪斜长片麻岩	锆石蒸发法和 SHRIMP	2 750~2 690 Ma	2 690~2 450 Ma 2 000~1 900 Ma	耿元生等, 2006, 2007
N40°32. 291' E106°8. 203'	透辉二长片麻岩	SHRIMP	1 802±12 Ma		耿元生等, 2010a
N40°31. 222' E106°12. 531'	长英质片麻岩	SHRIMP		1 926±23 Ma	耿元生等, 2010a
N40°32. 38' E106°8. 23'	黑云斜长片麻岩	SHRIMP	1 970±8 Ma		耿元生等, 2010a
N40°31. 695' E106°6. 291'	含辉石斜长角闪岩	角闪石 ³⁹ Ar- ⁴⁰ Ar 斜长石 ³⁹ Ar- ⁴⁰ Ar		1 918 Ma 1 722 Ma	耿元生等, 2010a
N40°27'3. 3" E106°8'0. 0"	基性片麻岩	SIMS		1 898±17 Ma 1 795±8 Ma	Dan et al., 2012a
N40°28'52" E106°4'53"	钾质花岗岩	SIMS	1 966±7 Ma	1 916~1 819 Ma	Dan et al., 2012a
N40°28'54" E106°4'45"	钾质花岗岩	SIMS	1 981±10 Ma	1 850~1 843 Ma	Dan et al., 2012a
	变辉长岩	SIMS	1 904±19 Ma	1 822±8 Ma	Wan B et al., 2015
	变辉长岩	SIMS	1 906±42 Ma		Wan B et al., 2015
	花岗质片麻岩	SIMS	1 887±29 Ma		Wan B et al., 2015
	变辉绿岩墙	SIMS	1 809±38 Ma		Wan B et al., 2015
N40°30'33. 9" E106°13'12. 04"	黑云斜长片麻岩	LA-ICP-MS		1 942±11 Ma 1 847±12 Ma	牛鹏飞等, 2019
N40°30'47. 41" E106°13'15. 44"	黑云斜长片麻岩	LA-ICP-MS	2 177~2 010 Ma (碎屑锆石)	1 969~1 811 Ma	邹雷等, 2020
	含石榴石磁铁石英岩	LA-ICP-MS		1 854±25 Ma 1 843±18 Ma	Zou et al., 2021
N40°32'34. 2" E106°10'31. 8"	角闪斜长片麻岩	LA-ICP-MS	2 476~2 080 Ma (碎屑锆石)	1 940±19 Ma 1 846±12 Ma	Niu et al., 2022
N40°30'55. 1" E106°12'44. 8"	黑云斜长片麻岩	LA-ICP-MS	2 461~2 121 Ma (碎屑锆石)	1 927±13 Ma 1 830±9 Ma	Niu et al., 2022
N40°34'40. 39" E106°17'5. 34"	斜长角闪岩	LA-ICP-MS	2 636±14 Ma	2 483±54 Ma 1 968±18 Ma	笔者未发表数据
	英云闪长质片麻岩	SHRIMP	2 836±20 Ma	1 951±12 Ma 1 867±12 Ma	Niu et al., 2022
	奥长花岗岩	SHRIMP	2 491±18 Ma	1 834±45 Ma	笔者未发表数据
	英云闪长岩	SHRIMP	2 540±38 Ma	1 784±24 Ma	笔者未发表数据
	花岗片麻岩	SHRIMP	2 763±42 Ma	1 852 ± 46 Ma 1 753~1 721 Ma	Niu et al., 2022
波罗斯坦庙杂岩					
庆格勒图北	黑云斜长片麻岩	TIMS	1 826±13 Ma		周红英等, 2007
	黑云母花岗片麻岩/ 黑云斜长片麻岩	SHRIMP	1 818±19 Ma 1 839±18 Ma		沈其韩等, 2005
巴彦乌拉山岩群					
苏海图南 10 km	花岗闪长质片麻岩	TIMS	2 082±22 Ma		李俊健等, 2004
N39°36. 85' E105°9. 62'	黑云斜长片麻岩	锆石蒸发法	2 271~2 264 Ma		沈其韩等, 2005; 耿元生等, 2007
N39°36'29" E105°10'4"	片麻状花岗岩	SHRIMP	2 323±20 Ma	1 923±28 Ma 1 856±12 Ma	董春艳等, 2007

续表 1

Continued Table 1

位置	岩性	测试方法	成岩年龄	变质年龄	文献
N39°36'44" E105°9'21"	斜长角闪岩	SIMS	2 329±9 Ma	1 926~1 812 Ma	Dan <i>et al.</i> , 2012a
N39°31'47" E105°5'12"	基性片麻岩	SIMS	2 344±12 Ma	1 792±17 Ma	Dan <i>et al.</i> , 2012a
N39°36'11" E105°12'44"	长英质副片麻岩	SIMS		1 891±9 Ma 1 796 ± 22 Ma	Dan <i>et al.</i> , 2012a
N39°31'47" E105°5'12"	正片麻岩	SIMS	2 303±5 Ma		Dan <i>et al.</i> , 2012a
E105°30'46" N39°47'17"	正片麻岩	SIMS	2 334±6 Ma	1 892±8 Ma 1 792±17 Ma	Dan <i>et al.</i> , 2012a
E105°29'42" N39°28'9"	正片麻岩	SIMS	2 324±10 Ma	1 767±23 Ma	Dan <i>et al.</i> , 2012a
E104°58'58" N39°36'24"	花岗岩脉	SIMS	1 895±28 Ma		Dan <i>et al.</i> , 2012a
E105°11'55" N39°40'57. 2"	斜长角闪岩	LA-ICP-MS	2 300 Ma	2 239~1 988 Ma 1 940~1 892 Ma	包创等, 2013
E105°15'13" N39°36'47. 1"	角闪斜长片麻岩	LA-ICP-MS	2 290±11 Ma	1 878±13 Ma	Wu <i>et al.</i> , 2014
E105°8'35. 5" N39°36'47. 1"	角闪斜长片麻岩	LA-ICP-MS	2 244±10 Ma	1 935±8. 9 Ma	Wu <i>et al.</i> , 2014
E105°8'35. 5" N39°36'39. 6"	花岗质片麻岩	LA-ICP-MS	2 301±20 Ma	1 858±25 Ma	Wu <i>et al.</i> , 2014
E105°9'6. 7" N39°36'38. 2"	黑云斜长片麻岩	LA-ICP-MS		1 904. 8±6 Ma	Wu <i>et al.</i> , 2014
E105°9'51. 3" N39°36'19. 3"	黑云斜长片麻岩	LA-ICP-MS	2 266±20 Ma	1 876±10 Ma	Wu <i>et al.</i> , 2014
E105°12'8. 5"	二长片麻岩	LA-ICP-MS	2 232±18 Ma	1 911. 4±4. 1 Ma	Wu <i>et al.</i> , 2014
				1 851±17 Ma	Wu <i>et al.</i> , 2014
				1 955±27 Ma	Wu <i>et al.</i> , 2014

测年将叠布斯格岩群中的变质表壳岩沉积时代限定在了 2.45~2.0 Ga 之间, 并获得了 1.98~1.97 Ga 的深成侵入岩岩浆结晶年龄和~1.89 Ga、~1.79 Ga 的变质年龄, 认为狼山地区不存在太古宙表壳岩, 但锆石 Hf 同位素结果显示并不排除地壳深部存在太古宙基底的可能性。Wan B 等(2015)在狼山地区发现了~1.9 Ga 的变辉长岩和花岗质片麻岩, 并获得了 1.82 Ga 左右的变质年龄。最近, 笔者对叠布斯格岩群进行了详细的锆石年代学和 Hf 同位素研究(牛鹏飞等, 2019; Niu *et al.*, 2022)(表 1), 获得副片麻岩中碎屑锆石的年龄主要集中在 2.7~2.1 Ga 之间, 并获得了 1.99~1.89 Ga 和 1.88~1.76 Ga 两期变质年龄(图 2), 结合碎屑锆石和变质锆石年龄, 将叠布斯格岩群部分变质表壳岩的沉积时代限定在了 2.1~2.0 Ga 之间; 副片麻岩中碎屑锆石的 Hf 同位素模式年龄为 3.4~2.7 Ga, 表明这些碎屑锆石主要来源于中新太古代古老陆壳物质。另外, 笔

者对叠布斯格岩群中广泛发育的钾质花岗岩也进行了锆石测年, 获得其岩浆结晶年龄为 1 860±8 Ma(笔者未发表数据), 与 Dan 等(2012a)获得的 1.98~1.97 Ga 的深成侵入岩结晶年龄共同成为该地区古元古代晚期的岩浆事件记录。

此外, 笔者最近在狼山地区识别出多件新太古代变质岩石, 其中斜长角闪岩中的岩浆锆石年龄为 2 697~2 566 Ma(加权平均年龄为 2 636±14 Ma), 两期变质年龄分别为 2 517~2 454 Ma(加权平均年龄为 2 483±54 Ma)和 1 988~1 952 Ma(加权平均年龄为 1 968±18 Ma); 奥长花岗质片麻岩的成岩年龄为 2 491±18 Ma, 变质锆石的年龄为 1 834±45 Ma, 代表了一期古元古代变质事件; 英云闪长质片麻岩的原岩岩浆结晶年龄为 2 540±38 Ma, 变质锆石年龄为 1 866~1 748 Ma, 一组捕获锆石的年龄为 2 667~2 588 Ma(笔者未发表数据), 不仅表明阿拉善地块东部出露新太古代岩石, 更说明该地区地壳深部

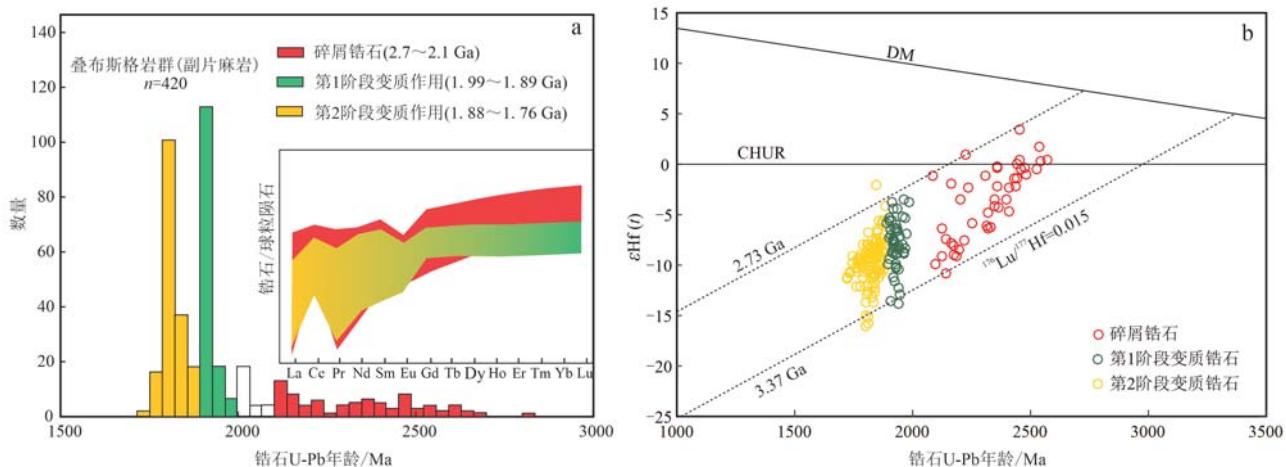


图 2 叠布斯格岩群副片麻岩锆石年龄柱状图(a)和Hf同位素组成(b)

Fig. 2 Zircon age histogram (a) and Hf isotopic composition (b) of the paragneisses from the Diebusige Complex

也是存在太古宙基底的。这套新太古代岩石从岩性到形成时代都与叠布斯格岩群有明显区别，因此，笔者将这套新太古代变质侵入体从叠布斯格岩群岩石组成中分离出来，独立于叠布斯格岩群之外作为狼山地区的新太古代变质基底。

锆石 Hf 同位素结果显示狼山地区 2.64 Ga 斜长角闪岩岩浆锆石的 εHf 值为 $-5.83 \sim -1.01$ ，二阶段模式年龄集中在 3.38 ~ 3.21 Ga (笔者未发表数据)，其原岩可能是古太古代古老陆壳物质在新太古代晚期的再循环产物 (图 3)，但不排除来自富集地幔的可能；新太古代末期 2.54 Ga 英云闪长质片麻岩岩浆锆石和捕获锆石的 εHf 值分别为 $-1.1 \sim 0.71$ 和 $0.8 \sim 3.46$ ，两类锆石的二阶段模式年龄均集中在 3.10 ~ 2.88 Ga 之间，而 2.49 Ga 奥长花岗质片麻岩

岩浆锆石的 εHf 值在 $-3.83 \sim -2.12$ 之间，二阶段模式年龄为 3.23 ~ 3.12 Ga (图 3)，均指示狼山地区 ~2.5 Ga TTG 片麻岩主要来自于中太古代陆壳物质的重熔或再造。宫江华等 (2012) 和 Zhang J X 等 (2013) 在西阿拉善地块的北大山地区也发现了 ~2.5 Ga 的 TTG 片麻岩，锆石 Hf 同位素特征显示其来自于 3.0 ~ 2.7 Ga 陆壳物质在新太古代末期的重熔或再造 (图 3)，相比狼山地区 ~2.5 Ga TTG 片麻岩具有稍年轻的源区时代。

2.2 波罗斯坦庙地区

位于阿拉善地块东北缘庆格勒图北部波罗斯坦庙地区的波罗斯坦庙杂岩 (图 1a、1b) 主体为一套古元古代晚期 (1.95 ~ 1.85 Ga) 的 TTG 质片麻岩 (耿元生等, 2010a)，主要出露黑云斜长片麻岩、黑云二长片麻岩、角闪斜长片麻岩等变质深成侵入岩，变质程度达到角闪岩相 (霍福臣等, 1987)。目前有关波罗斯坦庙杂岩的研究只有少量的年代学工作，沈其韩等 (2005) 利用锆石 SHRIMP U-Pb 测年获得花岗片麻岩的原岩侵入年龄为 1818 ± 19 Ma 和 1839 ± 18 Ma；周红英等 (2007) 通过锆石 TIMS U-Pb 测年获得黑云斜长片麻岩原岩的岩浆结晶年龄为 1826 ± 13 Ma。

2.3 巴彦乌拉山地区

巴彦乌拉山岩群呈条带状独立分布于巴彦乌拉山地区 (图 1a、1c)，与其他岩群未直接接触 (耿元生等, 2007)，主要由斜长角闪岩、花岗质片麻岩、黑云斜长片麻岩等组成，夹少量酸性变质火山岩、火山沉积岩、大理岩等，总体变质程度为高角闪岩相 (沈其韩等, 2004；Wu et al., 2014)。李俊健等 (2004)

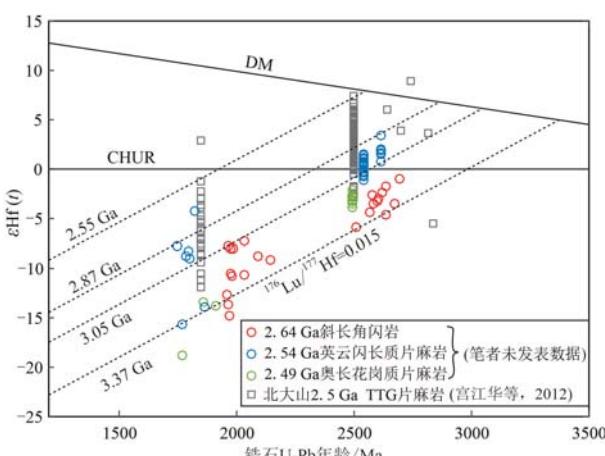


图 3 阿拉善地块东部太古宙岩石锆石 Hf 同位素组成

Fig. 3 Concordia diagram of and zircon Hf isotopic composition of Archean rocks from the eastern Alxa Block

获得巴彦乌拉山岩群中的花岗闪长质片麻岩的单颗粒锆石 U-Pb 年龄为 $2\ 082 \pm 22$ Ma, 表明该地区主要片麻岩的形成时代为古元古代; 沈其韩等(2005)和耿元生等(2007)利用锆石蒸发法, 获得斜长角闪岩的原岩年龄为 $2\ 271 \pm 18 \sim 2\ 264 \pm 3$ Ma; 李俊健(2006)对巴彦乌拉山岩群的斜长角闪岩进行 Sm-Nd 等时线测年, 获得了 $2\ 005 \pm 21 \sim 1\ 920 \pm 38$ Ma 的成岩年龄; 董春艳等(2007)通过锆石 SHRIMP U-Pb 测年获得片麻状花岗岩形成于 $2\ 323 \pm 20$ Ma, 并获得了 $1\ 923 \pm 28$ Ma 和 $1\ 856 \pm 12$ Ma 的变质年龄; Dan 等(2012a)通过锆石 SIMS U-Pb 测年, 获得巴彦乌拉山岩群多件正片麻岩的成岩年龄为 $2.34 \sim 2.30$ Ga, 并在长英质副片麻岩中获得了 $1\ 891 \pm 9$ Ma 和 $1\ 796 \pm 22$ Ma 的变质年龄, 锆石 Hf 同位素结果显示 2.30 Ga 左右的岩浆锆石的模式年龄为 $2.92 \sim 2.81$ Ga; 包创等(2013)对斜长角闪岩进行锆石 LA-ICP-MS 定年, 获得其成岩年龄为 2.30 Ga, 锆石 Hf 同位素结果显示其模式年龄为 $2.80 \sim 2.45$ Ga, 另外在斜

长角闪岩中也识别出 $2\ 239 \sim 1\ 988$ Ma 和 $1\ 940 \sim 1\ 892$ Ma 的变质年龄; Wu 等(2014)通过锆石 LA-ICP-MS 测年获得多件正片麻岩的岩浆结晶年龄集中在 $2.30 \sim 2.24$ Ga, 在正片麻岩和副片麻岩中获得的变质年龄集中在 $1.94 \sim 1.91$ Ga 和 $1.88 \sim 1.86$ Ga 之间。

3 阿拉善地块东部早前寒武纪构造热事件

如上所述, 前人在阿拉善地块东部地区的研究工作使得该地区早前寒武纪变质基底的年代学格架逐渐完善。统计分析前人报道的同位素年代学数据可知(表 1、图 4), 阿拉善地块东部地区的叠布斯格岩群、波罗斯坦庙杂岩和巴彦乌拉山岩群普遍受到古元古代晚期构造热事件的影响, 而新太古代晚期和古元古代早期的岩浆热事件分别发生于狼山地区和巴彦乌拉山地区。

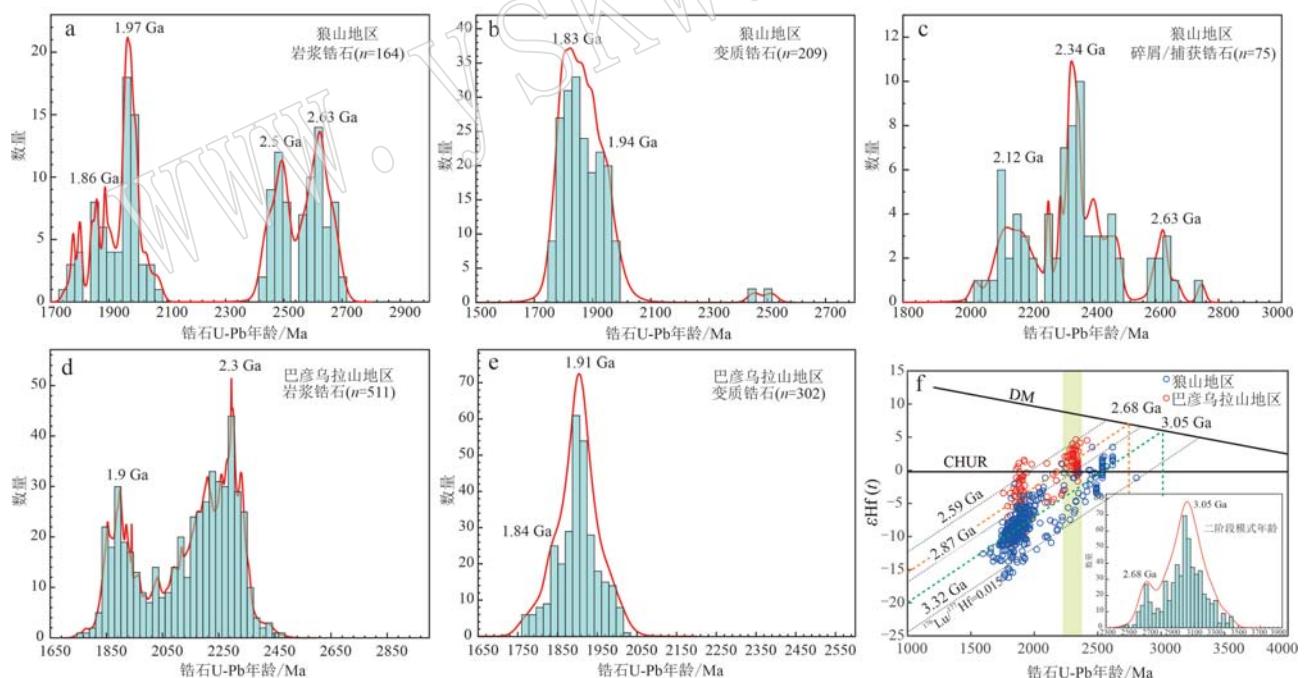


图 4 阿拉善地块东部地区前寒武纪变质基底变质锆石、正变质岩的岩浆锆石、捕获锆石和副变质岩的碎屑锆石年龄直方图和锆石 Hf 同位素组成

Fig. 4 Age histograms of the magmatic, detrital and metamorphic zircons from the early Precambrian basements of the Eastern Alxa Block and their Hf isotopic composition

叠布斯格岩群数据来自 Dan 等(2012a)、牛鹏飞等(2019)、Niu 等(2022)和未发表数据; 巴彦乌拉山岩群数据引自董春艳等(2007)、Dan 等(2012a)、包创等(2013)、Wu 等(2014)

data of the Diebusige Complex are from Dan et al., 2012a; Niu Pengfei et al., 2019; Niu et al., 2022; unpublished data;

data of the Bayanwulashan Complex are from Dong Chunyan et al., 2007; Dan et al., 2012a; Bao Chuang et al., 2013; Wu et al., 2014

3.1 古元古代早期之前(>2.45 Ga)陆壳生长和岩浆-变质事件

阿拉善地块东部新太古代岩浆-变质事件发生在2 697~2 454 Ma之间,包括狼山地区2.64 Ga的斜长角闪岩和2.54 Ga、2.49 Ga的TTG片麻岩的形成以及2.48 Ga的变质作用(笔者未发表数据,图4a、4b),与西阿拉善北大山地区出露的2.55~2.51 Ga的TTG片麻岩及其记录的2.52~2.47 Ga变质作用属于同一期岩浆-变质事件(宫江华等,2012; Zhang J X et al., 2013)。另外,阿拉善地块东部地区叠布斯格岩群和巴彦乌拉山岩群的锆石Hf同位素数据结果显示,早前寒武纪变质基底中锆石Hf模式年龄集中分布在3.4~2.6 Ga之间,并且大致存在2.68 Ga和3.05 Ga两个明显的峰值年龄(图4f),表明古太古代晚期至新太古代晚期阿拉善地块东部存在多期陆壳生长和再循环。

3.2 古元古代早期(~2.3 Ga)岩浆事件

这一期岩浆事件在西阿拉善龙首山地区和华北克拉通鄂尔多斯地块内均有发生(修群业等,2004;刘勇,2008;宫江华等,2011; Gong et al., 2016; Hu et al., 2013; Wan et al., 2013a; Wang et al., 2014; Zhang et al., 2015),也存在于阿拉善地块东部的巴彦乌拉山地区,表现为2.3 Ga左右的正片麻岩和斜长角闪岩出露(图4d; 沈其韩等,2005; 耿元生等,2007; 董春艳等,2007; Dan et al., 2012a; 包创等,2013; Wu et al., 2014),这可能代表了孔兹岩带附近的古元古代构造活动带向西的延伸(董春艳等,2007; 耿元生等,2009)。通过锆石年代学统计发现,叠布斯格岩群副片麻岩中的碎屑锆石存在2.3 Ga左右的年龄峰值(图4c),邹雷等(2020)根据这一特点认为巴彦乌拉山地区的古元古代变质基底可以作为叠布斯格岩群副片麻岩的潜在物源区,但锆石Hf同位素结果显示巴彦乌拉山岩群中~2.33 Ga的斜长角闪岩岩浆锆石的 ε_{Hf} 值为-2.2~4.5,模式年龄在2.92~2.45 Ga之间,而叠布斯格岩群副片麻岩中~2.3 Ga碎屑锆石的 ε_{Hf} 值和模式年龄分别为-8.4~-0.2和3.29~2.90 Ga,与巴彦乌拉山岩群具有明显差别(图4f),因此可以排除巴彦乌拉山岩群中的2.30 Ga变质基底岩石作为叠布斯格岩群副片麻岩物源的可能。

3.3 古元古代中晚期(2.0~1.8 Ga)岩浆-变质事件

古元古代晚期事件主要发生在2.0~1.8 Ga之

间,包括叠布斯格岩群深成侵入岩中岩浆锆石记录的1.98~1.97 Ga岩浆事件年龄(Dan et al., 2012a)(图4a)、斜长角闪岩变质锆石记录的1 918 Ma的变质事件年龄(耿元生等,2010a),巴彦乌拉山岩群中片麻状花岗岩锆石幔部和边部记录的1 923±28 Ma和1 856±12 Ma的变质事件年龄(董春艳等,2007)(图4d),以及波罗斯坦庙杂岩中正片麻岩记录的1.95~1.80 Ga岩浆事件年龄(沈其韩等,2005; 周红英等,2007; 耿元生等,2010a)等。~1.9 Ga和~1.8 Ga的古元古代变质事件在华北克拉通和阿拉善地块内广泛发育。Dan等(2012a)在阿拉善地块东部的叠布斯格岩群和巴彦乌拉山岩群正片麻岩中也识别出相似的两组古元古代变质年龄(~1.89 Ga和~1.79 Ga),并认为早期的变质事件可能与大陆碰撞有关,而晚期变质事件可能是区域性伸展造成的,分别与古元古代哥伦比亚超大陆的聚合和裂解是同时期的(Zhao et al., 2002, 2004; Hou et al., 2008)。Zhao等(2005)认为早期变质事件主要发生在西部的孔兹岩带中,而晚期变质事件与华北克拉通中部带的形成有关。但是近年来的年代学研究结果显示,孔兹岩带不仅记录了~1.9 Ga的构造热事件,同时也记录了~1.8 Ga的构造热事件(郭敬辉等,1999, 2002; Guo et al., 2005; Santosh et al., 2006; Yin et al., 2011; Ma et al., 2012; Wan et al., 2013a)。另外值得注意的是,古元古代晚期的变质事件年龄呈现出连续分布的特征(图2a、图4b、4e),虽然显示出两个年龄峰值,但是其究竟是代表着两期变质事件还是一期变质事件的两个阶段,以及反映了怎样的构造背景,还有待进一步研究。

4 区域对比及地质意义

如前文所述,阿拉善地块东部经历了3.4~2.6 Ga的陆壳生长,存在~2.5 Ga TTG岩浆-变质事件和古元古代中晚期2.3~2.0 Ga岩浆事件、~1.95 Ga和~1.85 Ga岩浆-变质事件(图5a-1)。阿拉善地块东部的狼山地区与西阿拉善北大山地区出露的~2.5 Ga的TTG片麻岩具有近一致的陆壳生长历史(图5a-2)。张建新等(2018)在波罗斯坦庙杂岩中也识别出一套~2.50 Ga的TTG片麻岩,狼山地区的太古宙变质基底可能向西经过波罗斯坦庙地区与西阿拉善的北大山地区相连。发育在阿拉善地块东部巴

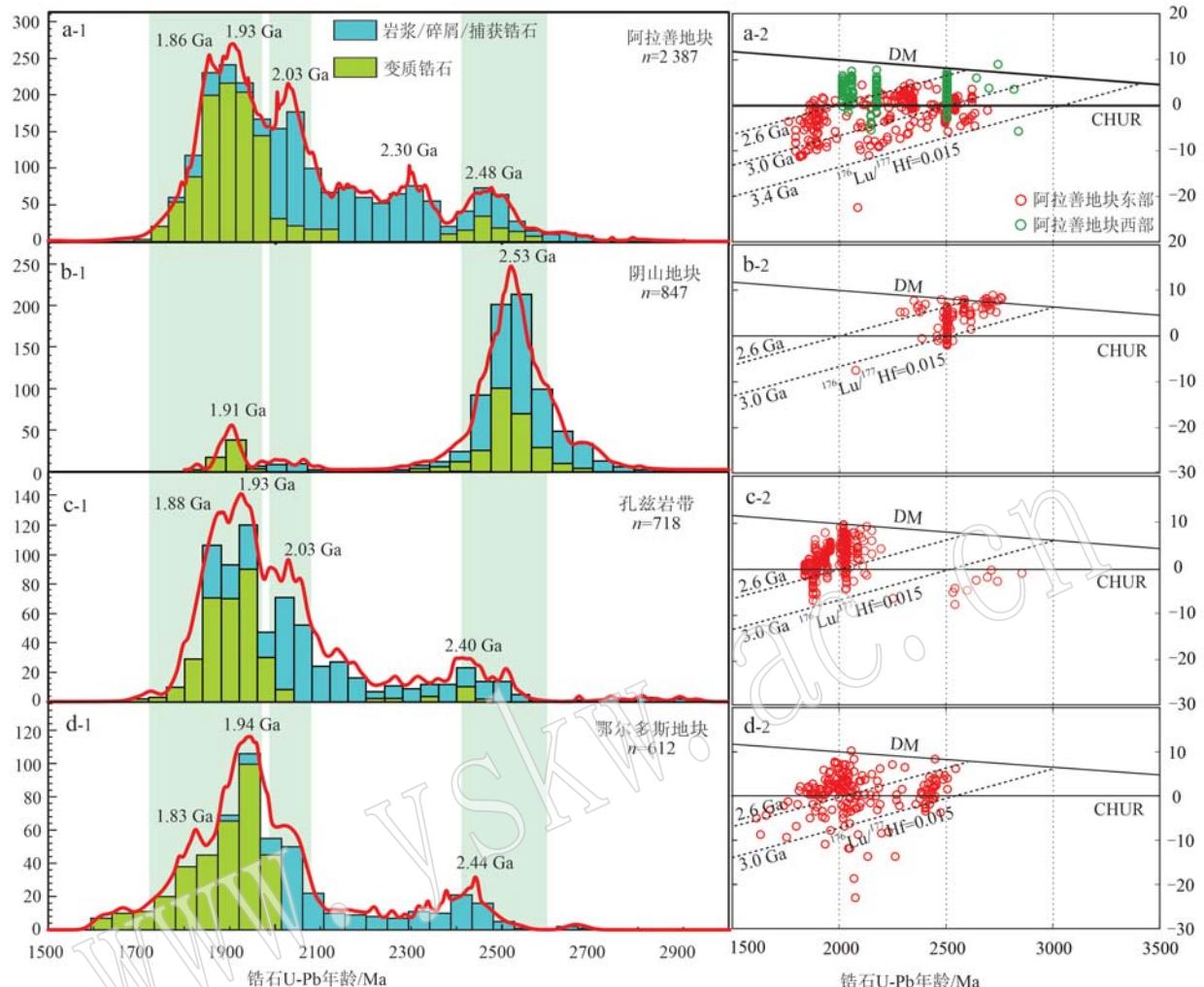


图 5 阿拉善地块与华北克拉通西部陆块早前寒武纪变质基底锆石年龄对比直方图及岩浆锆石的 Hf 同位素组成

Fig. 5 Age comparison histograms and corresponding magmatic zircon Hf isotopic composition of the early Precambrian basements in the Alxa Block, Yinshan Block, Khondalite Belt and Ordos Block

a-1 和 a-2 数据引自陆松年等(2002)、董国安等(2007)、董春艳等(2007)、宫江华等(2012)、Dan 等(2012a)、包创等(2013)、Wu 等(2014)、Gong 等(2016)、牛鹏飞等(2019)、Niu 等(2022)；b-1 和 b-2 数据引自张维杰等(2000)、简平等(2005)、范宏瑞等(2010)、董晓杰等(2012)、Jian 等(2012)、Ma 等(2013)、马铭株等(2013)、Wang 等(2015)、刘平华等(2016)；c-1 和 c-2 数据引自张维杰等(2000)、Xia 等(2006)、董春艳等(2007)、周喜文等(2009)、Yin 等(2009, 2011)、Dan 等(2012a, 2012b)、Dong 等(2014)、蔡佳等(2015)；d-1 和 d-2 数据引自 Wan 等(2013b)、Zhang 等(2015)、Gou 等(2016)、He 等(2016)

data of a-1 and a-2 are from Lu Songnian et al., 2002; Dong Guo'an et al., 2007; Dong Chunyan et al., 2007; Gong Jianghua et al., 2012; Dan et al., 2012a; Bao Chuang et al., 2013; Wu et al., 2014; Gong et al., 2016; Niu Pengfei et al., 2019; Niu et al., 2022; data of b-1 and b-2 are from Zhang Weijie et al., 2000; Jian Ping et al., 2005; Fan Hongrui et al., 2010; Dong Xiaojie et al., 2012; Jian et al., 2012; Ma et al., 2013; Ma Mingzhu et al., 2013; Wang et al., 2015; Liu Pinghua et al., 2016; data of c-1 and c-2 are from Zhang Weijie et al., 2000; Xia et al., 2006; Dong Chunyan et al., 2007; Zhou Xiwen et al., 2009; Yin et al., 2009, 2011; Dan et al., 2012a, 2012b; Dong et al., 2014; Cai Jia et al., 2015; data of d-1 and d-2 are from Wang et al., 2013b; Zhang et al., 2015; Gou et al., 2016; He et al., 2016

彦乌拉山地区的 2.3~2.0 Ga 岩浆事件在西阿拉善的龙首山地区也有记录(Gong et al., 2016)，二者古元古代变质表壳岩的形成时代也近一致(耿元生等, 2007; Gong et al., 2016; 张建新等, 2018)，指示巴彦乌拉山地区的古元古代变质基底向西可与龙首山

地区的变质基底相连。阿拉善地块的变质基底普遍记录了古元古代晚期的岩浆-变质事件，如阿拉善地块东部叠布斯格岩群中 1.98~1.97 Ga 的深成侵入岩及其记录的 1.89 Ga 和 1.79 Ga 两期变质事件(Dan et al., 2012a)、笔者在叠布斯格岩群副片麻岩

中获得的~1.95 Ga 及~1.85 Ga 的变质事件(牛鹏飞等, 2019; 笔者未发表数据)、波罗斯坦庙杂岩中 1.95~1.80 Ga 的 TTG 质片麻岩(沈其韩等, 2005; 耿元生等, 2010a)以及西阿拉善北大山地区 2.5 Ga TTG 片麻岩所记录的 1.85 Ga 的变质事件(Zhang J X et al., 2013)等。古元古代晚期的岩浆-变质事件在阿拉善地块内的广泛存在表明, 阿拉善地块可能整体都卷入了古元古代晚期的构造热事件中。而华北克拉通主体的前寒武纪变质基底主要形成于新太古代—古元古代, 以 2.8~2.7 Ga 的地壳生长时代(Jiang et al., 2010)和~2.5 Ga 的基底成岩时代(Zhao et al., 2005; Zhai and Santosh, 2011; Wang and Liu, 2012; Wan et al., 2013a)为特征, 同时也广泛记录了~1.95 Ga 和~1.85 Ga 两期变质事件(Wu et al., 2005; Wan et al., 2006, 2013a; Yin et al., 2009, 2011, 2014; Zhao et al., 2010, 2012; Dong et al., 2013)。由此可见, 阿拉善地块与华北克拉通主体共同经历了太古宙陆壳生长和新太古代—古元古代岩浆-变质事件, 但阿拉善地块与华北克拉通之间究竟具有怎样的关系和演化历史? 阿拉善地块在何时、以何种方式与华北克拉通拼合?

4.1 与阴山地块对比

在地理分界上, 阴山地块与阿拉善地块以狼山断裂带相隔。Zhao 等(2005)认为阿拉善地块是阴山地块的西延部分, 一些学者也曾试图将阴山地块内色尔腾山-固阳-集宁-桑干等地区的早前寒武纪基底地层统一起来(陈志勇等, 2004, 2007), 但目前没有资料显示阿拉善地块与华北克拉通之间存在不同陆块拼合的地质证据。近年来的研究表明, 阿拉善地块的早前寒武纪变质基底与阴山地块是有所区别的。在物质组成方面, 阴山地块不同地区不同程度地出露太古宙基底岩石, 以固阳、武川等地区出露最为完整和具代表性, 主要由新太古代花岗绿岩带以及 TTG 片麻岩和麻粒岩-紫苏花岗岩类等组成(Zhao et al., 2012; Jian et al., 2012; Zhao and Zhai, 2013)。而阿拉善地块东部的叠布斯格岩群、波罗斯坦庙杂岩和巴彦乌拉山岩群以及西阿拉善地块内的其他变质基底均形成于古元古代, 未见有大面积的太古宙变质表壳岩和花岗绿岩带出露(李文渊等, 2004; Dan et al., 2012a)。在岩浆-变质作用记录方面, 阿拉善地块内仅有北大山地区的 2.5 Ga TTG 片麻岩(宫江华等, 2012; Zhang J X et al., 2013)和笔者最近在狼山地区识别出的新太古代岩

石可以对比阴山地块内广泛分布的新太古代 TTG 岩石。阴山地块自~2.7 Ga 就已有陆壳的生长(董晓杰等, 2012; 马铭株等, 2013), 并在 2.58~2.53 Ga 出现大量岩浆活动, 在 2.5~2.45 Ga 遭受了变质作用的改造(简平等, 2005; Jian et al., 2012; Ma et al., 2014 及相关参考文献)(图 5b-1、b-2)。虽然阿拉善地块与阴山地块均记录有新太古代晚期岩浆-变质事件, 但广泛发育于阿拉善地块早前寒武纪变质基底中的 2.3~2.0 Ga 的岩浆事件和 2.0~1.8 Ga 的岩浆-变质事件在阴山地块内报道较少(Jian et al., 2012; 马铭株等, 2013)(图 5b-1)。鉴于在物质组成和岩浆-变质作用记录方面的明显差异, 阿拉善地块与阴山地块应属于不同的构造单元, 至少在古元古代之前具有不同的地质演化历史。

4.2 与孔兹岩带对比

华北克拉通的孔兹岩带自西向东主要分布在集宁-乌拉山-大青山-千里山-贺兰山地区, 由阴山地块与鄂尔多斯地块在~1.95 Ga 拼合形成, 代表了鄂尔多斯地块周缘的被动大陆边缘沉积(Zhao et al., 2005)。孔兹岩带基底岩石主要由中压泥质麻粒岩、高压与超高温变质的麻粒岩、少量 TTG 片麻岩、同构造紫苏花岗岩和 S 型花岗岩等组成(Zhao et al., 2005; Santosh et al., 2007; 翟明国, 2009; Guo et al., 2012; Ma et al., 2012; Zhao and Zhai, 2013; Cai et al., 2014; Yin et al., 2014)。前人对孔兹岩带基底岩石开展了大量的年代学工作, 近年来在大青山地区古元古代早期片麻岩中识别出 2.5~2.44 Ga 的花岗岩和辉长岩包体, 并经历了 2.5 Ga、2.45 Ga 变质作用的改造(Wan et al., 2009, 2013a; 刘平华等, 2013; Liu et al., 2014, 2017), 表明孔兹岩带基底经历了与阿拉善地块和阴山地块近一致的新太古代晚期-古元古代早期岩浆-变质事件。孔兹岩带变沉积岩中的碎屑锆石年龄主要集中在 2.25~2.0 Ga 之间, 还包括少量 2.9~2.25 Ga 之间的锆石年龄, 结合~1.95 Ga 和~1.86 Ga 的变质年龄, 将孔兹岩系的沉积时代限定在了 2.0~1.95 Ga 之间(Zhao et al., 2005; Wan et al., 2006, 2009; Yin et al., 2011; Dan et al., 2012b; 蔡佳等, 2015), 与阿拉善地块东部地区巴彦乌拉山岩群形成时代(2.27~1.95 Ga)近一致, 略年轻于叠布斯格岩群(2.1~2.0 Ga)。通过对阿拉善地块与孔兹岩带早前寒武纪变质基底的锆石 Hf 同位素数据发现, 孔兹岩带 2.25~2.0 Ga 的碎屑锆石与阿拉善地块龙首山地区

2.3~2.0 Ga 侵入岩的岩浆锆石在年龄和 Hf 同位素特征方面近一致(图 5c-1、c-2), 表明古元古代晚期西阿拉善龙首山地区可能作为物源区之一, 为孔兹岩系岩石提供成岩物质。此外值得一提的是, 阿拉善地块东部新太古代岩石中的部分岩浆锆石与孔兹岩带贺兰山地块基底岩石中的少量碎屑锆石的 Hf 同位素组成一致, 同样表明在古元古代中晚期阿拉善地块东部地区可能也是孔兹岩带的物源区之一。

孔兹岩带广泛记录的 1.95~1.9 Ga 的变质事件被认为是阴山地块与鄂尔多斯地块碰撞拼合形成华北克拉通西部陆块的响应, 而~1.85 Ga 的变质年龄对应于西部陆块与东部陆块沿古元古代中部造山带拼合形成统一的华北克拉通的时代(Zhao et al., 2005)。近年来的年代学数据显示, 孔兹岩带东部集宁、怀安地区的中、高压泥质变质岩、变辉长岩和紫苏花岗岩等记录了 1.96~1.92 Ga 和 1.90~1.81 Ga 的变质事件, 并发育 1.95 Ga 的淡色花岗岩和 1.92~1.89 Ga 的 S 型花岗岩(Wan et al., 2006; Zhao et al., 2010; Jiao et al., 2013a; Santosh et al., 2013; Peng et al., 2010, 2014); 中部大青山地区的古元古代表壳岩中记录了~1.96 Ga 和 1.88~1.83 Ga 的变质事件, 也存在 1.95 Ga 的正长花岗岩和紫苏闪长岩的岩浆活动(Ma et al., 2012; Dong et al., 2013; Jiao et al., 2013b; Wang et al., 2017); 西部贺兰山-千里山地区的泥质麻粒岩中也存在 1.96~1.92 Ga 和 1.87 Ga 两组变质年龄和 1.9 Ga、1.88~1.86 Ga 的 S 型花岗岩侵入事件(Yin et al., 2009, 2011; Zhang et al., 2017)。这些数据表明, ~1.85 Ga 的岩浆-变质事件也存在于孔兹岩带中, 因此古元古代晚期的这两期变质事件应具有更广泛的构造意义, 很可能是对 Columbia 超大陆古元古代地质演化的响应(Dan et al., 2012a; Ge et al., 2013)。但无论如何, 阿拉善地块内广泛记录的~1.95 Ga 和~1.85 Ga 的岩浆-变质事件与孔兹岩带极为相似(图 5a-1、c-1), 表明二者至少在古元古代中晚期是具有亲缘性的。

4.3 与鄂尔多斯地块对比

孔兹岩带以南的鄂尔多斯地块被认为是一个具有统一基底的太古宙稳定陆块(张抗, 1982; 贾进斗等, 1997; 邓军等, 2005), 是华北克拉通西部陆块的重要组成部分(Zhao et al., 2005; Zhao and Zhai, 2013), 航磁异常显示其具有整体刚性块体特征(Gao et al., 2015)。地块基底被巨厚的中新元古代

碎屑沉积岩、早古生代浅海碎屑岩和碳酸盐沉积以及晚古生代-新生代陆相碎屑岩覆盖。前人对地块内少量基底岩芯的研究显示, 地块内基底岩石主要由副片麻岩、变粒岩、片岩、大理岩和花岗质片麻岩等组成(Hu et al., 2013; Wan et al., 2013a; Zhang et al., 2015), 变质程度达到中压麻粒岩相(Wang et al., 2014; Gou et al., 2016; He et al., 2016)。

Zhang 等(2015)在鄂尔多斯地块北部、中东部和中部的基底岩芯中识别出了少量~2.5 Ga 的岩浆活动信息, 与阿拉善地块不同的是在该地块内未发现同时期的变质作用(图 5d-1), ~2.5 Ga 岩浆锆石的 Hf 同位素结果显示鄂尔多斯地块同样存在 3.0~2.6 Ga 的陆壳生长(图 5d-2), 与阴山地块和阿拉善地块的陆壳生长历史非常相似。而且, 张成立等(2018)在沉积盖层中也统计出~2.7 Ga 的碎屑锆石, 表明鄂尔多斯地块很可能在新太古代之前就已有一定的陆壳物质形成。鄂尔多斯地块北部 2.2~2.0 Ga 的岩浆活动十分强烈(胡建民等, 2012; Zhang et al., 2015; 张成立等, 2018), 在时代和锆石 Hf 同位素特征方面与阿拉善地块巴彦乌拉山地区和龙首山地区非常相似(图 5a-1、a-2、d-1、d-2), 这些岩浆活动可能共同形成了孔兹岩带附近规模较大的古元古代活动构造带, 并为孔兹岩的形成提供物质来源(董春艳等, 2007; 张建新等, 2018; 张成立等, 2018)。此外, 鄂尔多斯地块北部的基底岩石与阿拉善地块经历了同时期的~1.95 Ga 和~1.85 Ga 岩浆-变质事件(Wang et al., 2014; Zhang et al., 2015; 吴素娟等, 2015; Gou et al., 2016; He et al., 2016), 也表明鄂尔多斯地块古元古代的地壳演化与阿拉善地块具有较高的相似度, 但二者是否具有亲缘性还需要更多的证据。

综上所述, 阿拉善地块与华北克拉通西部陆块共同经历了中新太古代(3.0~2.6 Ga)陆壳生长, 而后又被新太古代晚期-古元古代早期(~2.5 Ga)岩浆-变质事件改造, 但鄂尔多斯地块内尚未发现~2.5 Ga 的变质作用记录。2.3~2.0 Ga 的岩浆活动在阿拉善地块和鄂尔多斯地块内十分强烈, 而阴山地块和孔兹岩带缺失该时期的岩浆活动, 表明阿拉善地块与鄂尔多斯地块北部基底很可能在古元古代早中期处于同一活动构造带内(董春艳等, 2007)。阿拉善地块内广泛发育的古元古代中晚期(2.0~1.8 Ga)岩浆-变质事件与孔兹岩带最为相似, 鄂尔多斯地块北部的变质基底中也存在该岩浆

-变质事件的年龄峰值，但在阴山地块内鲜有报道。因此，阿拉善地块很可能在古元古代中晚期作为孔兹岩带基底的一部分卷入到造山作用中。

5 结语

结合已有数据和笔者近些年在阿拉善地块东部地区获取的研究资料，综合对比分析阿拉善地块和华北克拉通早前寒武纪变质基底的锆石年代学和Hf同位素特征，得出如下认识：

(1) 狼山地区出露的2.64 Ga斜长角闪岩和~2.5 Ga TTG片麻岩，明确指示了阿拉善地块东部地区存在新太古代岩石，它们与阿拉善地块西部北大山地区的~2.5 Ga的TTG片麻岩共同组成了阿拉善地块太古宙变质基底。

(2) 阿拉善地块东部的早前寒武纪变质基底主要记录了3.4~2.6 Ga的陆壳生长、~2.5 Ga TTG岩浆-变质事件、2.3~2.0 Ga岩浆事件、~1.95 Ga和~1.85 Ga岩浆-变质事件，与西阿拉善和华北克拉通变质基底具有相似的地壳演化历史。

(3) 通过对比华北克拉通西部基底发现，阿拉善地块早前寒武纪变质基底与孔兹岩带共同经历了古元古代中晚期(~1.85 Ga和~1.95 Ga)岩浆-变质事件，表明阿拉善地块很可能在古元古代中晚期作为孔兹岩带基底的一部分卷入到了古元古代造山作用中。

谨以此文献给著名变质地质学家、前寒武纪地质学家沈其韩院士百岁寿诞，向先生严谨的治学态度和为地质事业奉献终身的崇高精神致敬，并祝愿先生健康长寿！

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